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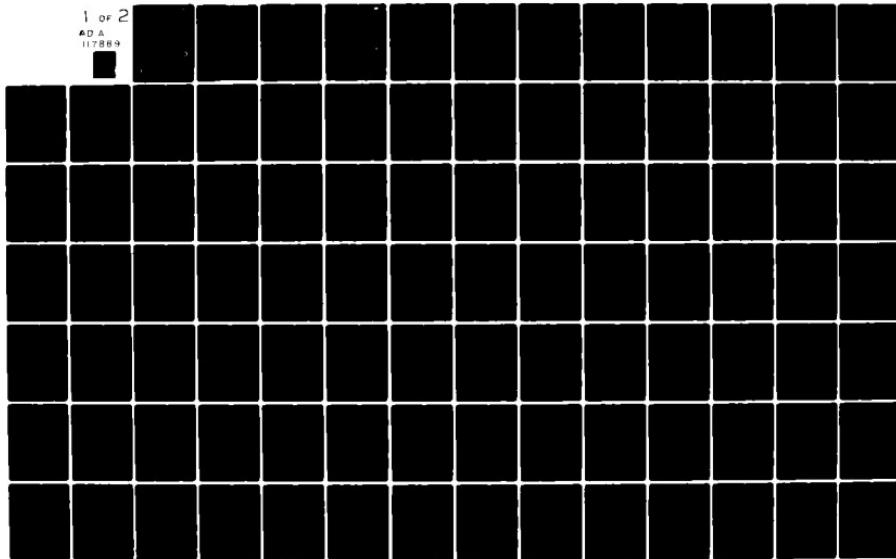
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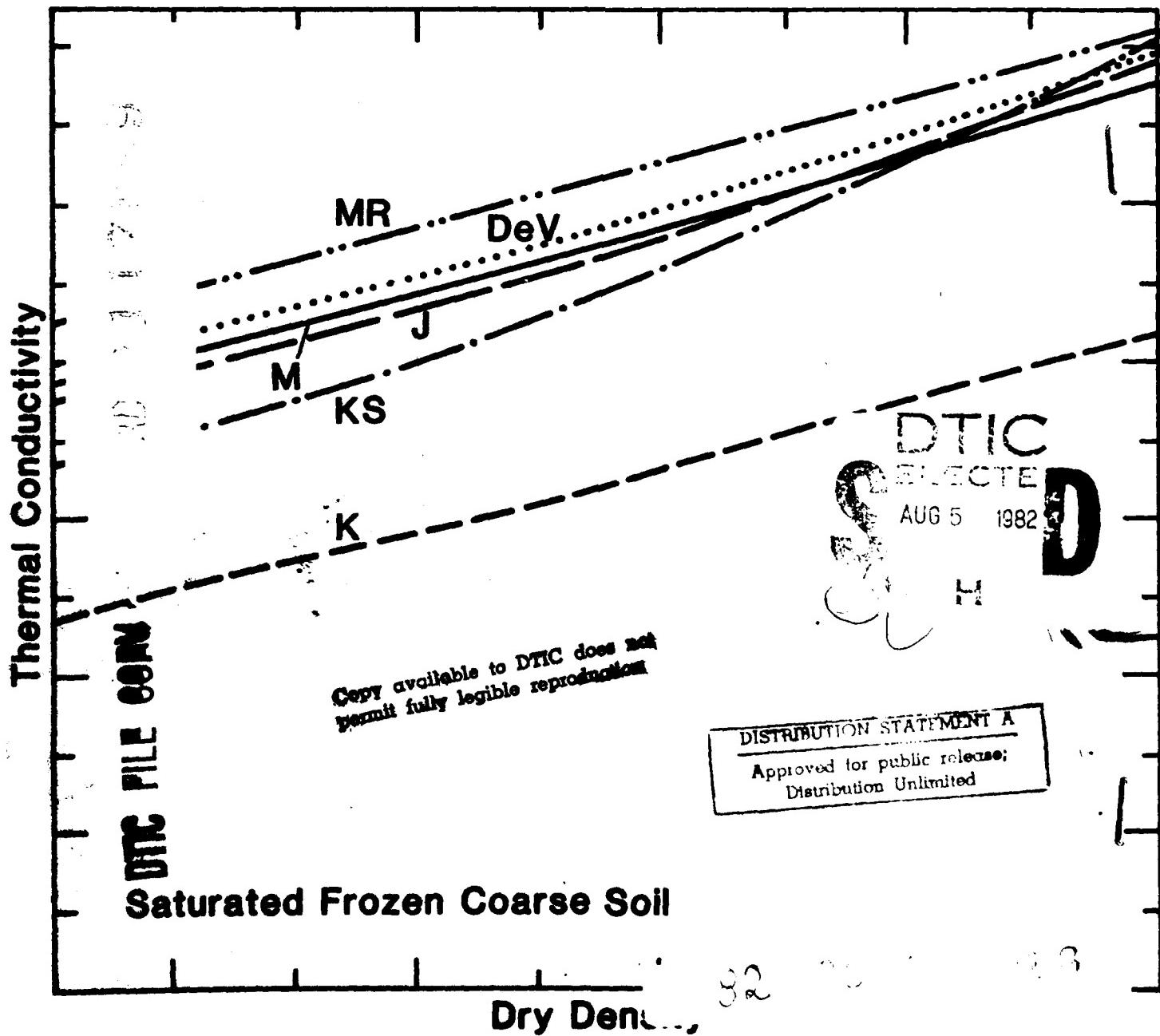


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Cold Regions Research &
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Evaluation of methods for calculating soil thermal conductivity



*For conversion of SI metric units to U.S./
British customary units of measurement
consult ASTM Standard E380, Metric Practice
Guide, published by the American Society
for Testing and Materials, 1916 Race St.,
Philadelphia, Pa. 19103.*

*Cover: Comparison of thermal conductivity
values calculated by various methods (MR—
modified resistor, DeV—De Vries, M—Mick-
ley, J—Johansen, KS—Kunii—Smith, K—
Kersten).*

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Evaluation of methods for calculating soil thermal conductivity

Omar Farouki

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PREFACE

This report was prepared by Dr. Omar T. Farouki, Senior Lecturer, Department of Civil Engineering, The Queen's University of Belfast, Northern Ireland, United Kingdom. It was started at the U.S. Army Cold Regions Research and Engineering Laboratory while the author was on sabbatical leave from the University.

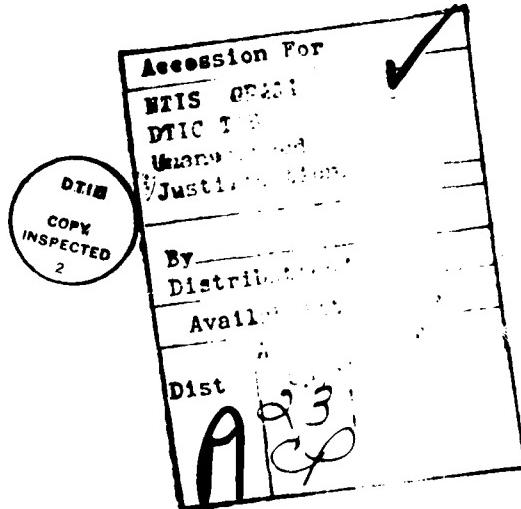
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This report was technically reviewed by Prof. Frederick J. Sanger and Dr. George K. Swinow. The author thanks them for their efforts and for their comments and suggestions.

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NOMENCLATURE

k	soil thermal conductivity	s_γ	sensitivity of thermal conductivity of a soil to variations in dry density at constant moisture content
k_a	thermal conductivity of air	t	temperature
k_f	thermal conductivity of fluid in soil	T	absolute temperature
k_i	thermal conductivity of ice	UWC	unfrozen water content as a fraction by volume of total soil volume
k_q	thermal conductivity of quartz	w	moisture content of soil
k_o	thermal conductivity of soil solids other than quartz	x_a	volume fraction of air in unit soil volume
k_s	thermal conductivity of soil solids	x_f	volume fraction of fluid in unit soil volume
k_w	thermal conductivity of water	x_s	volume fraction of solids in unit soil volume
k_{\parallel}	thermal conductivity of quartz parallel to the c-axis	x_w	volume fraction of water in unit soil volume
k_{\perp}	thermal conductivity of quartz perpendicular to the c-axis	α'	thermal structure value (Smith method)
n	soil porosity (fractional)	γ_d	soil dry density
n_c	volume of series fluid in unit soil volume	γ_k/k_s	sensitivity of thermal conductivity of a soil to variation in soil solids thermal conductivity at constant degree of saturation
p	clay content as a fraction of soil solids		
S_r	degree of saturation of soil (fractional)		
s_w	sensitivity of thermal conductivity of a soil to variations in moisture content at constant dry density		

EVALUATION OF METHODS FOR CALCULATING SOIL THERMAL CONDUCTIVITY

Omar T. Farouki

INTRODUCTION

The U.S. Army Cold Regions Research and Engineering Laboratory monograph entitled *Thermal Properties of Soils* (Farouki 1981) describes the various methods that are available for calculating the thermal conductivity of soils. In chronological order, these are the methods of Smith (1942), Kersten (1949), Mickley (1951), Gemant (1952), De Vries (1952 and 1963), Van Rooyen and Winterkorn (1959), Kunii and Smith (1960), the modified resistor equation (Woodside and Messmer 1961), McGaw (1969) and Johansen (1975).

This report shows in detail the predicted thermal conductivity values given by these methods for appropriate types and conditions of soils. This is done for unfrozen and for frozen soils having a range of moisture contents and dry densities. For each method and each soil condition, the sensitivity of the calculated thermal conductivity value to variations in moisture content and dry density is determined.

These methods are evaluated by comparing their predictions with experimental data from soils of known composition. A computer program is used to calculate the deviations of the predicted values from the experimental values for soils with certain properties. The conditions of applicability and the extent of validity of each of these methods are thus determined. From this follows a recommendation of the method or methods to apply to soils of different types, which may be frozen or unfrozen, range from dry to saturated and have varied dry densities.

Because of the extreme complexity of soils, of their behavior, a semi-empirical approach to calculation of their thermal conductivity is essential and is followed here. Hence the methods mentioned, being either theoretically based with empirical modifications or totally empirical, are tested against experimental data to determine the conditions of their validity in practice and to enhance their trustworthiness under these conditions.

ANALYSIS OF METHODS FOR CALCULATING THERMAL CONDUCTIVITY

Introduction

In order to perform a detailed analysis of the thermal conductivity equations resulting from the various methods, computer programs of these equations were prepared.* The main input parameters were:

1. Specific gravity of the soil solids, taken as 2.70
2. Temperature
3. Thermal conductivity of water k_w and of ice k_i at that temperature
4. Effective thermal conductivity of the air (allowing for moisture migration in the manner suggested by De Vries 1963)
5. Thermal conductivity of the soil solids k_s
6. Moisture content w
7. Dry density γ_d .

*The programming was done by S.A. James Clarke and Albert Smith on the ICL 1906S computer of the Queen's University of Belfast using FORTRAN language.

The last two properties were varied over a wide range so that their influence on the soil thermal conductivity could be determined.

The computer printout provided the thermal conductivity values predicted by the equations for the given input data. Calculations were made for both the unfrozen and the frozen conditions. In the latter case the unfrozen water content was assumed to be zero for this sensitivity analysis. k_s was set at 8.0 W/m K for coarse soils and at 2.0 W/m K for fine soils.

The moisture content w was varied at constant dry density γ_d to determine the sensitivity of the thermal conductivity to variations in w . In another series of calculations, γ_d was varied at constant w to evaluate the influence of γ_d . The sensitivity of the thermal conductivity to k_s was also determined by varying k_s from 2.0 to 8.0 W/m K, keeping both the degree of saturation S_r and the dry density constant.

For the unsaturated frozen condition, only the methods of Kersten, Mickley, De Vries and Johansen could be applied. For the unsaturated unfrozen condition, these were applied together with the Gemant, Van Rooyen and McGaw methods. In the saturated state, frozen or unfrozen, the method of Kunii-Smith and the modified resistor equation could additionally be applied. For dry soils, the methods of Smith, Mickley, De Vries and adjusted De Vries, Van Rooyen, Kunii-Smith, modified resistor, McGaw and Johansen were applicable.

The results of the sensitivity analysis are given and discussed in the next three sections.

Influence of moisture content on thermal conductivity

Thermal conductivity values were calculated at a constant γ_d for moisture contents varying from the dry to the near saturated condition in increments of 5%. This procedure was repeated for different dry densities in increments of 0.1 g/cm³ from a γ_d of 1.1 g/cm³ to one of 2.1 g/cm³. The results for unfrozen coarse soil ($k_s = 8.0$ W/m K) at 4°C are shown in Figure 1 while those for unfrozen fine soil ($k_s = 2.0$ W/m K) are given in Figure 2. Seven methods are applicable to unfrozen soil: Kersten, Mickley, Gemant, De Vries, Van Rooyen, McGaw and Johansen*, except that Kersten and Gemant do not apply for the dry or nearly dry condition. For frozen coarse or fine soil four methods are applicable: Kersten, Mickley, De Vries and Johansen. The resulting curves are shown in Figures 3 and 4.

The sensitivity of the thermal conductivity to $w(%)$ at constant γ_d is given in Tables 1-4 for four representative values of γ_d . These tables give the

absolute sensitivity expressed as the absolute increase in the thermal conductivity per 1% increase in w (designated by s_w). By expressing this value as a percentage of the thermal conductivity (taken in the middle of the associated moisture content range) a value for the "relative sensitivity" can be obtained. Except for Mickley and McGaw, s_w decreases as the moisture content increases.

Figures 5-8 compare the absolute and relative sensitivities for the seven methods at a constant γ_d of 1.4 g/cm³. With one or two exceptions the pattern of relative sensitivities is similar to that of absolute sensitivities. For unfrozen coarse soil (Fig. 5) Johansen, De Vries and Van Rooyen show comparatively high sensitivities from the dry condition to a w of around 7.5%. Beyond this w , all the sensitivities decrease in a roughly similar manner, except for those from Mickley and McGaw which are small and increase slightly.

In the case of unfrozen fine soil (Fig. 6) all the methods, except Mickley and McGaw, give roughly similar trends. The sensitivities generally decrease appreciably as w increases above 7.5%. It may be noted, however, that Johansen gives a maximum s_w at about $w = 7.5\%$, while Van Rooyen gives a maximum at about $w = 15\%$. As for coarse soil Mickley and McGaw give the lowest sensitivities; these do not vary much with the moisture content.

A comparison of the sensitivities for unfrozen coarse soil to the corresponding values for unfrozen fine soil shows that all the methods give an s_w for coarse soil that is higher than for fine soil, with the exception of Kersten which gives a slightly lower value for coarse soil (compare Tables 1 and 2 at a dry density of 1.4 g/cm³).

For frozen soils both Kersten and Johansen give linear relations between the thermal conductivity and w at constant γ_d , implying a constant s_w for each method. In fact, these two methods give nearly the same s_w for frozen fine soil, Johansen giving slightly larger values at the higher dry densities. For frozen coarse soil, however, Johansen gives an appreciably larger s_w , being about 50% more than that from Kersten at a γ_d of 1.4 g/cm³. De Vries shows the highest s_w at low w but this decreases to become the lowest at high w . Mickley, on the other hand, shows a reverse trend (see Fig. 7 and 8).

It may also be noted that the s_w values for frozen fine soil are less than those for frozen coarse soil. The percentage sensitivities, however, do not differ much, except at low w values where Johansen and De Vries show higher values for frozen coarse soil.

Tables 1-4 show that for Kersten, Johansen, De Vries and Mickley, the effect of increasing γ_d is to increase the value of s_w when moisture content is similar.

Kersten, Mickley, De Vries and Johansen are applicable to both the unfrozen and frozen conditions.

*Mention of the method name in the text implies application of the associated equations.

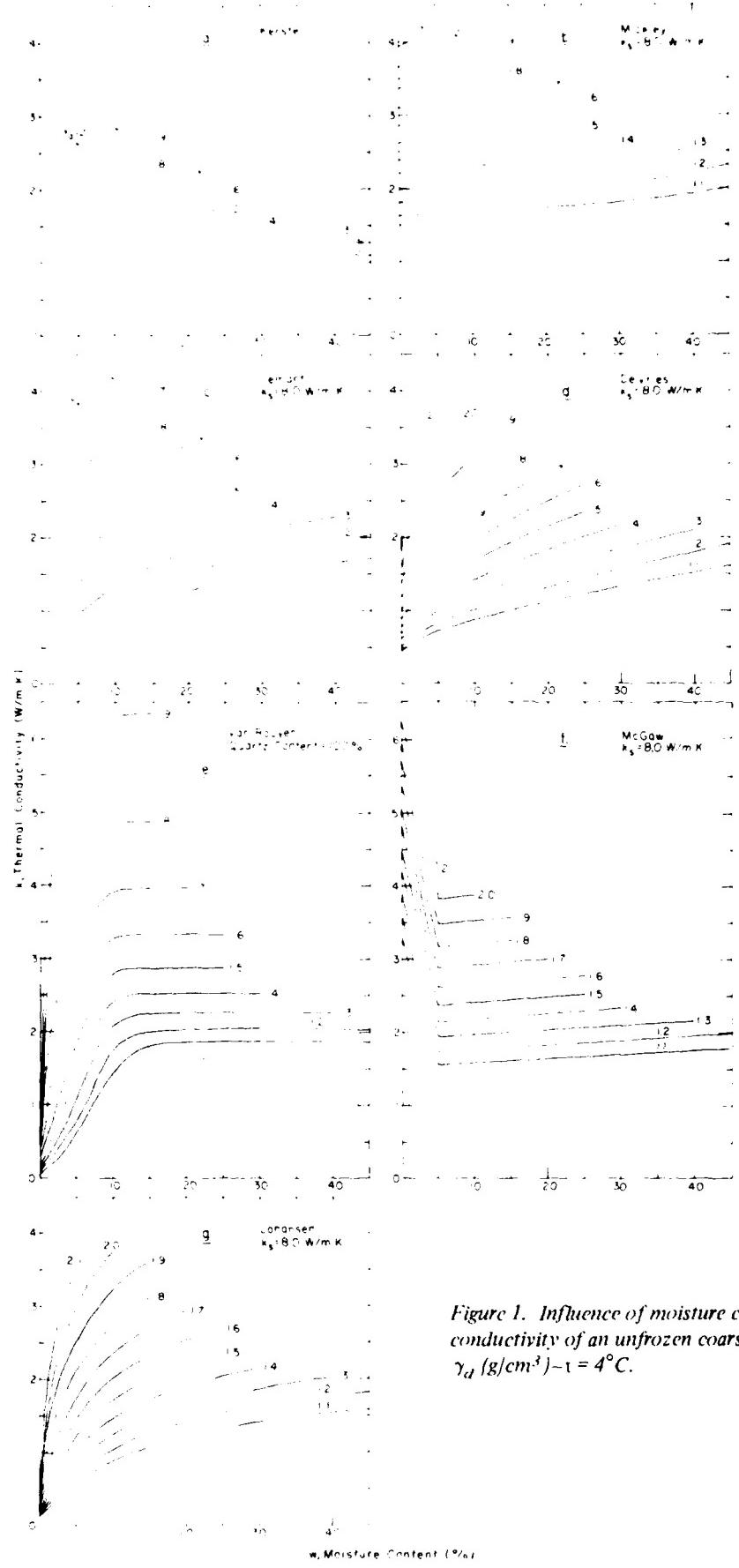


Figure 1. Influence of moisture content on calculated thermal conductivity of an unfrozen coarse soil at constant dry density γ_d (g/cm^3)— $t = 4^\circ\text{C}$.

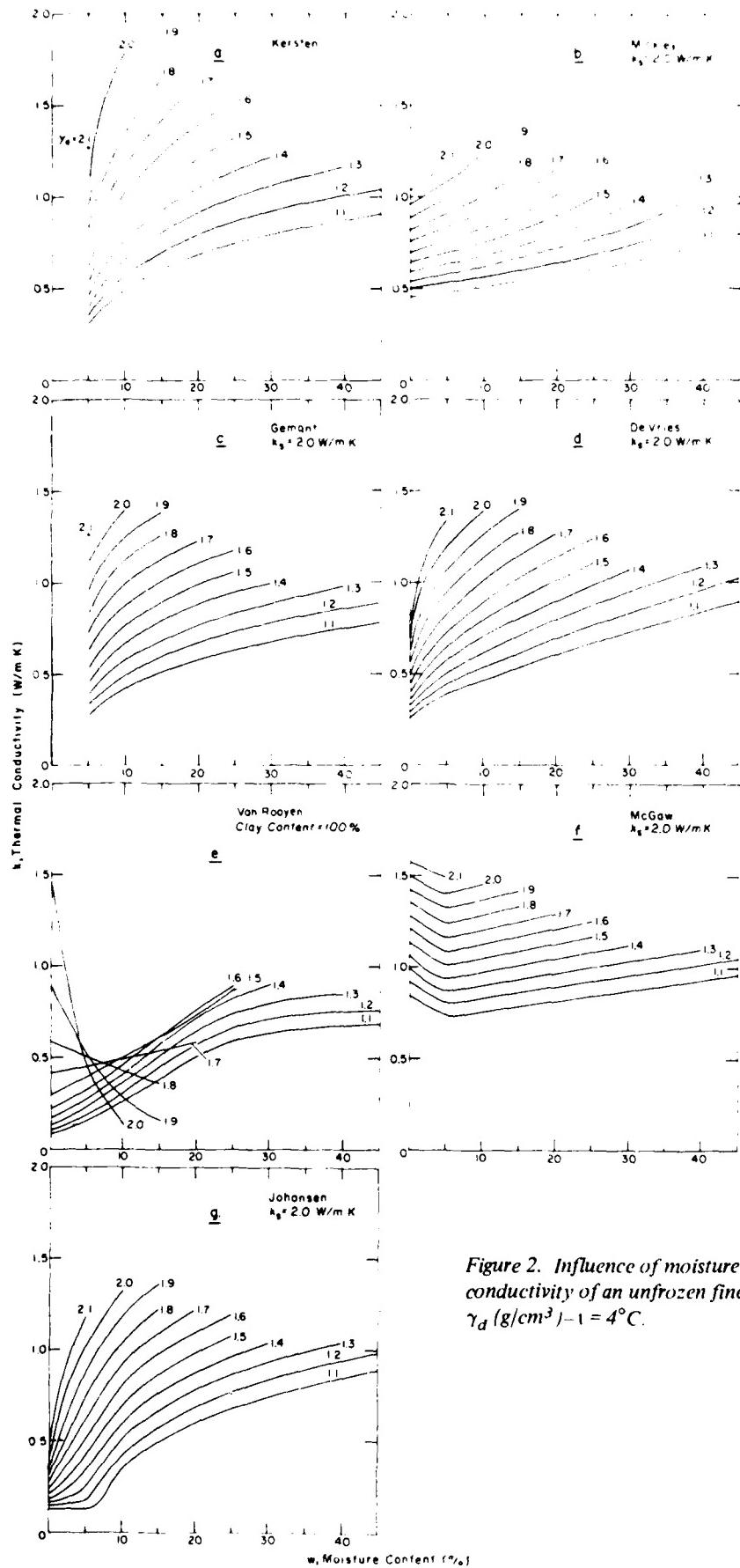


Figure 2. Influence of moisture content on calculated thermal conductivity of an unfrozen fine soil at constant dry density γ_d (g/cm^3) - 1 = 4°C.

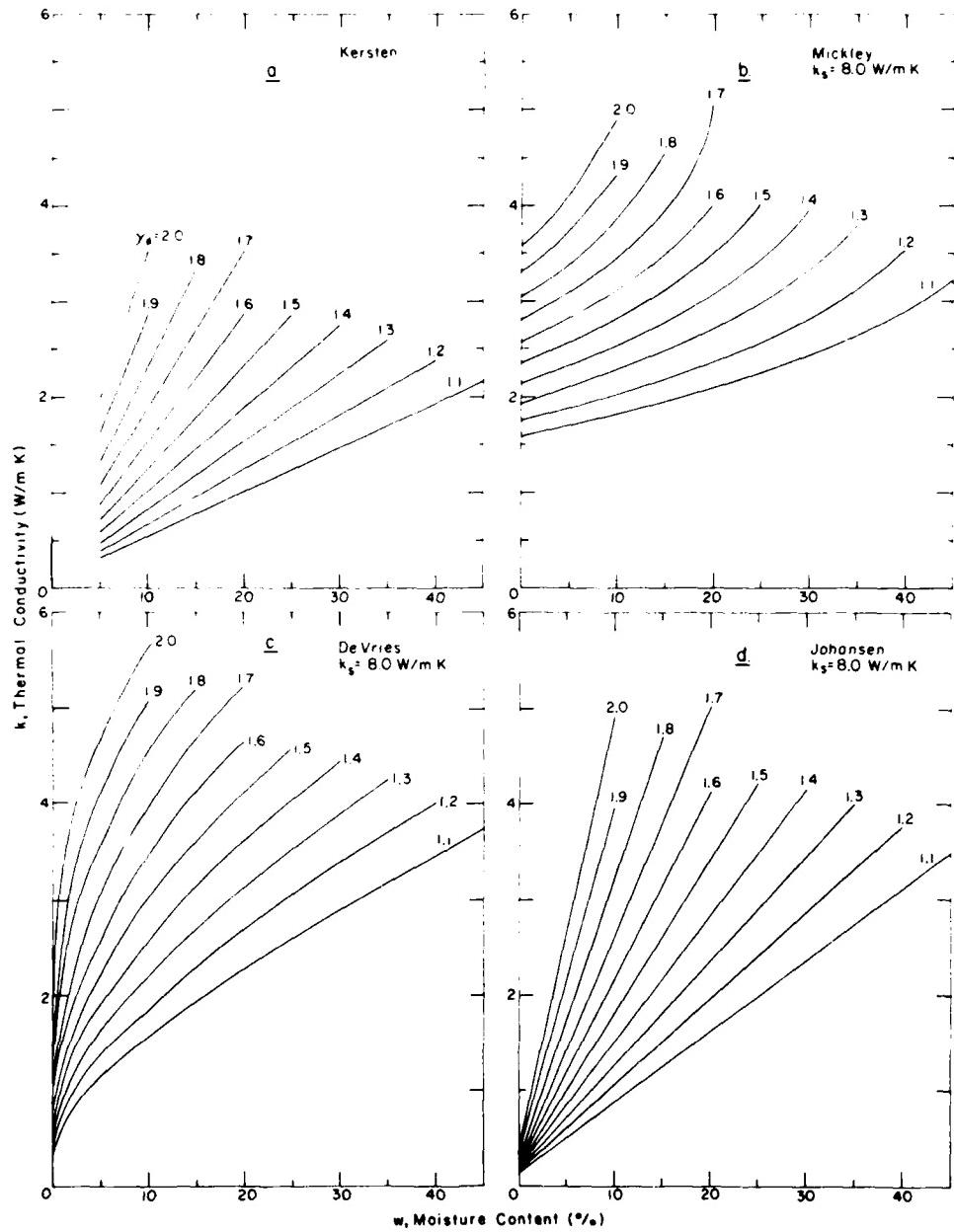


Figure 3. Influence of moisture content on calculated thermal conductivity of a frozen coarse soil at constant dry density γ_d (g/cm^3)— $t = -4^\circ\text{C}$.

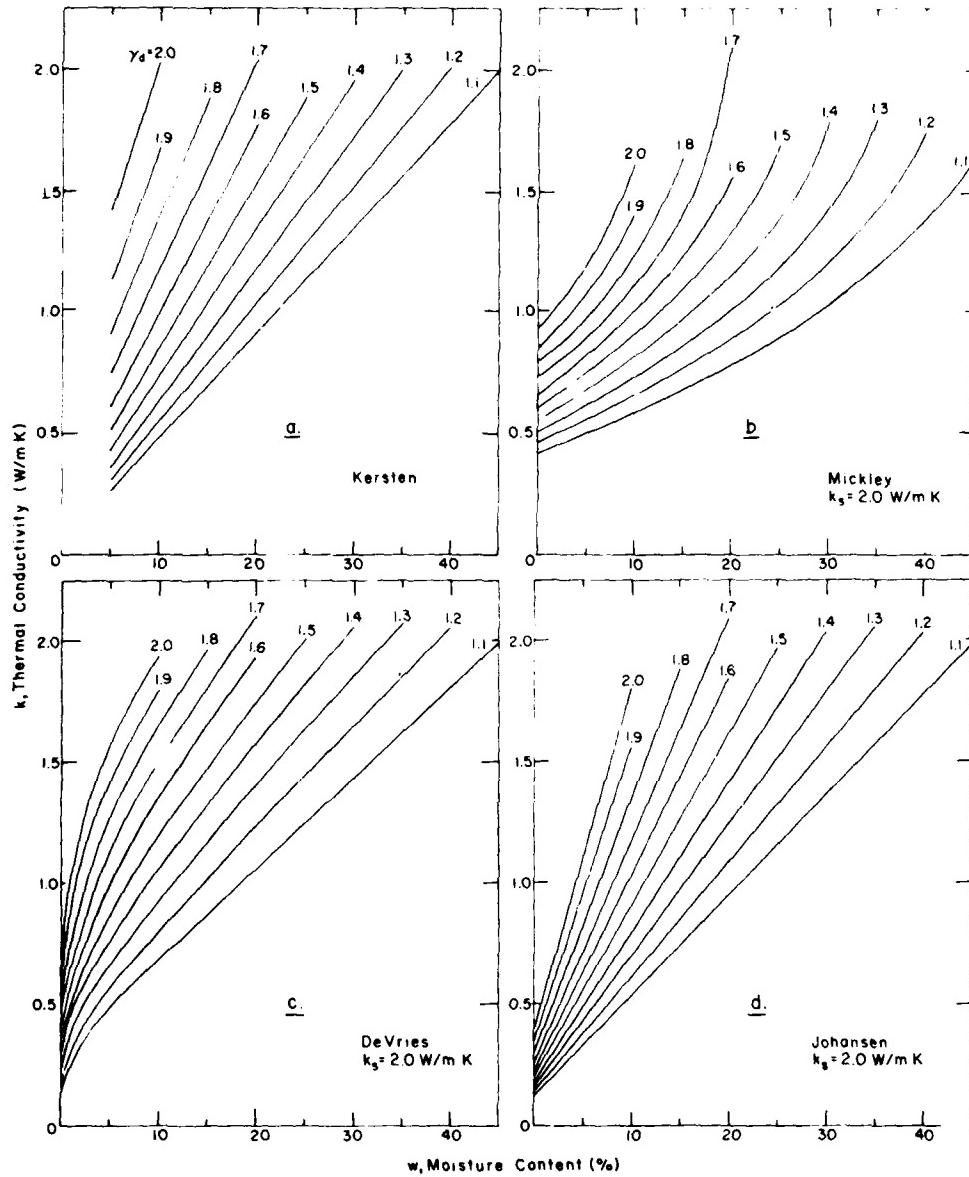


Figure 4. Influence of moisture content on calculated thermal conductivity of a frozen fine soil at constant dry density γ_d (g/cm^3) - $t = -4^\circ\text{C}$.

Table 1. Sensitivity of thermal conductivity k (W/m K) of an unfrozen coarse soil to moisture content w at constant dry density γ_d for different methods of calculating k (s_w is the increase in k per 1% increase in w at constant γ_d).

$k_s = 8.0 \text{ W/m K}$							
	Kersten	Johansen	De Vries	Mickley	McGaw	Gemant	Van Rooyen
$\gamma_d = 1.1 \text{ g/cm}^3$							
s_w for $w = 0$ to 5%	-	0.085	0.052	0.007	negative	-	0.097
k at $w = 2.5\%$	-	0.378	0.610	1.625	-	-	0.308
s_w as % of k	-	22.6	8.5	0.4	-	-	31.6
s_w for $w = 5$ to 15%	0.023	0.051	0.027	0.007	0.006	0.051	0.123
k at $w = 10\%$	0.771	0.881	0.865	1.665	1.586	0.950	1.482
s_w as % of k	3.0	5.8	3.1	0.4	0.4	5.4	8.3
s_w for $w = 15$ to 45%	0.008	0.017	0.021	0.012	0.007	0.022	0.002
k at $w = 30\%$	1.005	1.395	1.332	1.823	1.700	1.497	1.860
s_w as % of k	0.8	1.2	1.5	0.7	0.4	1.5	0.1
$\gamma_d = 1.4 \text{ g/cm}^3$							
s_w for $w = 0$ to 5%	-	0.170	0.094	0.011	-	-	0.256
k at $w = 2.5\%$	-	0.683	0.950	2.213	-	-	0.788
s_w as % of k	-	24.9	9.9	0.5	-	-	32.5
s_w for $w = 5$ to 10%	0.046	0.086	0.053	0.012	0.007	0.096	0.192
k at $w = 7.5\%$	1.078	1.280	1.300	2.263	2.163	1.383	2.070
s_w as % of k	4.2	6.8	4.1	0.5	0.3	7.0	9.3
s_w for $w = 10$ to 20%	0.023	0.043	0.042	0.014	0.007	0.050	0.010
k at $w = 15\%$	1.320	1.721	1.649	2.344	2.218	1.891	2.525
s_w as % of k	1.7	2.5	2.5	0.6	0.3	2.6	0.4
s_w for $w = 20$ to 30%	0.013	0.025	0.033	0.019	0.007	0.031	0.000
k at $w = 25\%$	1.487	2.040	2.008	2.500	2.291	2.275	2.532
s_w as % of k	0.9	1.2	1.6	0.8	0.3	1.4	0.0
$\gamma_d = 1.7 \text{ g/cm}^3$							
s_w for $w = 0$ to 5%	-	0.302	0.160	0.018	-	-	0.565
k at $w = 2.5\%$	-	1.262	1.475	2.870	-	-	2.612
s_w as % of k	-	23.9	10.8	0.6	-	-	21.6
s_w for $w = 5$ to 20%	0.047	0.077	0.072	0.027	0.009	0.091	0.264
k at $w = 10\%$	1.827	2.357	2.297	3.026	2.921	2.604	3.939
s_w as % of k	2.6	3.3	3.1	0.9	0.3	3.5	6.7
$\gamma_d = 2.0 \text{ g/cm}^3$							
s_w for $w = 0$ to 5%	-	0.513	0.298	0.033	-	-	-
k at $w = 2.5\%$	-	2.308	2.480	3.720	-	-	-
s_w as % of k	-	22.2	12.0	3.9	-	-	-
s_w for $w = 5$ to 10%	0.108	0.152	0.119	0.044	0.010	0.181	-
k at $w = 5\%$	2.273	2.984	3.030	3.797	3.829	3.222	-
s_w as % of k	4.7	5.1	3.9	1.2	0.3	5.6	-

Table 2. Sensitivity of thermal conductivity k (W/m K) of an unfrozen fine soil to moisture content w at constant dry density γ_d for different methods of calculating k (s_w is the increase in k per 1% increase in w at constant γ_d).

$k_s \approx 2.0 \text{ W/m K}$							
<i>Kersten</i>	<i>Johansen</i>	<i>De Vries</i>	<i>Mickley</i>	<i>McGaw</i>	<i>Germann</i>	<i>Van Rooyen</i>	
$\gamma_d = 1.1 \text{ g/cm}^3$							
s_w for $w = 0$ to 5%	-	0.000	0.026	0.006	-0.021	-	0.015
k at $w = 2.5\%$	-	0.131	0.333	0.460	0.777	-	0.116
s_w as % of k	-	0.000	7.7	1.3	-2.7	-	12.6
s_w for $w = 5$ to 10%	0.038	0.044	0.014	0.006	0.004	0.030	0.022
k at $w = 7.5\%$	0.405	0.200	0.435	0.490	0.740	0.362	0.210
s_w as % of k	9.4	22.0	3.3	1.2	0.5	8.2	10.3
s_w for $w = 10$ to 20%	0.019	0.025	0.014	0.007	0.006	0.015	0.024
k at $w = 15\%$	0.602	0.494	0.534	0.536	0.786	0.512	0.392
s_w as % of k	3.2	5.0	2.6	1.2	0.7	2.9	6.2
s_w for $w = 20$ to 45%	0.009	0.012	0.012	0.056	0.006	0.008	0.007
k at $w = 30\%$	0.792	0.740	0.724	0.645	0.871	0.676	0.634
s_w as % of k	1.1	1.6	1.6	8.7	0.6	1.2	1.1
$\gamma_d = 1.4 \text{ g/cm}^3$							
s_w for $w = 0$ to 5%	-	0.029	0.043	0.009	-0.025	-	0.022
k at $w = 2.5\%$	-	0.240	0.486	0.610	0.978	-	0.219
s_w as % of k	-	12.2	8.8	1.5	-2.6	-	9.9
s_w for $w = 5$ to 10%	0.058	0.054	0.025	0.010	0.007	0.041	0.028
k at $w = 7.5\%$	0.636	0.478	0.645	0.659	0.955	0.579	0.342
s_w as % of k	9.2	11.4	3.8	1.5	0.8	7.1	8.3
s_w for $w = 10$ to 20%	0.029	0.027	0.020	0.012	0.007	0.021	0.030
k at $w = 15\%$	0.926	0.765	0.807	0.740	1.010	0.787	0.576
s_w as % of k	3.2	3.6	2.5	1.6	0.7	2.6	5.2
s_w for $w = 20$ to 30%	0.017	0.016	0.017	0.016	0.004	0.012	0.018
k at $w = 25\%$	1.142	0.966	0.987	0.873	1.082	0.942	0.827
s_w as % of k	1.5	1.6	1.7	1.9	0.3	1.3	2.2
$\gamma_d = 1.7 \text{ g/cm}^3$							
s_w for $w = 0$ to 5%	-	0.071	0.064	0.015	-0.023	-	0.008
k at $w = 2.5\%$	-	0.455	0.698	0.796	1.205	-	0.435
s_w as % of k	-	15.6	9.2	1.8	-1.9	-	1.8
s_w for $w = 5$ to 10%	0.090	0.059	0.038	0.017	0.009	0.005	0.099
k at $w = 7.5\%$	0.972	0.785	0.922	0.872	1.182	0.876	0.473
s_w as % of k	9.2	7.5	4.1	1.9	0.7	0.5	20.9
s_w for $w = 10$ to 20%	0.045	0.030	0.025	0.025	0.009	0.024	0.009
k at $w = 15\%$	1.426	1.096	1.146	1.016	1.249	1.131	0.540
s_w as % of k	3.2	2.7	2.2	2.4	0.7	2.1	1.7
$\gamma_d = 2.0 \text{ g/cm}^3$							
s_w for $w = 0$ to 5%	-	0.120	0.098	0.024	-0.018	-	-0.202
k at $w = 2.5\%$	-	0.800	1.057	1.020	1.440	-	0.907
s_w as % of k	-	15.0	9.3	2.4	-1.3	-	-22.3
s_w for $w = 5$ to 10%	0.138	0.062	0.039	0.033	0.010	0.057	-0.063
k at $w = 7.5\%$	1.563	1.180	1.296	1.165	1.430	1.270	0.265
s_w as % of k	8.9	5.2	3.0	2.8	0.7	4.5	-23.8

Table 3. Sensitivity of thermal conductivity k (W/m K) of a frozen coarse soil to moisture content w at constant dry density γ_d for different methods of calculating k (s_w is the increase in k per 1% increase in w at constant γ_d).

$k_s = 8.0 \text{ W/m K}$				
	Kerten	Johansen	De Fries	Mickler
$\gamma_d = 1.1 \text{ g/cm}^3$				
s_w for $w = 0$ to 5%	-	0.074	0.162	0.024
k at $w = 2.5\%$	-	0.305	0.863	1.630
s_w as % of k	-	24.3	18.7	1.5
s_w for $w = 5$ to 15%	0.046	0.074	0.076	0.026
k at $w = 10\%$	0.549	0.873	1.541	1.813
s_w as % of k	8.4	8.5	4.9	1.5
s_w for $w = 15$ to 25%	0.046	0.074	0.067	0.031
k at $w = 20\%$	1.013	1.616	2.273	2.096
s_w as % of k	4.6	4.6	2.9	1.5
s_w for $w = 25$ to 45%	0.046	0.074	0.057	0.048
k at $w = 35\%$	1.709	2.730	3.179	2.637
s_w as % of k	2.7	2.7	1.8	1.8
$\gamma_d = 1.4 \text{ g/cm}^3$				
s_w for $w = 0$ to 5%	-	0.133	0.278	0.039
k at $w = 2.5\%$	-	0.505	1.488	2.453
s_w as % of k	-	26.2	18.7	1.6
s_w for $w = 5$ to 15%	0.087	0.133	0.123	0.045
k at $w = 10\%$	1.021	1.513	2.591	2.540
s_w as % of k	8.5	8.8	4.7	1.8
s_w for $w = 15$ to 25%	0.087	0.133	0.089	0.063
k at $w = 20\%$	1.891	2.839	3.617	3.059
s_w as % of k	4.6	4.7	2.1	2.1
s_w for $w = 25$ to 30%	0.087	0.133	0.082	0.110
k at $w = 27.5\%$	2.538	3.825	4.238	3.655
s_w as % of k	3.4	3.5	1.9	3.0
$\gamma_d = 1.7 \text{ g/cm}^3$				
s_w for $w = 0$ to 5%	-	0.238	0.452	0.063
k at $w = 2.5\%$	-	0.858	2.463	2.950
s_w as % of k	-	27.7	18.4	2.1
s_w for $w = 5$ to 15%	0.163	0.238	0.159	0.085
k at $w = 10\%$	1.897	2.651	3.996	3.475
s_w as % of k	8.6	9.0	4.0	2.4
s_w for $w = 15$ to 20%	0.163	0.237	0.114	0.220
k at $w = 17.5\%$	3.108	4.418	4.968	4.318
s_w as % of k	5.2	5.4	2.3	5.1
$\gamma_d = 2.0 \text{ g/cm}^3$				
s_w for $w = 0$ to 5%	-	0.451	0.689	0.107
k at $w = 2.5\%$	-	0.970	4.063	3.830
s_w as % of k	-	46.5	16.9	2.8
s_w for $w = 5$ to 10%	0.307	0.450	0.193	0.159
k at $w = 7.5\%$	2.750	3.750	5.238	4.488
s_w as % of k	11.1	12.0	3.7	3.6

Table 4. Sensitivity of thermal conductivity k (W/m K) of a frozen fine soil to moisture content w at constant dry density γ_d for different methods of calculating k (s_w is the increase in k per 1% increase in w at constant γ_d).

$$k_s = 2.0 \text{ W/m K}$$

	<i>Kersten</i>	<i>Johansen</i>	<i>De Vries</i>	<i>Mickley</i>
$\gamma_d = 1.1 \text{ g/cm}^3$				
s_w for $w = 0$ to 5%	-	0.041	0.068	0.016
k at $w = 2.5\%$	-	0.230	0.360	0.450
s_w as % of k	-	17.8	18.8	3.6
s_w for $w = 5$ to 15%	0.043	0.041	0.039	0.018
k at $w = 10\%$	0.481	0.541	0.684	0.579
s_w as % of k	9.0	7.6	5.6	3.2
s_w for $w = 15$ to 25%	0.044	0.041	0.038	0.022
k at $w = 20\%$	0.915	0.952	1.077	0.779
s_w as % of k	4.8	4.3	3.5	2.8
s_w for $w = 25$ to 45%	0.043	0.041	0.036	0.036
k at $w = 35\%$	1.567	1.569	1.629	1.181
s_w as % of k	2.8	2.6	2.2	3.1
$\gamma_d = 1.4 \text{ g/cm}^3$				
s_w for $w = 0$ to 5%	-	0.062	0.109	0.024
k at $w = 2.5\%$	-	0.337	0.605	0.609
s_w as % of k	-	18.3	18.0	3.9
s_w for $w = 5$ to 15%	0.061	0.062	0.057	0.029
k at $w = 10\%$	0.734	0.803	1.066	0.808
s_w as % of k	8.4	7.7	5.3	3.6
s_w for $w = 15$ to 25%	0.061	0.062	0.048	0.043
k at $w = 20\%$	1.347	1.419	1.584	1.152
s_w as % of k	4.6	4.3	3.0	3.7
s_w for $w = 25$ to 30%	0.061	0.062	0.047	0.078
k at $w = 27.5\%$	1.805	1.880	1.936	1.550
s_w as % of k	3.4	3.3	2.4	5.1
$\gamma_d = 1.7 \text{ g/cm}^3$				
s_w for $w = 0$ to 5%	-	0.091	0.161	0.035
k at $w = 2.5\%$	-	0.495	0.870	0.800
s_w as % of k	-	18.5	18.5	4.4
s_w for $w = 5$ to 15%	0.087	0.091	0.068	0.050
k at $w = 10\%$	1.177	1.187	1.507	1.110
s_w as % of k	7.4	7.7	4.5	4.5
s_w for $w = 15$ to 20%	0.086	0.091	0.058	0.140
k at $w = 17.5\%$	1.820	1.865	1.951	1.665
s_w as % of k	4.7	4.9	3.0	8.4
$\gamma_d = 2.0 \text{ g/cm}^3$				
s_w for $w = 0$ to 5%	-	0.139	0.216	0.053
k at $w = 2.5\%$	-	0.756	1.285	1.041
s_w as % of k	-	18.3	16.8	5.1
s_w for $w = 5$ to 10%	0.122	0.139	0.076	0.085
k at $w = 7.5\%$	1.715	1.450	1.765	1.375
s_w as % of k	7.1	9.6	4.3	6.2

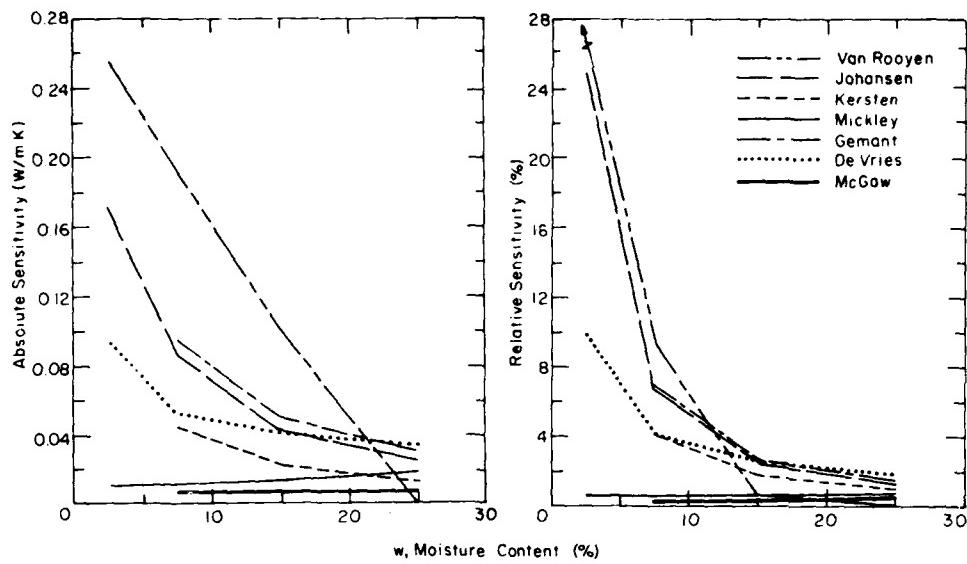


Figure 5. Absolute and relative sensitivities of calculated thermal conductivity of an unfrozen coarse soil vs moisture content ($k_s = 8.0 \text{ W/m K}$, $\gamma_d = 1.4 \text{ g/cm}^3$). Absolute sensitivity is the change in thermal conductivity (W/m K) due to 1% change in moisture content. Relative sensitivity is the absolute sensitivity expressed as a percentage of the thermal conductivity.

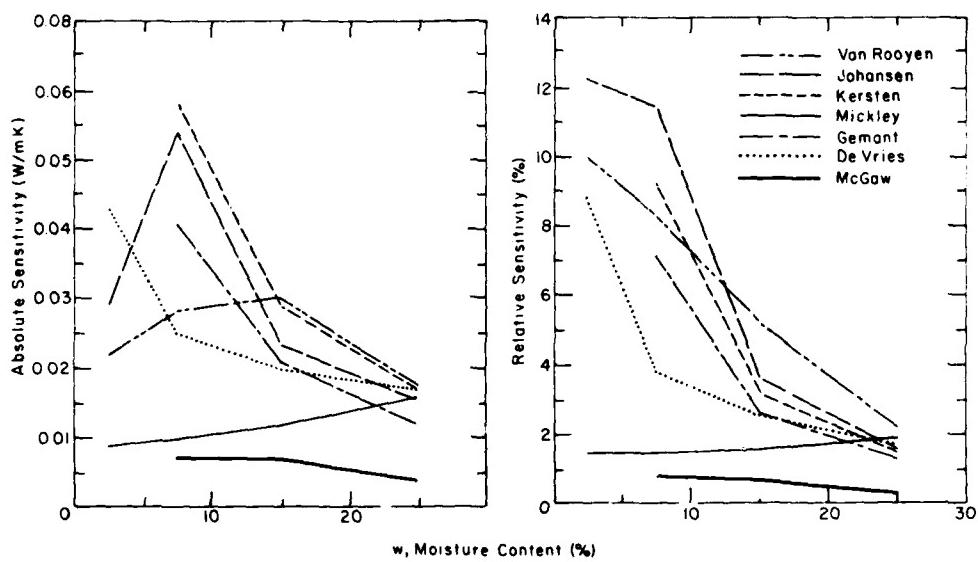


Figure 6. Absolute and relative sensitivities of calculated thermal conductivity of an unfrozen fine soil vs moisture content ($k_s = 2.0 \text{ W/m K}$, $\gamma_d = 1.4 \text{ g/cm}^3$). Absolute sensitivity is the change in thermal conductivity (W/m K) due to 1% change in moisture content. Relative sensitivity is the absolute sensitivity expressed as a percentage of the thermal conductivity.

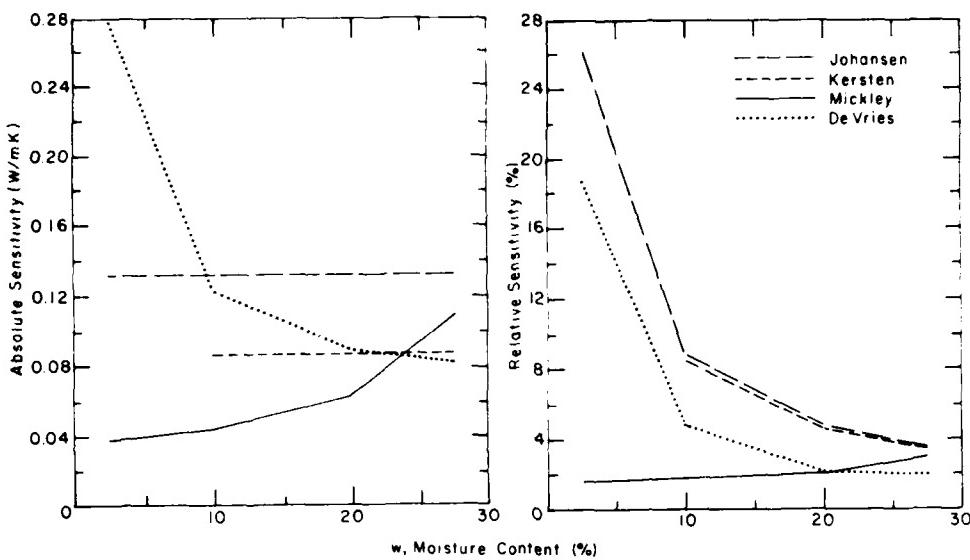


Figure 7. Absolute and relative sensitivities of calculated thermal conductivity of a frozen coarse soil vs moisture content ($k_s = 8.0 \text{ W/m K}$, $\gamma_d = 1.4 \text{ g/cm}^3$). Absolute sensitivity is the change in thermal conductivity (W/m K) due to 1% change in moisture content. Relative sensitivity is the absolute sensitivity expressed as a percentage of the thermal conductivity.

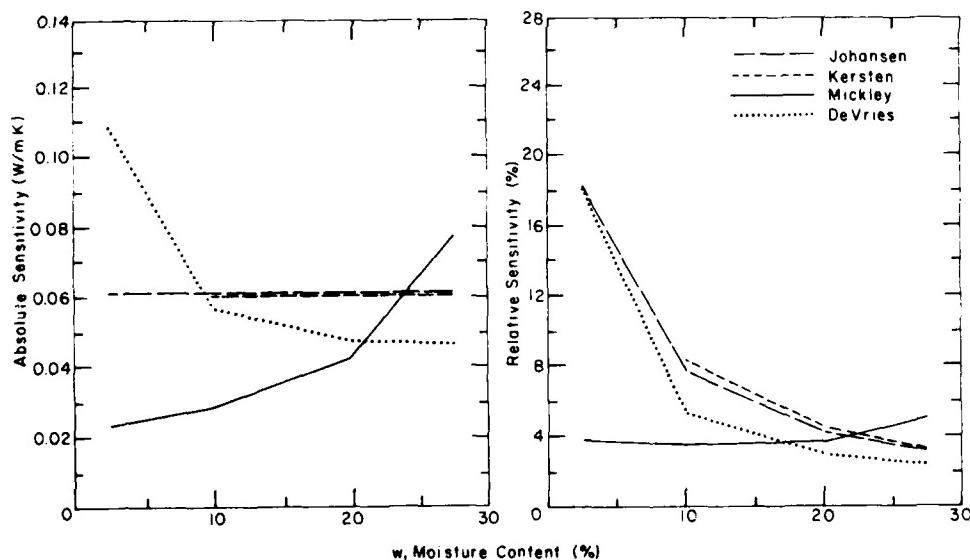


Figure 8. Absolute and relative sensitivities of calculated thermal conductivity of a frozen fine soil vs moisture content ($k_s = 2.0 \text{ W/m K}$, $\gamma_d = 1.4 \text{ g/cm}^3$). Absolute sensitivity is the change in thermal conductivity (W/m K) due to 1% change in moisture content. Relative sensitivity is the absolute sensitivity expressed as a percentage of the thermal conductivity.

From a comparison of the results for these conditions, it is evident that each of these common methods gives an s_w value which is appreciably greater for frozen soil than for unfrozen soil. This may be expected because ice has a much higher thermal conductivity than water. At low value of w (0-5% range), Johansen is an exception, giving a higher s_w for the unfrozen condition than for the frozen state.

Influence of dry density on thermal conductivity

To determine the effect of γ_d , thermal conductivity values were calculated at a constant w for values of γ_d varying from 1.1 to 2.0 g/cm³ in increments of 0.1 g/cm³. This was done for w values from dry to near saturation. At the higher w values, saturation corresponds to lower values of γ_d . As in the previous section, results were obtained from seven methods for the unfrozen state (Fig. 9 and 10) and from four methods for the frozen state (Fig. 11 and 12).

Van Rooyen clearly shows incorrect trends at high values of γ_d or w . Also, both Mickley and McGaw give values for the dry thermal conductivity that are obviously much too high. In Mickley's case, this was expected by him, while for McGaw, the interfacial efficiency ϵ was assumed to be 1.0 in the calculations which is too high for the dry or nearly dry condition.

The sensitivity of the thermal conductivity to γ_d (at constant w) is given in Tables 5-8 for several representative values of w . These tables give the absolute sensitivity s_γ expressed as the increase in thermal conductivity per 0.1 g/cm³ increase in γ_d at constant w . This increase has also been expressed in relative terms as a percentage of the thermal conductivity value in the middle of the dry density range.

Van Rooyen is considerably out of step with the others for the unfrozen soils (Fig. 9 and 10). For these soils the other six methods give an increased s_γ as γ_d increases at constant w . It is also evident that s_γ increases as w increases for a given γ_d range; however, the values from Mickley and McGaw do not vary much.

For unfrozen coarse soil Kersten gives the lowest s_γ throughout (Table 5). The relative sensitivity given by Kersten is constant at about 14.5% over

the whole range of γ_d and w . On the other hand for unfrozen fine soil Kersten tends to give the highest s_γ at the higher values of γ_d (Table 6). The values of s_γ for unfrozen fine soil are lower than the corresponding values for unfrozen coarse soil, as expected, because of the higher k_s value for the latter.

Tables 7 and 8 for frozen soils show the marked increase in s_γ caused by γ_d increasing at constant w . As for unfrozen soils, increased w gives increased sensitivity over a similar range of γ_d . For frozen coarse soil Kersten generally gives the lowest s_γ while Johansen and Mickley give the largest at high values of w . For frozen fine soil, Kersten and Johansen give s_γ values near each other at low γ_d values, but Kersten tends to give higher values in the higher γ_d range.

Because the degree of saturation S_f may be more important than the absolute value of w , especially in affecting moisture migration, thermal conductivity values have been calculated at a constant S_f value but with varying γ_d . The curves corresponding to several S_f values are shown in Figures 13-16. Along a constant S_f curve, varying γ_d implies a varying w ; the effect of a varying w is a contributing factor to changes in thermal conductivity.

Apart from Van Rooyen, these curves (at constant S_f) for the unfrozen soils show a more or less constant rate of increase in the thermal conductivity with increasing γ_d . This implies that s_γ (at constant S_f) is approximately constant. For a given method the curves corresponding to different S_f values run roughly parallel to each other. The same trends are indicated for the four methods applicable to the frozen coarse soil (Fig. 15). With regard to the frozen fine soil, however, the trends shown by these methods are somewhat different (Fig. 16).

The effect of S_f is interesting to note. For the unfrozen soils, the curves given by Kersten, Geman, De Vries and Johansen show that the sensitivity of the thermal conductivity to S_f at constant γ_d decreases as S_f increases (i.e. the curves become closer together). The opposite trend is shown by Mickley, but this appears to be physically incorrect. For the frozen soils, Kersten, De Vries and Johansen give a nearly constant s_w as S_f increases at constant dry density.

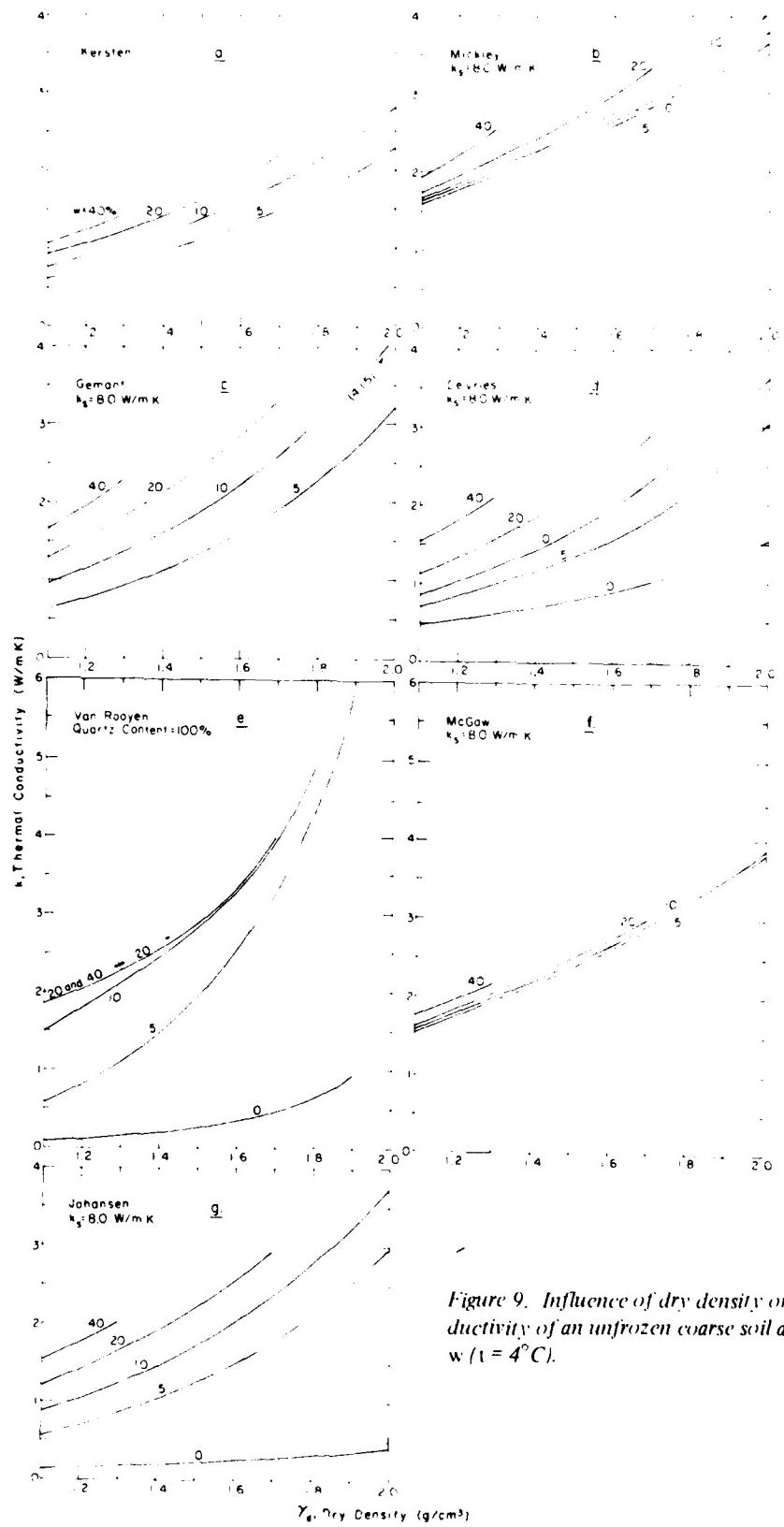


Figure 9. Influence of dry density on calculated thermal conductivity of an unfrozen coarse soil at constant moisture content w ($t = 4^\circ\text{C}$).

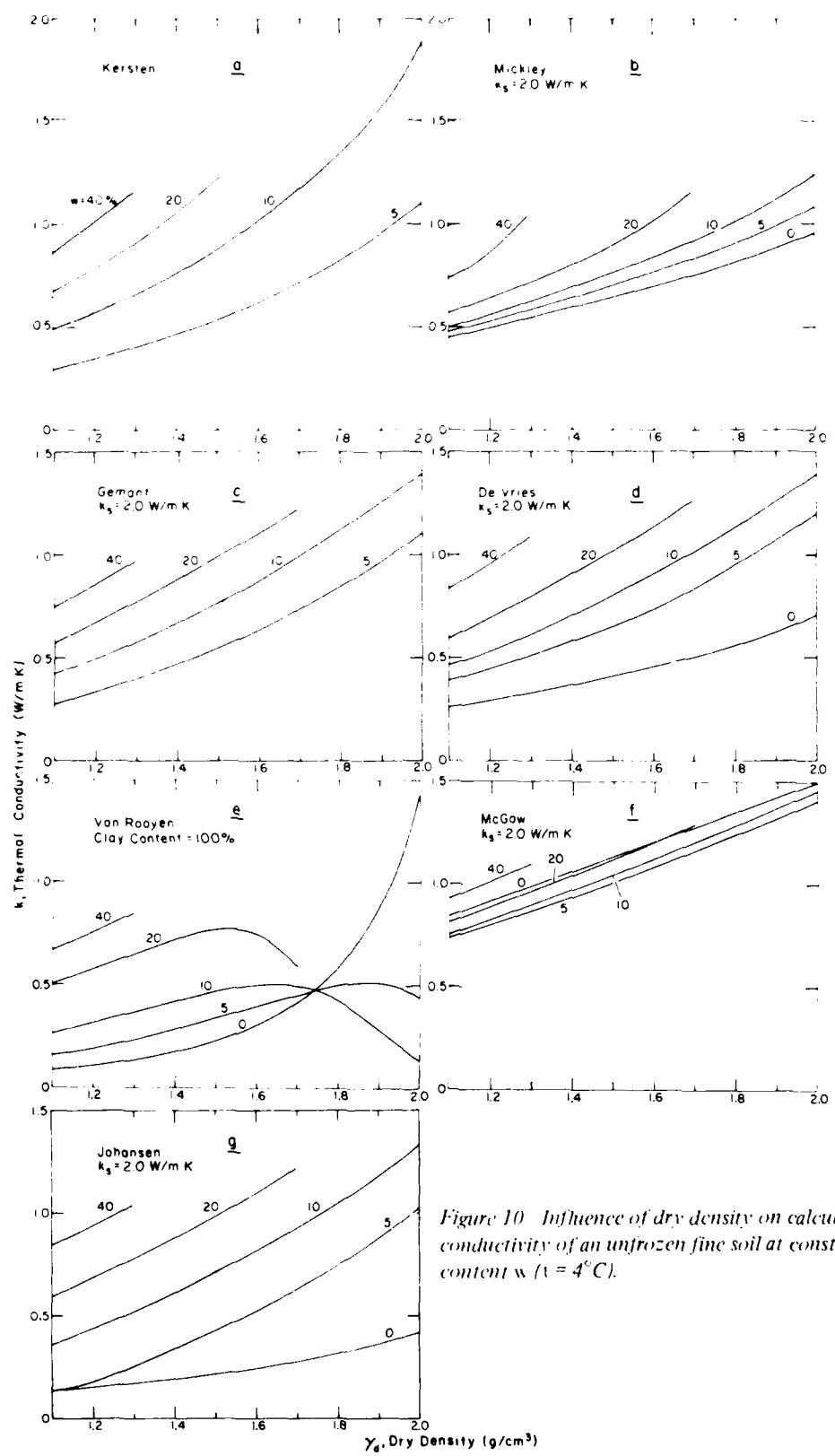


Figure 10. Influence of dry density on calculated thermal conductivity of an unfrozen fine soil at constant moisture content w ($t = 4^\circ\text{C}$).

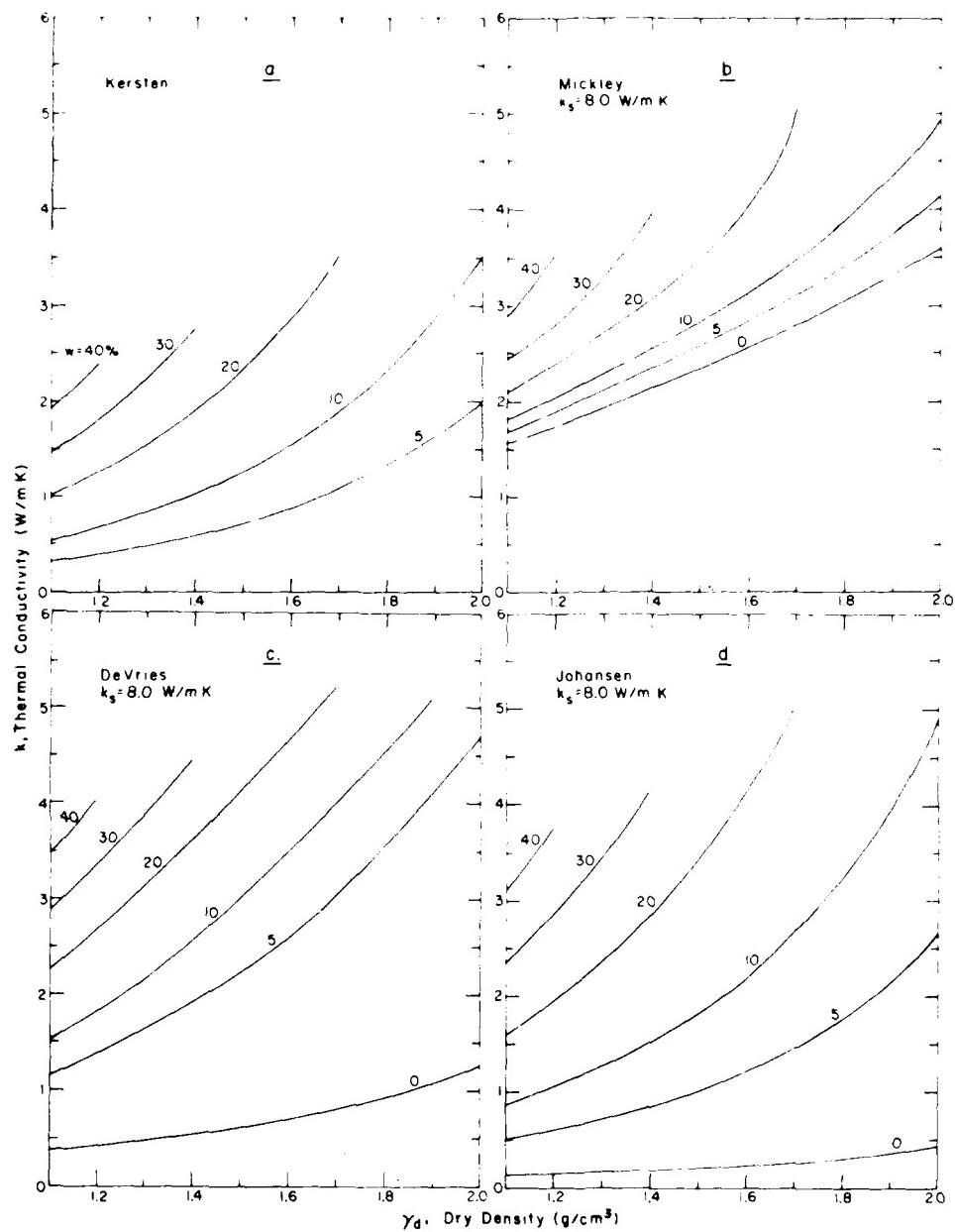


Figure 11. Influence of dry density on calculated thermal conductivity of a frozen coarse soil at constant moisture content w ($t = -4^\circ\text{C}$).

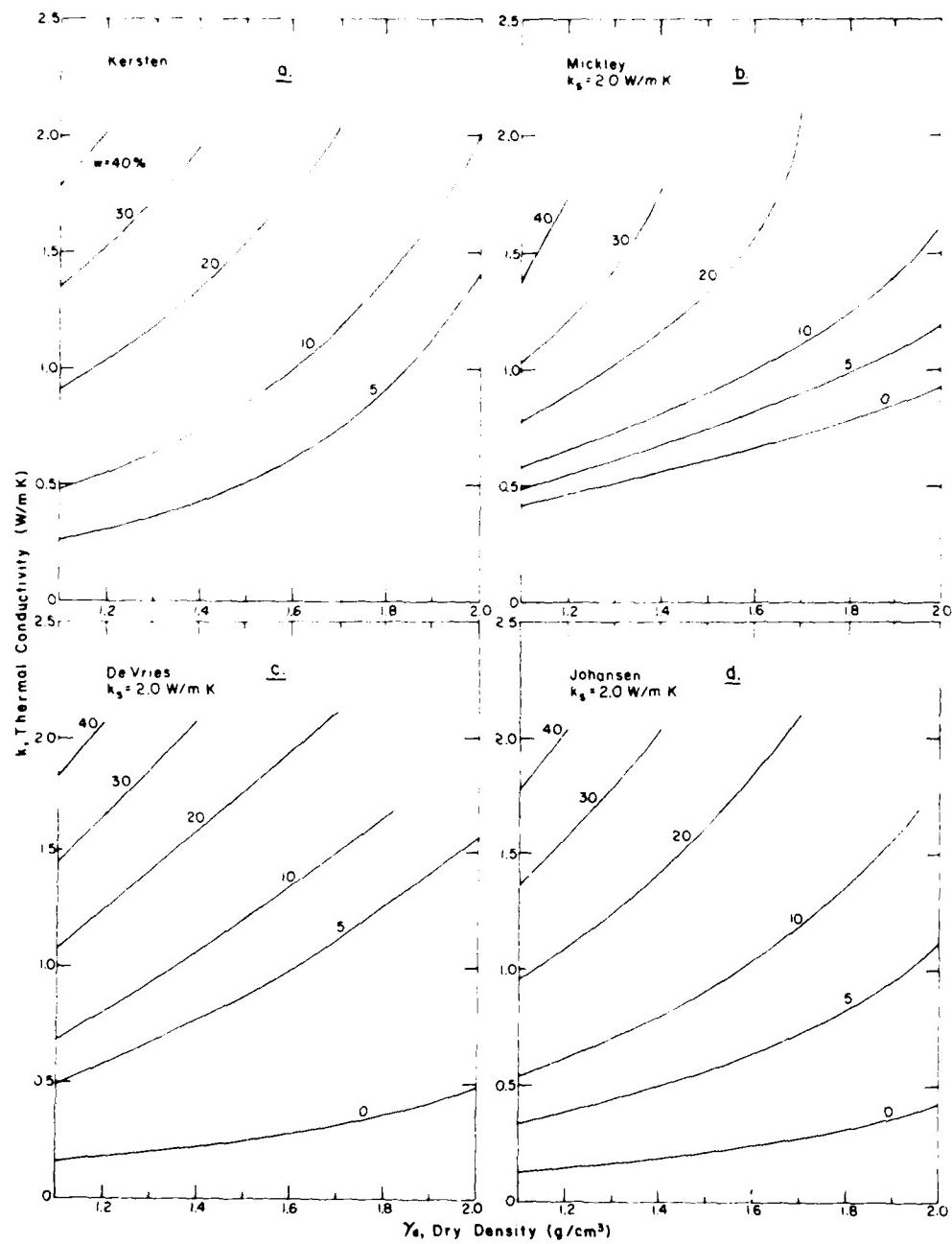


Figure 12. Influence of dry density on calculated thermal conductivity of a frozen fine soil at constant moisture content w ($t = -4^\circ\text{C}$).

Table 5. Sensitivity of thermal conductivity k (W/m K) of an unfrozen coarse soil to dry density γ_d at constant moisture content w for different methods of calculating k (s_γ is the increase in k per 0.1 g/cm³ increase in γ_d at constant w).

$k_s = 8.0 \text{ W/m K}$							
	<i>Kersten</i>	<i>Johansen</i>	<i>De Vries</i>	<i>Mckley</i>	<i>McGaw</i>	<i>Geman</i>	<i>Van Rooyen</i>
$w = 5\%$							
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm ³	0.121	0.173	0.154	0.203	0.203	0.180	0.342
k at $\gamma_d = 1.3 \text{ g/cm}^3$	0.831	0.853	0.995	2.014	1.937	0.937	1.093
s_γ as % of k	14.6	29.3	15.5	10.1	10.5	19.2	31.3
s_γ for $\gamma_d = 1.5$ to 1.9 g/cm ³	0.215	0.317	0.307	0.260	0.277	0.341	0.977
k at $\gamma_d = 1.7 \text{ g/cm}^3$	1.477	1.782	1.808	2.922	2.877	1.918	3.291
s_γ as % of k	14.6	17.8	17.0	8.9	9.6	17.8	29.7
$w = 10\%$							
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm ³	0.150	0.211	0.202	0.212	0.206	0.236	0.334
k at $\gamma_d = 1.3 \text{ g/cm}^3$	1.028	1.246	1.209	2.063	1.971	1.354	2.095
s_γ as % of k	14.6	16.9	16.7	10.3	10.5	17.4	15.9
s_γ for $\gamma_d = 1.5$ to 1.9 g/cm ³	0.266	0.371	0.362	0.283	0.280	0.414	0.881
k at $\gamma_d = 1.7 \text{ g/cm}^3$	1.827	2.357	2.297	3.026	2.921	2.604	3.939
s_γ as % of k	14.6	15.7	15.8	9.4	9.6	15.9	22.4
$w = 20\%$							
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm ³	0.179	0.249	0.257	0.236	0.223	0.293	0.256
k at $\gamma_d = 1.3 \text{ g/cm}^3$	1.225	1.639	1.568	2.173	2.037	1.800	2.259
s_γ as % of k	14.6	15.2	16.4	10.9	10.9	16.3	11.3
s_γ for $\gamma_d = 1.5$ to 1.7 g/cm ³	0.272	0.367	0.372	0.326	0.261	0.414	0.541
k at $\gamma_d = 1.6 \text{ g/cm}^3$	1.885	2.541	2.493	2.979	2.738	2.846	3.333
s_γ as % of k	14.4	14.4	14.9	10.9	9.5	14.5	16.2
$w = 40\%$							
s_γ for $\gamma_d = 1.1$ to 1.3 g/cm ³	0.128	0.247	0.294	0.318	0.204	0.313	0.200
k at $\gamma_d = 1.2 \text{ g/cm}^3$	1.231	1.765	1.800	2.200	1.958	1.957	2.041
s_γ as % of k	14.5	14.0	16.3	14.5	10.4	16.0	9.8

Table 6. Sensitivity of thermal conductivity k (W/m K) of an unfrozen fine soil to dry density γ_d at constant moisture content w for different methods of calculating k (s_γ is the increase in k per 0.1 g/cm 3 increase in γ_d at constant w).

$k_s = 2.0 \text{ W/m K}$							
	Kersten	Johannsen	De Vries	Macklin	McGraw	Germann	Van Rossum
$w = 8\%$							
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm^3	0.058	0.073	0.065	0.056	0.068	0.066	0.044
k at $\gamma_d = 1.3 \text{ g/cm}^3$	0.401	0.251	0.509	0.579	0.862	0.396	0.231
s_γ as % of k	14.5	29.2	12.7	9.6	7.8	16.7	18.9
s_γ for $\gamma_d = 1.5$ to 1.9 g/cm^3	0.113	0.119	0.109	0.077	0.078	0.113	0.023
k at $\gamma_d = 1.7 \text{ g/cm}^3$	0.713	0.627	0.823	0.832	1.162	0.732	0.451
s_γ as % of k	15.8	19.0	13.3	9.3	6.7	15.4	5.0
$w = 10\%$							
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm^3	0.095	0.088	0.083	0.063	0.073	0.086	0.017
k at $\gamma_d = 1.3 \text{ g/cm}^3$	0.654	0.515	0.613	0.621	0.901	0.577	0.366
s_γ as % of k	14.6	17.1	13.6	10.1	8.1	14.9	4.6
s_γ for $\gamma_d = 1.5$ to 2.0 g/cm^3	0.183	0.125	0.119	0.099	0.081	0.125	-0.068
k at $\gamma_d = 1.7 \text{ g/cm}^3$	1.162	0.923	1.014	0.915	1.205	0.993	0.494
s_γ as % of k	15.8	13.5	11.2	10.8	6.7	12.6	13.7
$w = 20\%$							
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm^3	0.132	0.097	0.104	0.082	0.078	0.103	0.065
k at $\gamma_d = 1.3 \text{ g/cm}^3$	0.907	0.779	0.794	0.716	0.968	0.771	0.649
s_γ as % of k	14.6	12.5	13.0	13.5	8.1	13.3	10.0
s_γ for $\gamma_d = 1.5$ to 1.7 g/cm^3	0.202	0.117	0.124	0.132	0.083	0.120	-0.090
k at $\gamma_d = 1.6 \text{ g/cm}^3$	1.396	1.097	1.135	1.012	1.209	1.104	0.748
s_γ as % of k	14.4	10.7	10.9	13.0	6.9	10.8	-12.0
$w = 40\%$							
s_γ for $\gamma_d = 1.1$ to 1.3 g/cm^3	0.145	0.098	0.126	0.161	0.086	0.110	0.089
k at $\gamma_d = 1.2 \text{ g/cm}^3$	1.005	0.940	0.961	0.862	1.014	0.858	0.751
s_γ as % of k	14.4	10.4	13.1	18.6	8.4	12.8	11.8

Table 7. Sensitivity of thermal conductivity k (W/m K) of a frozen coarse soil to dry density γ_d at constant moisture content w for different methods of calculating k (s_γ is the increase in k per 0.1 g/cm³ increase in γ_d at constant w).

$k_s = 8.0 \text{ W/m K}$				
	Korzeni	Johansen	De Vries	Mikkelsen
$w = 5\%$				
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm^3	0.100	0.129	0.237	0.224
k at $\gamma_d = 1.3 \text{ g/cm}^3$	0.477	0.713	1.640	2.190
s_γ as % of k	21.0	18.1	16.3	10.5
s_γ for $\gamma_d = 1.5$ to 2.0 g/cm^3	0.255	0.332	0.490	0.312
k at $\gamma_d = 1.8 \text{ g/cm}^3$	1.325	1.770	3.554	3.411
s_γ as % of k	19.3	18.7	13.8	9.4
$w = 10\%$				
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm^3	0.177	0.144	0.371	0.252
k at $\gamma_d = 1.3 \text{ g/cm}^3$	0.830	1.260	2.200	2.280
s_γ as % of k	21.3	11.4	16.9	11.1
s_γ for $\gamma_d = 1.5$ to 2.0 g/cm^3	0.454	0.621	0.525	0.420
k at $\gamma_d = 1.8 \text{ g/cm}^3$	2.332	3.227	4.527	3.867
s_γ as % of k	19.5	19.3	11.6	10.9
$w = 20\%$				
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm^3	0.329	0.452	0.462	0.346
k at $\gamma_d = 1.3 \text{ g/cm}^3$	1.536	2.356	3.140	2.700
s_γ as % of k	21.4	19.2	14.7	12.8
s_γ for $\gamma_d = 1.5$ to 1.7 g/cm^3	0.600	0.800	0.551	0.785
k at $\gamma_d = 1.6 \text{ g/cm}^3$	2.867	4.142	4.653	4.006
s_γ as % of k	20.9	19.3	11.8	19.6
$w = 30\%$				
s_γ for $\gamma_d = 1.1$ to 1.4 g/cm^3	0.428	0.602	0.518	0.508
k at $\gamma_d = 1.3 \text{ g/cm}^3$	2.242	3.451	3.893	3.285
s_γ as % of k	19.1	17.4	13.3	15.5
$w = 40\%$				
s_γ for $\gamma_d = 1.1$ to 1.2 g/cm^3	0.451	0.658	0.561	0.647
k at $\gamma_d = 1.2 \text{ g/cm}^3$	2.392	3.760	4.019	3.529
s_γ as % of k	18.9	17.5	14.0	18.3

Table 8. Sensitivity of thermal conductivity k (W/m K) of a frozen fine soil to dry density γ_d at constant moisture content w for different methods of calculating k (s_γ is the increase in k per 0.1 g/cm³ increase in γ_d at constant w).

$$k_s = 2.0 \text{ W/m K}$$

	<i>Kersten</i>	<i>Johansen</i>	<i>De Vries</i>	<i>Mickley</i>
$w = 5\%$				
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm ³	0.061	0.057	0.094	0.063
k at $\gamma_d = 1.3$ g/cm ³	0.361	0.435	0.671	0.610
s_γ as % of k	17.0	13.0	14.0	10.2
s_γ for $\gamma_d = 1.5$ to 2.0 g/cm ³	0.181	0.110	0.138	0.090
k at $\gamma_d = 1.7$ g/cm ³	0.744	0.730	1.129	0.896
s_γ as % of k	24.4	15.1	12.2	10.0
$w = 10\%$				
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm ³	0.093	0.093	0.132	0.080
k at $\gamma_d = 1.3$ g/cm ³	0.635	0.705	0.930	0.726
s_γ as % of k	14.6	13.2	14.1	11.0
s_γ for $\gamma_d = 1.5$ to 2.0 g/cm ³	0.235	0.179	0.146	0.144
k at $\gamma_d = 1.7$ g/cm ³	1.177	1.187	1.507	1.110
s_γ as % of k	19.9	15.1	9.7	12.9
$w = 20\%$				
s_γ for $\gamma_d = 1.1$ to 1.5 g/cm ³	0.157	0.166	0.170	0.136
k at $\gamma_d = 1.3$ g/cm ³	1.181	1.245	1.412	1.010
s_γ as % of k	13.3	13.3	12.0	13.5
s_γ for $\gamma_d = 1.5$ to 1.7 g/cm ³	0.251	0.242	0.173	0.387
k at $\gamma_d = 1.6$ g/cm ³	1.771	1.841	1.927	1.554
s_γ as % of k	14.1	13.1	9.0	24.9
$w = 30\%$				
s_γ for $\gamma_d = 1.1$ to 1.4 g/cm ³	0.203	0.224	0.203	0.253
k at $\gamma_d = 1.3$ g/cm ³	1.728	1.785	1.850	1.426
s_γ as % of k	11.8	12.5	11.0	17.7
$w = 40\%$				
s_γ for $\gamma_d = 1.1$ to 1.2 g/cm ³	0.229	0.260	0.239	0.370
k at $\gamma_d = 1.2$ g/cm ³	2.013	2.034	2.049	1.739
s_γ as % of k	11.4	12.8	11.7	21.3

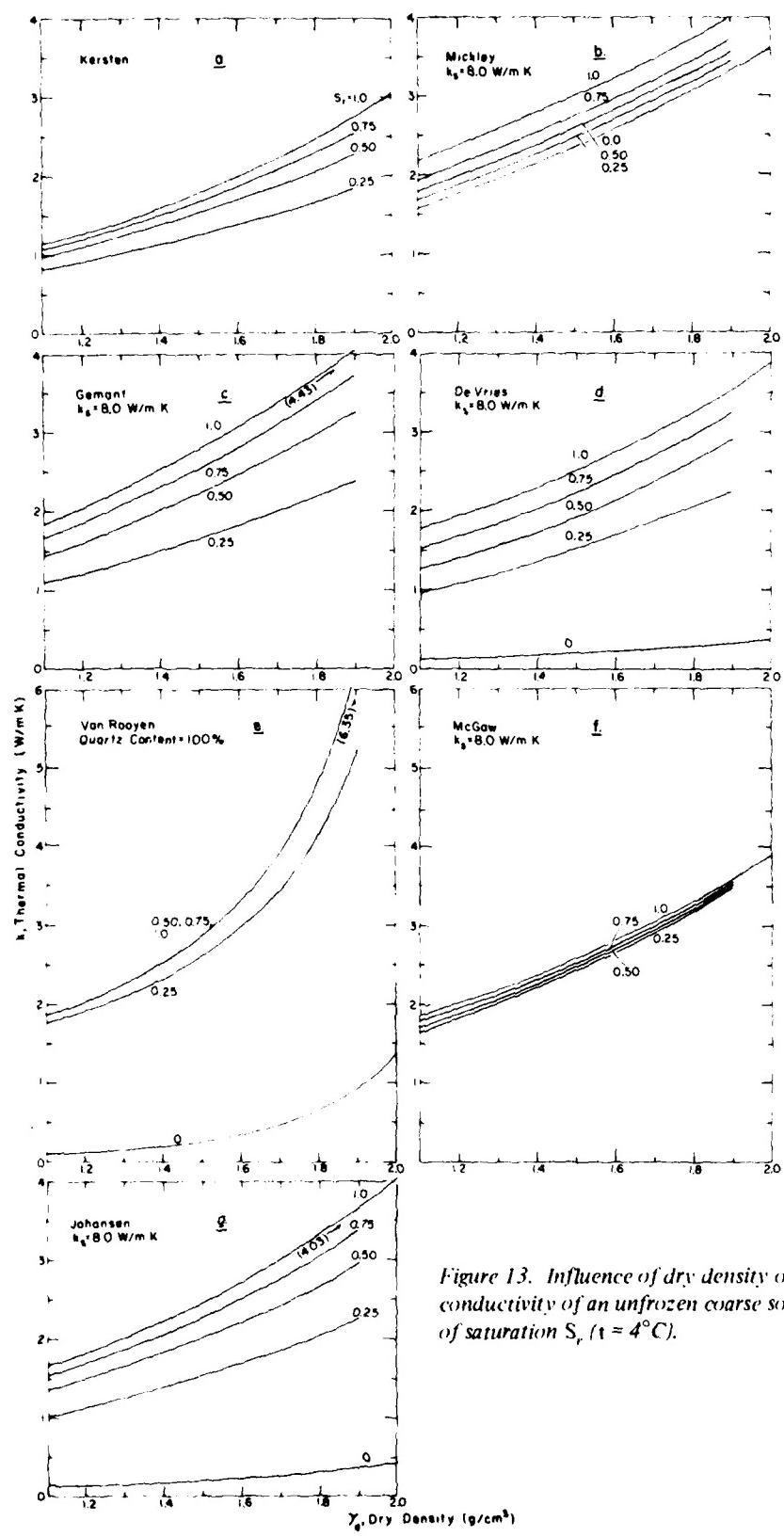


Figure 13. Influence of dry density on calculated thermal conductivity of an unfrozen coarse soil at a constant degree of saturation S_r ($t = 4^\circ\text{C}$).

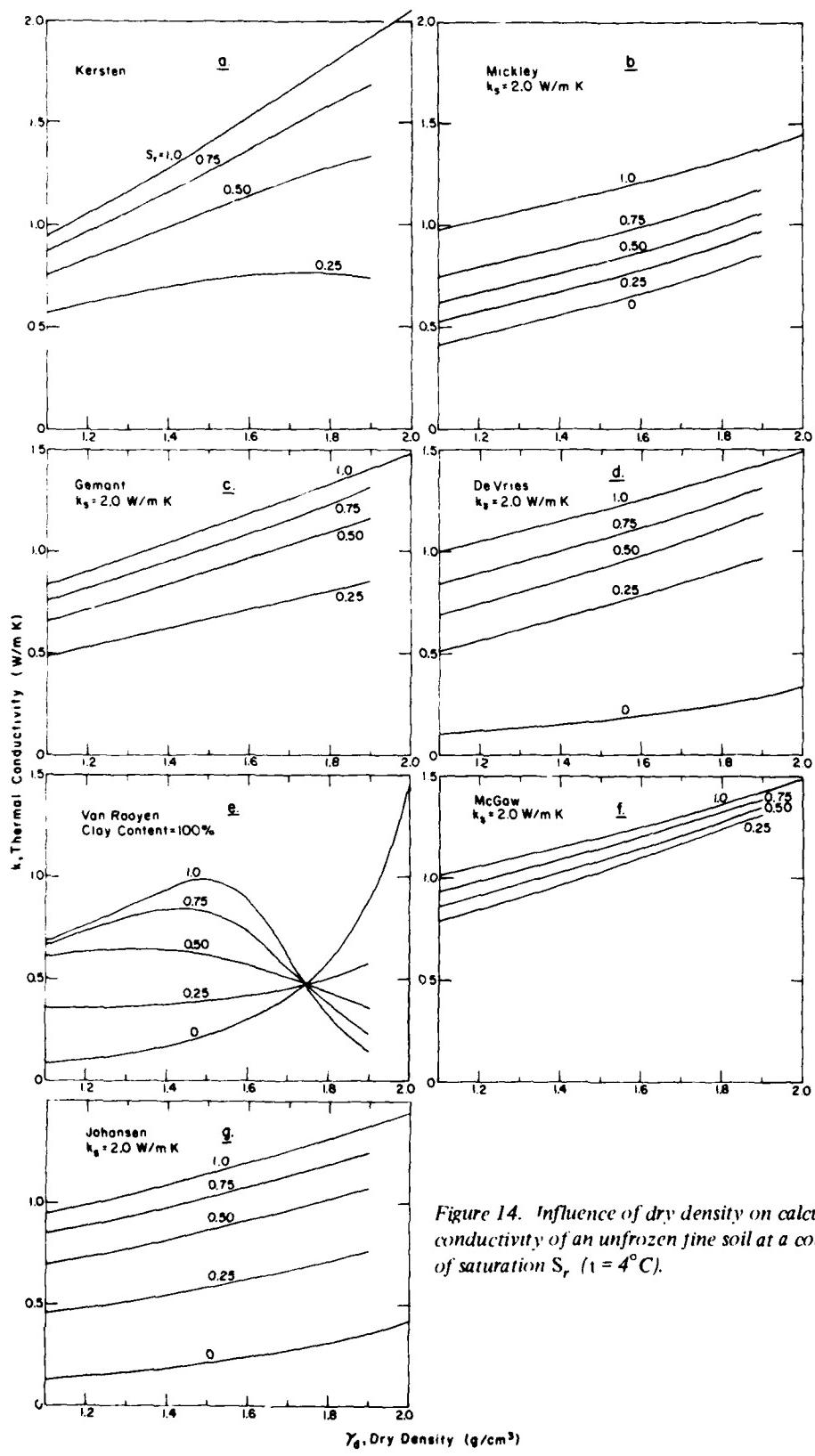


Figure 14. Influence of dry density on calculated thermal conductivity of an unfrozen fine soil at a constant degree of saturation S_r ($t = 4^\circ\text{C}$).

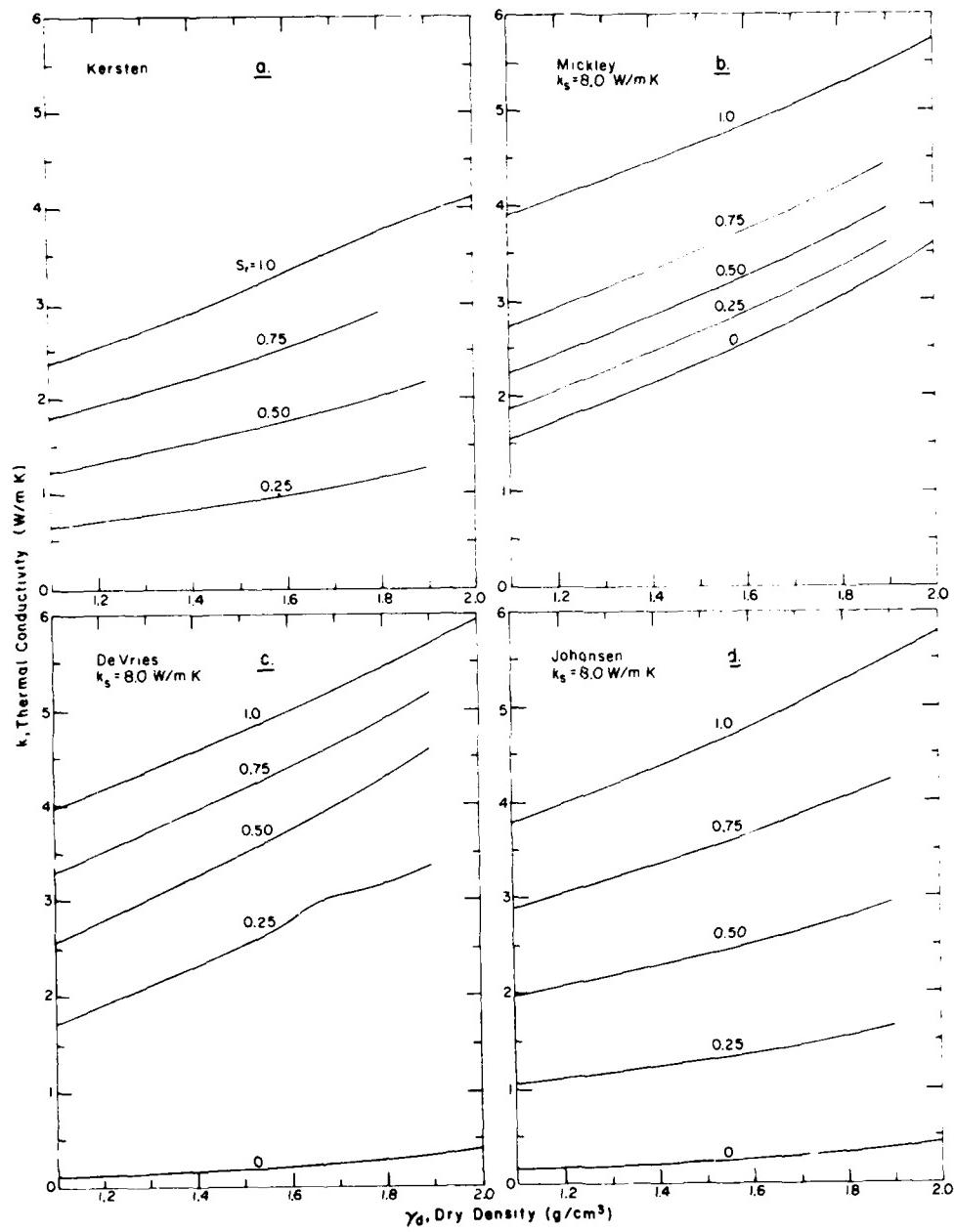


Figure 15. Influence of dry density on calculated thermal conductivity of a frozen coarse soil at a constant degree of saturation S_r ($t = -4^\circ\text{C}$).

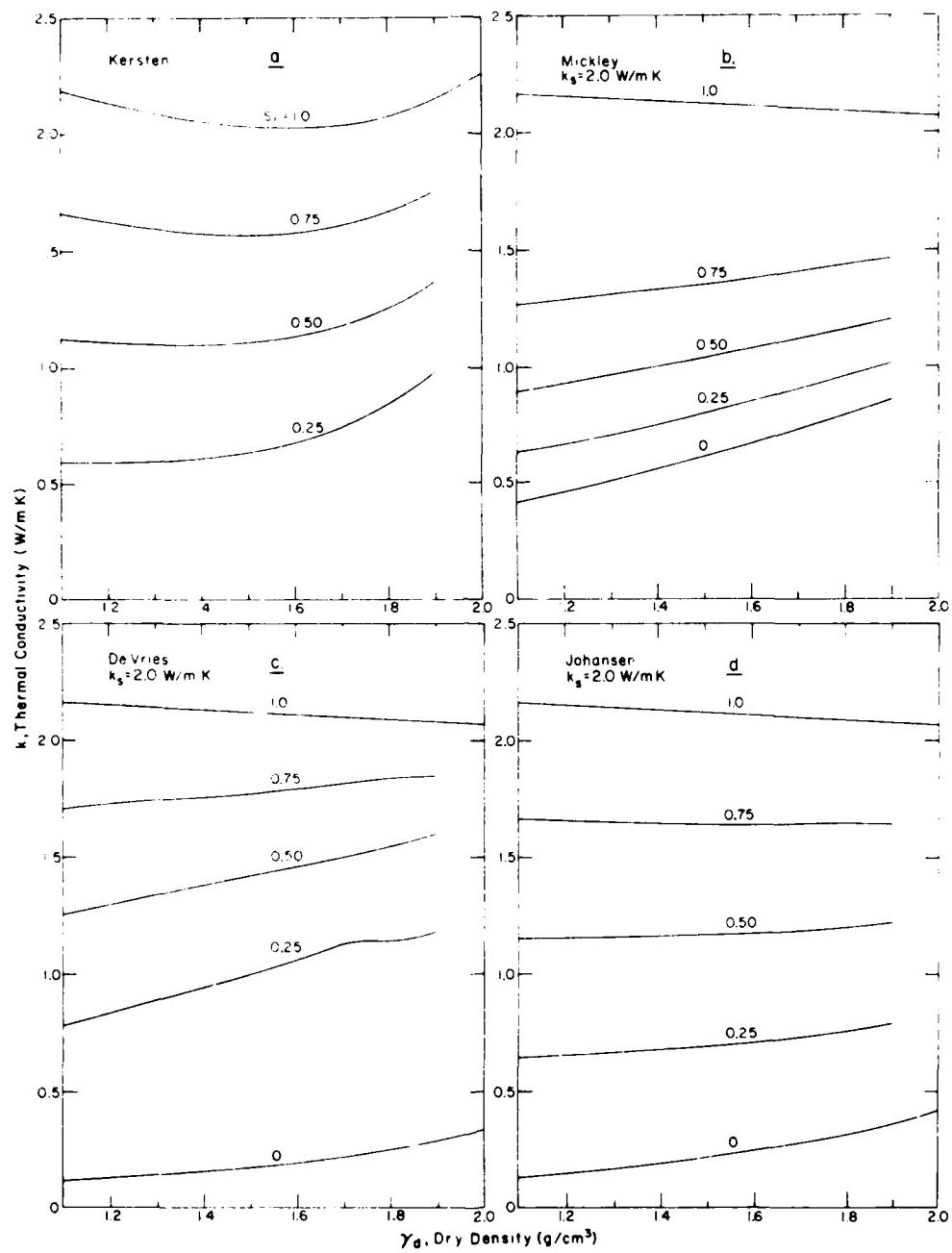


Figure 16. Influence of dry density on calculated thermal conductivity of a frozen fine soil at a constant degree of saturation S_r ($t = -4^\circ\text{C}$).

Table 9. Sensitivity of soil thermal conductivity k (W/m K) of an unfrozen soil to solids conductivity k_s at constant degree of saturation S_f for different methods.

$\gamma_d = 1.6 \text{ g/cm}^3$									
Johansen					Van Rossum				
Coarse	Fine	De Vries	Mod. res.	Kunii-Smith	Mickey	McGraw	Gemant	Coarse	
$S_f = 1.0$									
$k_s = 8.0 \text{ W/m K}$	2.719	-	2.736	3.265	2.270	3.218	2.779	3.066	2.327
$k_s = 4.0$	-	-	-	-	-	-	-	-	0.643
$k_s = 3.0$	1.521	1.521	-	-	-	-	-	-	0.415
$k_s = 2.0$	-	1.197	1.247	1.324	1.129	1.208	1.253	1.181	-
Sensitivity k/k_s	0.240	0.324	0.248	0.324	0.190	0.335	0.254	0.314	0.671*
$S_f = 0.5$									
$k_s = 8.0 \text{ W/m K}$	2.200	-	2.131	-	-	2.811	2.678	2.444	3.327
$k_s = 4.0$	-	-	-	-	-	-	-	-	0.643
$k_s = 3.0$	1.253	1.137	-	-	-	-	-	-	0.415
$k_s = 2.0$	-	0.909	0.973	-	-	0.874	1.149	0.959	-
Sensitivity k/k_s	0.189	0.228	-	-	-	0.323	0.255	0.248	0.511*
$S_f = 0.25$									
$k_s = 8.0 \text{ W/m K}$	1.677	-	1.697	-	-	2.696	2.625	1.823	2.960
$k_s = 4.0$	-	-	-	-	-	-	-	-	-
$k_s = 3.0$	0.982	0.751	-	-	-	-	-	-	0.628
$k_s = 2.0$	-	0.621	0.790	-	-	0.780	1.097	0.716	0.409
Sensitivity k/k_s	0.139	0.130	0.151	-	-	0.319	0.255	0.185	0.583*
$S_f = 0$ (dry)									
$k_s = 8.0 \text{ W/m K}$	0.240	-	0.216	0.302	0.259	-	-	-	0.339
$k_s = 4.0$	-	-	-	-	-	-	-	-	-
$k_s = 3.0$	0.240	0.240	-	-	-	-	-	-	0.198
$k_s = 2.0$	-	0.240	0.193	0.260	0.183	-	-	-	-
Sensitivity k/k_s	0.000	0.000	0.0038	0.0070	0.0127	-	-	-	0.0282†

*For k_s varying from 4.0 to 8.0 W/m K

†For k_s varying from 3.0 to 8.0 W/m K

Influence of soil solids' thermal conductivity

A determination was made of the variation in the calculated soil thermal conductivity due to changes in the value of the solids' thermal conductivity k_s at a constant γ_d of 1.6 g/cm³ and for a constant S_f . Calculations were made for several values of S_f , ranging from dry to saturated. The resulting curves for each S_f value are shown in Figure 17 for unfrozen soil and in Figure 18 for frozen soil. The Kersten method could not be applied as it does not explicitly take k_s into account so that it cannot allow for variation in k_s .

Six methods could be generally applied at all S_f values to unfrozen soils, while only three of these were applicable to frozen soils. In the fully saturated state, two additional methods, Kunii-Smith and the modified resistor equation, were also applicable to both frozen and unfrozen soils. When Johansen was applied at $k_s = 3.0 \text{ W/m K}$ and below, the equation appropriate to fine soil was used. This gives rise to different sensitivities with this method for coarse and for fine soils.

Tables 9 and 10 show the values obtained for the absolute sensitivity of the soil thermal conductivity

Table 10. Sensitivity of soil thermal conductivity k (W/m K) of a frozen soil to solids conductivity k_s at constant degree of saturation S_f for different methods.

$\gamma_d = 1.6 \text{ g/cm}^3$						
<i>Johansen</i>						
	<i>Coarse</i>	<i>Fine</i>	<i>De Vries</i>	<i>Mod. res.</i>	<i>Kuui-Smith</i>	<i>Mickley</i>
$S_f = 1.0$						
$k_s = 8.0 \text{ W/m K}$	4.793	-	4.996	5.300	4.524	4.837
$k_s = 3.0$	2.682	2.682	-	-	-	-
$k_s = 2.0$	-	2.110	2.111	2.113	2.136	2.110
Sensitivity $ k /k_s$	0.422	0.572	0.481	0.531	0.398	0.455
$S_f = 0.75$						
$k_s = 8.0 \text{ W/m K}$	3.654	-	4.395	-	-	3.724
$k_s = 3.0$	2.070	2.070	-	-	-	-
$k_s = 2.0$	-	1.641	1.792	-	-	1.371
Sensitivity $ k /k_s$	0.317	0.429	0.434	-	-	0.392
$S_f = 0.5$						
$k_s = 8.0 \text{ W/m K}$	2.522	-	3.727	-	-	3.247
$k_s = 3.0$	1.463	1.463	-	-	-	-
$k_s = 2.0$	-	1.176	1.463	-	-	1.069
Sensitivity $ k /k_s$	0.212	0.287	0.377	-	-	0.363
$S_f = 0.25$						
$k_s = 8.0 \text{ W/m K}$	1.381	-	2.792	-	-	2.875
$k_s = 3.0$	0.851	0.851	-	-	-	-
$k_s = 2.0$	-	0.708	1.064	-	-	0.845
Sensitivity $ k /k_s$	0.106	0.143	0.288	-	-	0.338

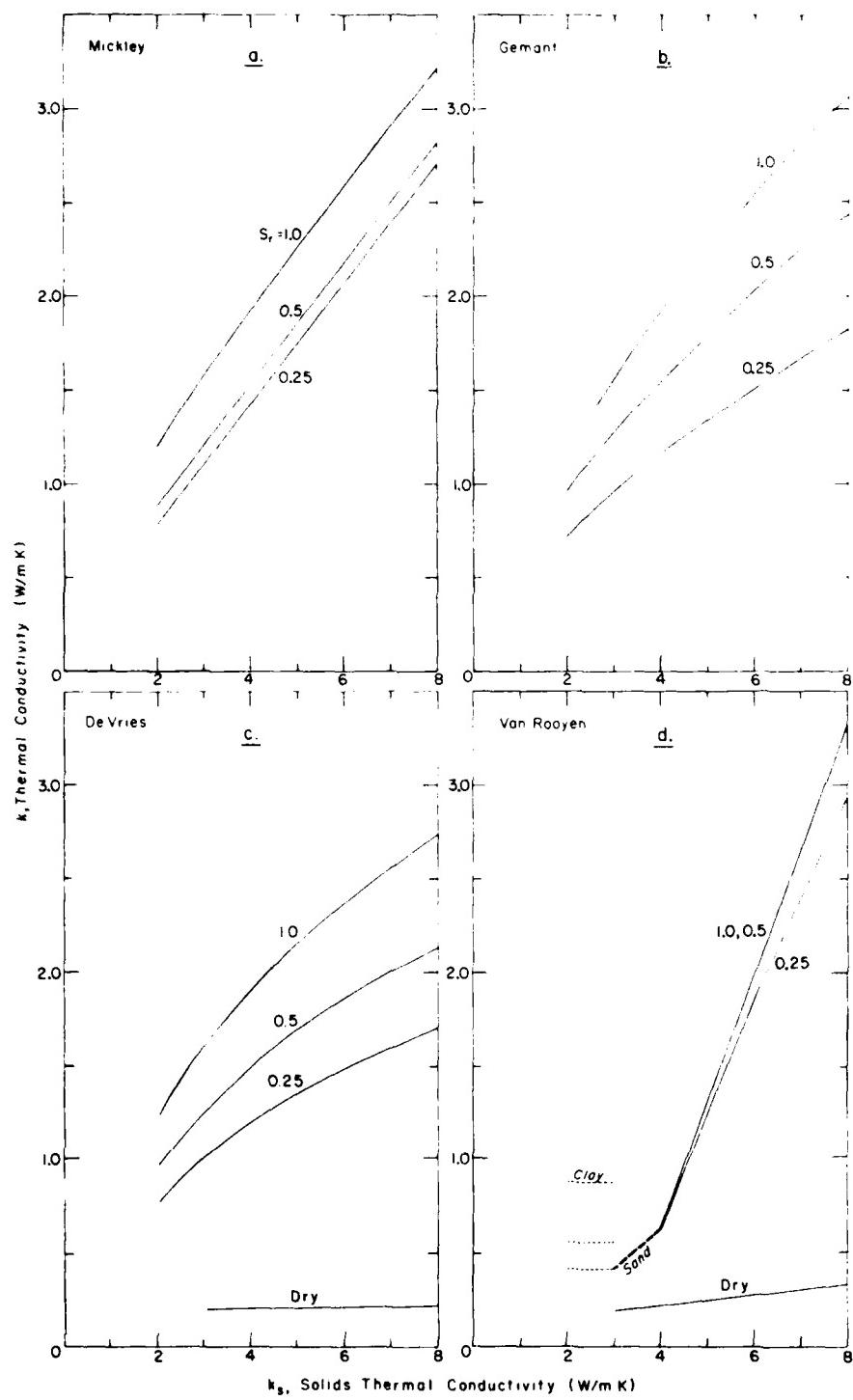


Figure 17. Influence of soil solids' thermal conductivity on calculated thermal conductivity of an unfrozen soil at a constant degree of saturation S_r ($t = 4^\circ\text{C}$, $\gamma_d = 1.6 \text{ g/cm}^3$).

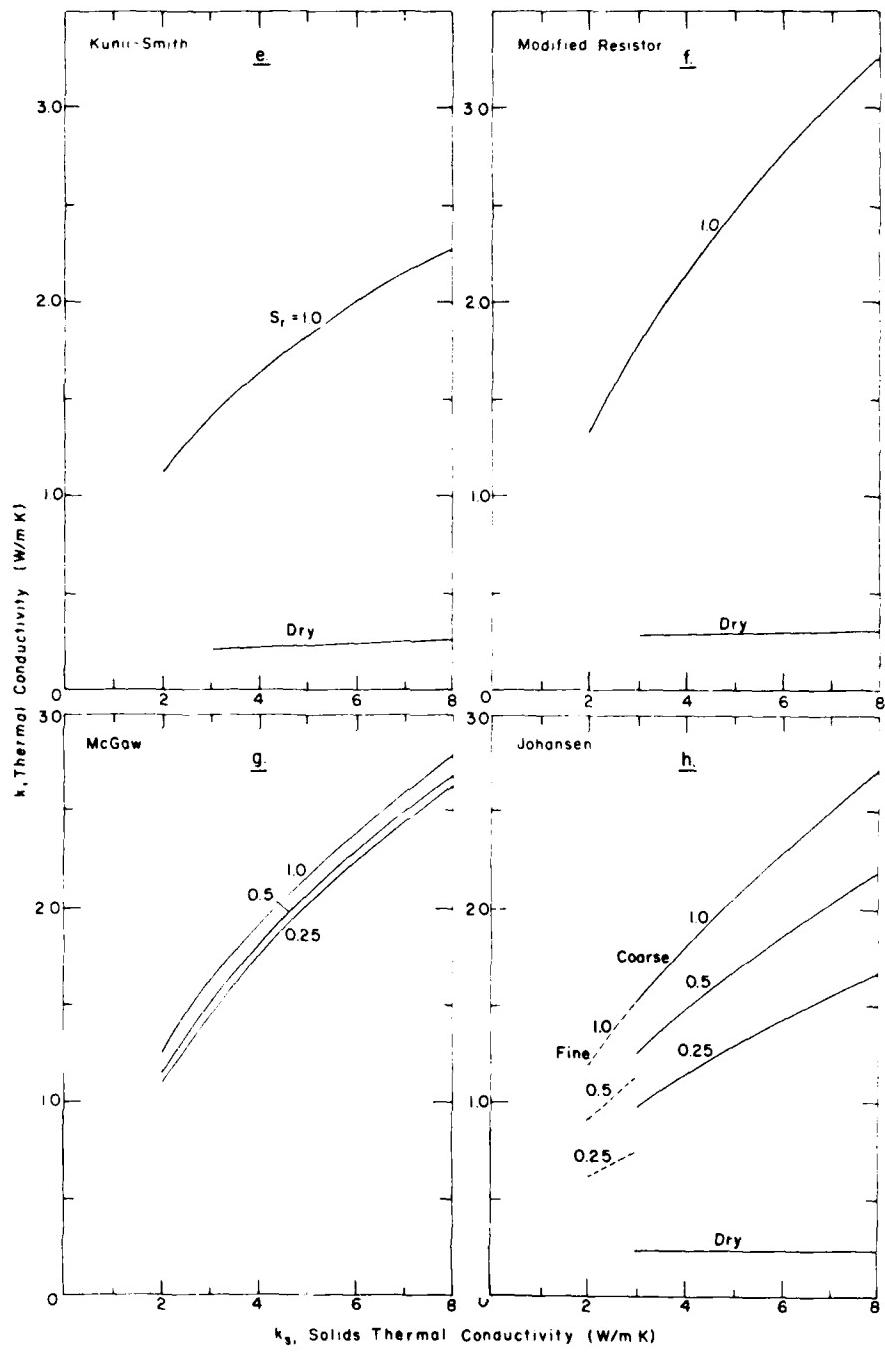


Figure 17 (cont'd).

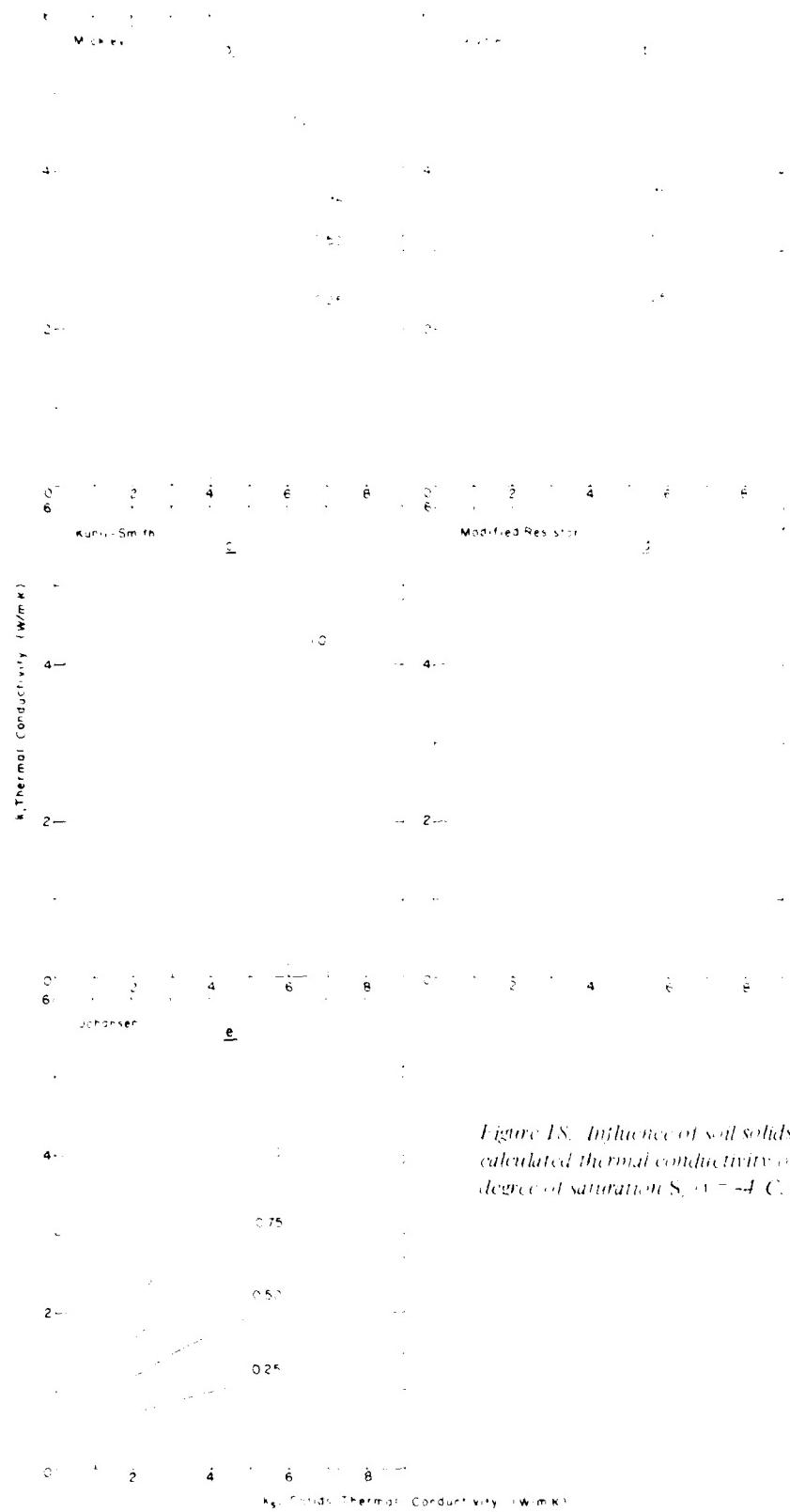


Figure 18. Influence of soil solids' thermal conductivity on calculated thermal conductivity of a frozen soil at a constant degree of saturation S_s ; $T = -4^\circ\text{C}$; $\gamma_f = 1.6 \text{ g cm}^{-3}$

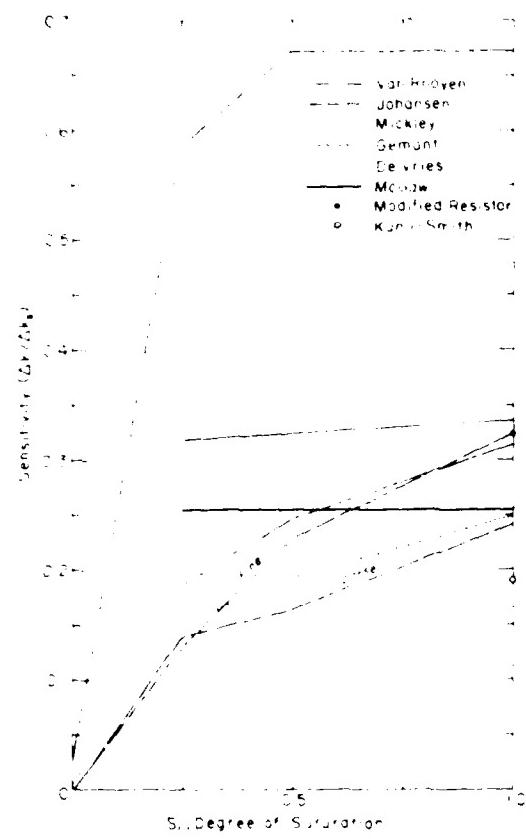


Figure 19. Sensitivity of calculated thermal conductivity to soil solids' thermal conductivity ($k - k_s$) in an unfrozen soil as influenced by degree of saturation S_r , $\gamma_d = 1.6 \text{ g cm}^{-3}$.

$\frac{\partial k}{\partial k_s}$ is the change in the soil thermal conductivity k due to unit change in k_s at constant S_r . The dependence of $\frac{\partial k}{\partial k_s}$ on S_r is shown in Figure 19 for unfrozen soil and in Figure 20 for frozen soil.

In the case of unfrozen soil, Van Rooyen gives extremely high sensitivities and differs from the other methods considerably. Johansen, De Vries and Gemant produce sensitivities which increase markedly as S_r increases, while the sensitivities from Mickley and McGraw are almost constant with S_r . For saturated soils Johansen (coarse), De Vries and McGraw give nearly the same sensitivity of about 0.25 while Johansen (fine), Gemant, Mickley and modified resistor give a higher value of about 0.32 (Fig. 19).

For frozen soil Johansen (fine and coarse), De Vries and Mickley give higher sensitivities (i.e. $k - k_s$) than for unfrozen soil. Also the rate of increase in the sensitivity with increasing S_r is more marked for the frozen soil. Johansen (fine) shows a particularly large rate of increase (Fig. 20). It is noteworthy

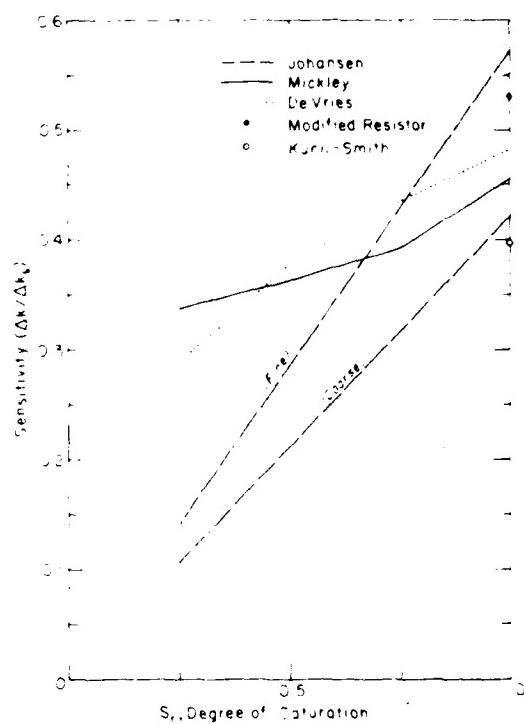


Figure 20. Sensitivity of calculated thermal conductivity to soil solids' thermal conductivity ($k - k_s$) in a frozen soil as influenced by degree of saturation S_r , $\gamma_d = 1.6 \text{ g cm}^{-3}$.

that the sensitivities for saturated frozen soil are much higher than the corresponding values for saturated unfrozen soil. This holds for Johansen (coarse and fine), modified resistor, Kunn-Smith, De Vries and Mickley.

Comparison of the various methods

The thermal conductivity values predicted by the various methods are compared for the saturated condition, for $S_r = 0.5$ and for the dry condition in Figures 21-30. This has been done over a γ_d range from 1.1 to 2.0 g cm^{-3} . Apart from Kersten the values chosen for k_s were $8.0 \text{ W m}^{-1} \text{ K}$ for coarse soil and $2.0 \text{ W m}^{-1} \text{ K}$ for fine soil.

The nine methods applicable to saturated unfrozen coarse soil all show a similar trend for the increase in thermal conductivity with increasing γ_d (Fig. 21). Kersten and Van Rooyen give the lowest thermal conductivity values while Mickley, modified resistor and Gemant give the highest values. Johansen, De Vries and McGraw give almost coincident curves.

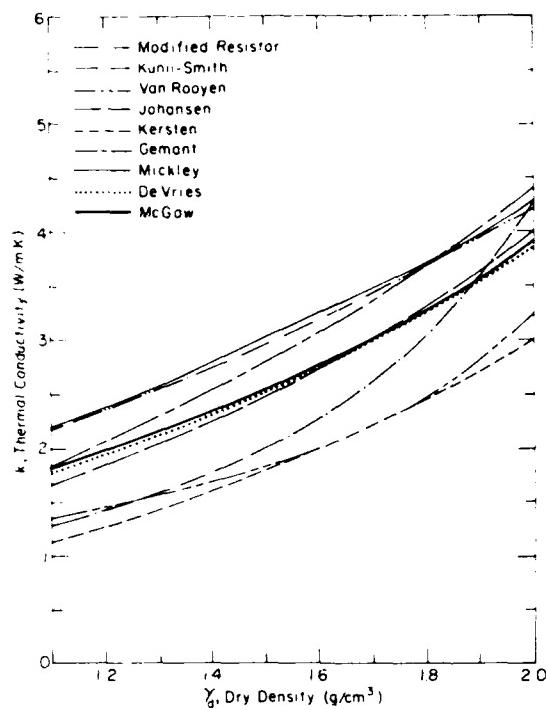


Figure 21. Comparison of thermal conductivity values calculated by the various methods for a saturated unfrozen coarse soil at different dry densities ($t = 4^\circ C$, $k_s = 8.0 \text{ W/m K}$).

For saturated unfrozen fine soil (Fig. 22), Kersten gives the highest values above a γ_d of 1.3 g/cm^3 . Van Rooyen provides the lowest values and is unreliable at values of γ_d over 1.5 g/cm^3 . Johansen, De Vries, modified resistor, Kunii-Smith, Mickley and Geman differ noticeably at low γ_d values, but show a similar, almost linear, trend and closely approach each other at high γ_d .

In the case of saturated frozen coarse soil (Fig. 23), Kersten gives much lower values than the other five methods. Modified resistor gives the highest values, showing a linear increase to which De Vries, Mickley and Johansen are approximately parallel.

With regard to saturated frozen fine soil (Fig. 24), the six methods give values which differ little from each other. In fact, Johansen, De Vries, modified resistor and Mickley all give coincident curves. Also, the thermal conductivity does not vary much with γ_d as may be expected because the replacement of ice, having a k_i of 2.2 W/m K , by the mineral solid, having a k_s of 2.0 W/m K , should not make much difference to the overall soil thermal conductivity.

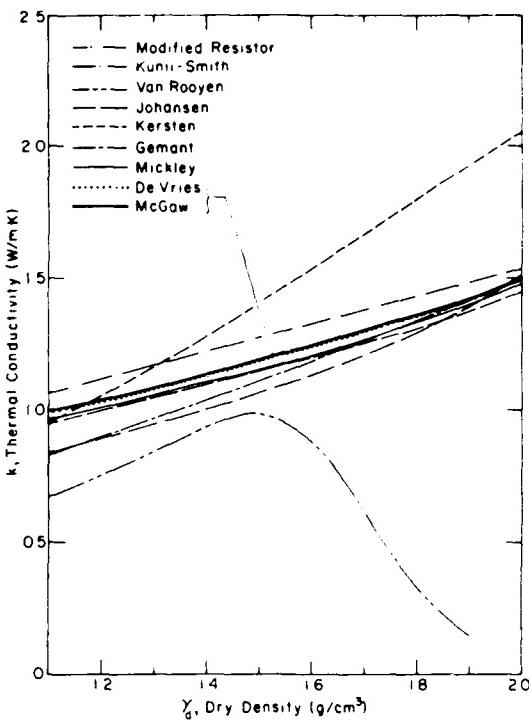


Figure 22. Comparison of thermal conductivity values calculated by the various methods for a saturated unfrozen fine soil at different dry densities ($t = 4^\circ C$, $k_s = 2.0 \text{ W/m K}$).

The seven curves for the partially saturated unfrozen sand ($S_r = 0.5$) all show a similar trend, except for Van Rooyen's, which rises rapidly above a γ_d of 1.4 g/cm^3 (Fig. 25). As with the saturated coarse soil, Kersten gives the lowest values while Mickley gives the highest (apart from Van Rooyen, which shows odd behavior). The curves given by Johansen and De Vries are close together and represent roughly average values.

With partially saturated unfrozen fine soil, some opposite trends are apparent (Fig. 26). Van Rooyen and Mickley give the lowest values, with Van Rooyen showing an incongruent decrease at γ_d above 1.3 g/cm^3 . Contrary to the coarse soil case, Kersten gives one of the highest curves, while De Vries, Geman and Johansen give curves at intermediate positions.

For partially saturated frozen coarse soil, the four applicable methods give curves which differ by large amounts, although they show similar trends (Fig. 27). Kersten gives the lowest thermal conductivity values while De Vries gives the highest. At high γ_d , De Vries gives values that are more than twice as high as those given by Kersten.

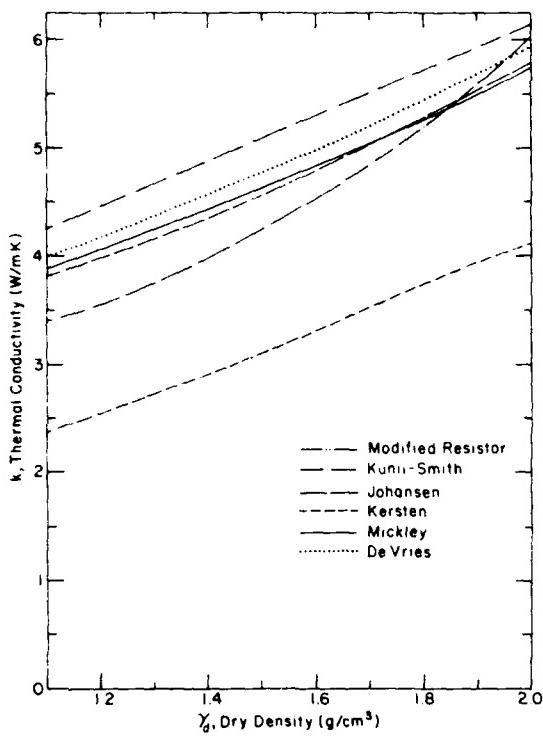


Figure 23. Comparison of thermal conductivity values calculated by the various methods for a saturated frozen coarse soil at different dry densities ($t = -4^\circ\text{C}$, $k_s = 8.0 \text{ W/m K}$).

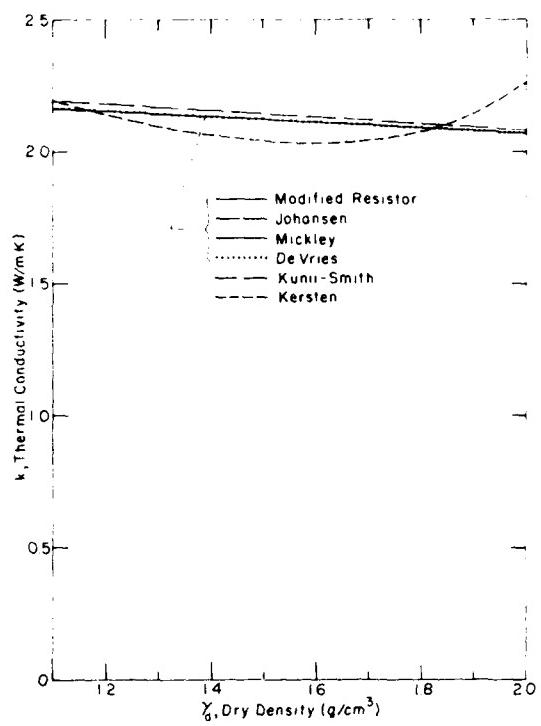


Figure 24. Comparison of thermal conductivity values calculated by the various methods for a saturated frozen fine soil at different dry densities ($t = -4^\circ\text{C}$, $k_s = 2.0 \text{ W/m K}$).

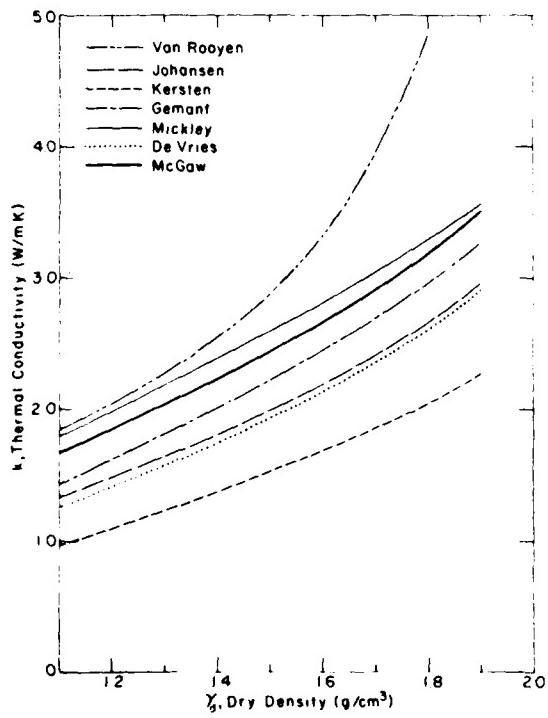


Figure 25. Comparison of thermal conductivity values calculated by the various methods for an unsaturated unfrozen coarse soil at different dry densities with $S_r = 0.5$ ($t = 4^\circ\text{C}$, $k_s = 8.0 \text{ W/m K}$).

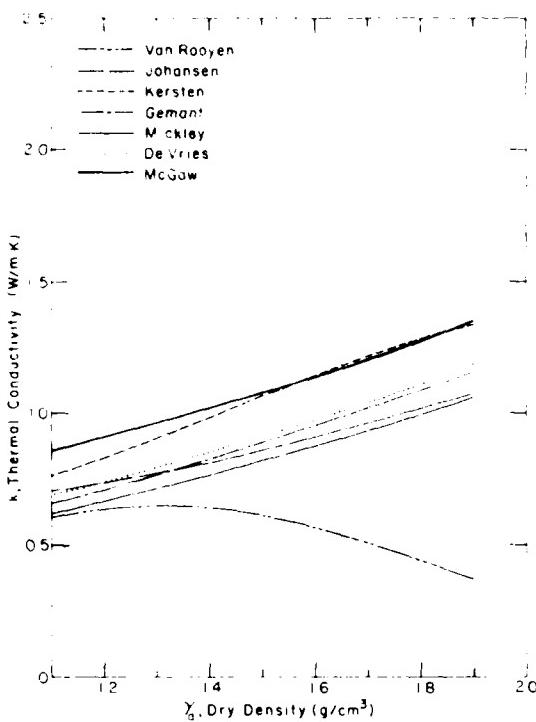


Figure 26. Comparison of thermal conductivity values calculated by the various methods for an unsaturated unfrozen fine soil at different dry densities with $S_r = 0.5$ ($t = 4^\circ\text{C}$, $k_s = 2.0 \text{ W/m K}$).

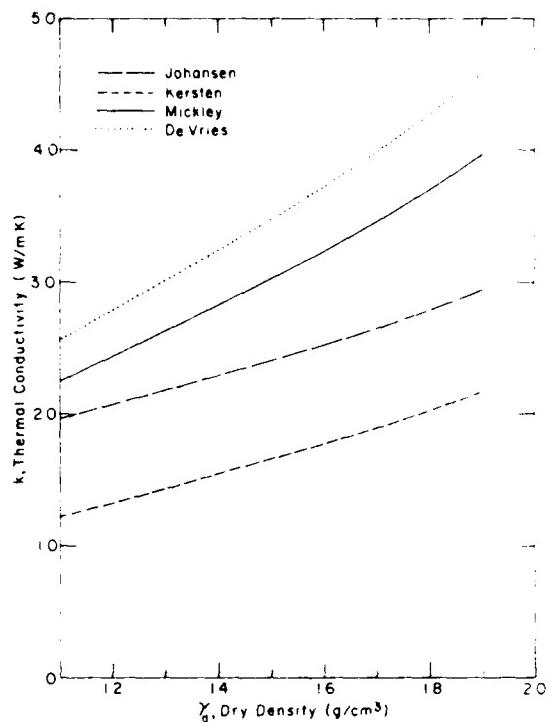


Figure 27. Comparison of thermal conductivity values calculated by the various methods for an unsaturated frozen coarse soil at different dry densities with $S_r = 0.5$ ($t = -4^\circ\text{C}$, $k_s = 8.0 \text{ W/m K}$).

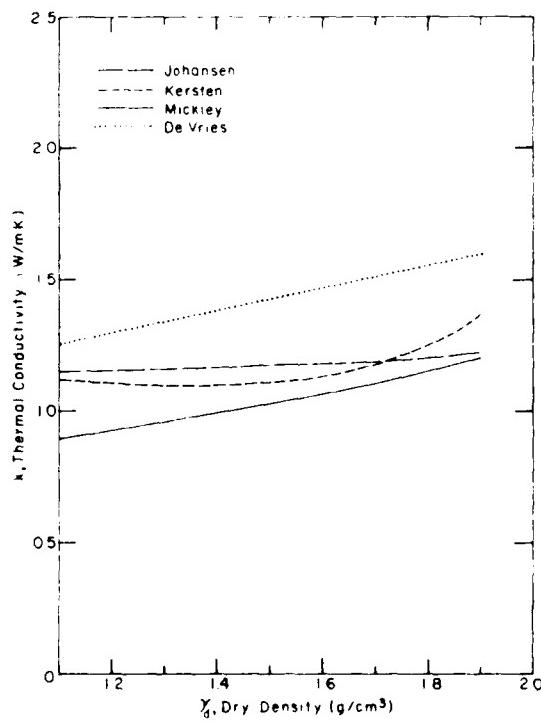


Figure 28. Comparison of thermal conductivity values calculated by the various methods for an unsaturated frozen fine soil at different dry densities with $S_r = 0.5$ ($t = -4^\circ\text{C}$, $k_s = 2.0 \text{ W/m K}$).

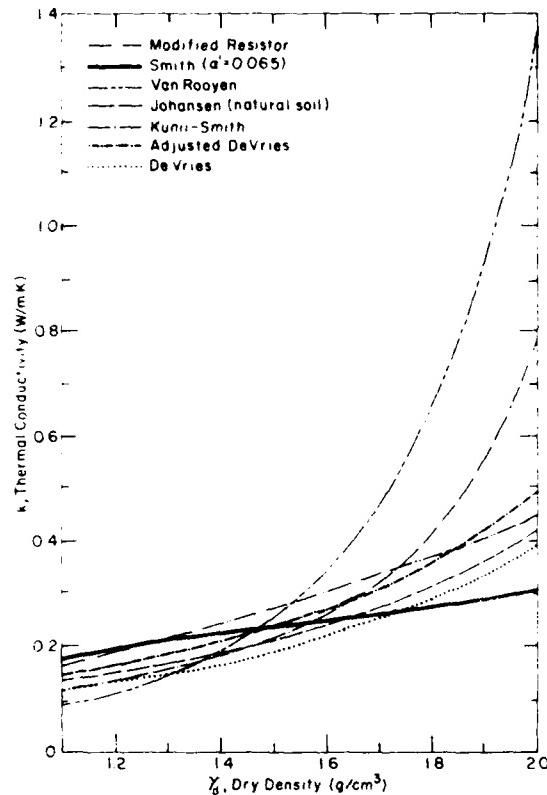


Figure 29. Comparison of thermal conductivity values calculated by the various methods for a dry coarse soil at different dry densities ($t = 4^\circ\text{C}$, $k_s = 8.0 \text{ W/m K}$).

The differences are less pronounced for frozen fine soil at partial saturation (Fig. 28). Kersten and Johansen do not differ by much over the whole γ_d range, and Mickley approaches them at high γ_d . De Vries gives the highest values, about 15% greater than Kersten and 30% greater than Johansen or Mickley at the highest γ_d (1.9 g/cm^3).

For dry coarse soil all the applicable methods show a similar trend, except that the curves of Van Rooyen and Kunii-Smith start to rise quickly above $\gamma_d = 1.5 \text{ g/cm}^3$ (Fig. 29). Johansen and De Vries give curves that are close together and nearly parallel, and modified resistor gives a similar curve that is higher. The adjusted De Vries equation gives higher values than the modified resistor equation at the highest γ_d . Smith shows a somewhat different trend, giving the lowest values at high γ_d . For this method the value chosen for α' (thermal structure factor) was 0.065.

In the case of dry fine soil (Fig. 30), Van Rooyen again shows an excessive rate of increase in the

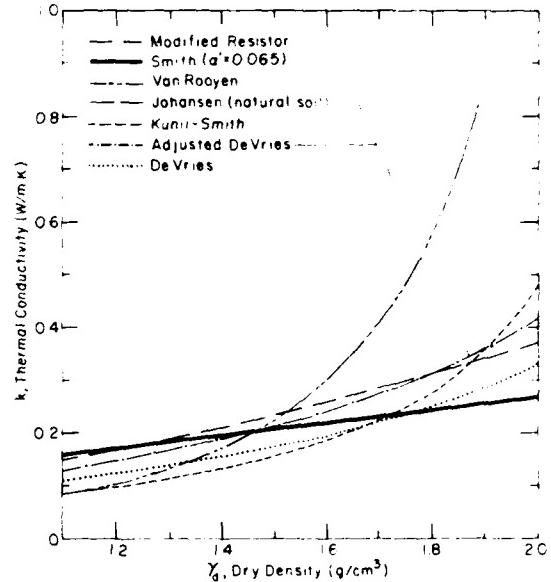


Figure 30. Comparison of thermal conductivity values calculated by the various methods for a dry fine soil at different dry densities ($t = 4^\circ\text{C}$, $k_s = 2.0 \text{ W/m K}$).

thermal conductivity above $\gamma_d = 1.5 \text{ g/cm}^3$, but Kunii-Smith appears reasonable up to $\gamma_d = 1.9 \text{ g/cm}^3$. The curves from Johansen and adjusted De Vries are practically coincident, while Smith again gives lower values at a γ_d above 1.8 g/cm^3 .

The temperature difference of 8°C between the frozen and unfrozen dry conditions does not produce appreciable changes in the soil thermal conductivity.

In the next section there is a detailed testing of all the methods against actual experimental data for different soil types and conditions. The methods are evaluated so as to determine under which conditions each method gives good predictions and in order to recommend the best method(s) applicable to soils in various conditions.

EVALUATION OF METHODS FOR CALCULATING THERMAL CONDUCTIVITY

This section presents an evaluation of the various proposed methods for calculating the thermal

conductivity of soils. This evaluation was carried out using a computer program which analyzed data obtained by various experimenters on soils with known characteristics. In particular, knowledge of the quartz content was important. The thermal conductivity predicted by each method was then computed at appropriate values of the moisture content and dry density. The deviation of this computed value from the measured value was then obtained. (It should be noted that measured values may not always be accurate).

Comparison of the deviations produced by the various methods indicates which methods give good agreement under the relevant conditions. The evaluation is done for moist coarse and fine soils, unfrozen or frozen, and for dry soils.

Soils data used for evaluation

All the methods for calculating soil thermal conductivity, except Kersten, depend on knowledge of the solids thermal conductivity k_s . The solids component which could have a major influence on the value of k_s is quartz, because quartz has a thermal conductivity appreciably larger than any other possible soil solid component. In order, therefore, to get a reliable value of k_s for use in any of these methods, it is essential to know the quartz content of the soil solids. So the different methods were evaluated by comparing their predictions with the thermal conductivity values measured on soils with known quartz content. The quartz content was supplied by Kersten (1949) and Johansen (pers. comm.) for each of the soils and soil materials they tested.

The evaluation was also carried out on other soils for which values of k_s were either suggested by the respective experimenters (e.g. Penner 1970, Smith and Byers 1938) or could be suitably chosen, depending on the type of soil. In these cases uncertainty about the actual k_s value makes the evaluation inaccurate to some extent. However, such an evaluation should show the main trends in the predictions and deviations, providing at least a general comparison.

Computer program

The computer program* calculates the value of the soil thermal conductivity for the given soil input data. This is done for each of the methods appropriate to the particular soil condition, frozen or unfrozen, and saturated, unsaturated or dry. The input data consist of the soil sample's dry density, moisture content, unfrozen water content (where

*Available from the author.

applicable) and the thermal conductivity value measured for that sample.

Other basic parameters to be input include the quartz content q (as a fraction of the solids content) and the values of the thermal conductivities of the soil components at the temperature of measurement. These are the thermal conductivity values for the soil air, water and ice, and those for quartz (k_q) and the solid components other than quartz (k_o). From k_o and k_q the program calculates k_s using the geometric mean equation

$$k_s = k_o^{(1-q)} k_q^q.$$

Alternatively, k_s may be input directly if it is known or if a suitable value is assumed.

When Van Rooyen is applied to fine soils, a value for the clay content has to be input as known or as estimated. Van Rooyen also differentiates fine soils into clay and silt. When Smith is applied to dry soils (coarse or fine), a suitable value must be input for the thermal structural factor α' .

The computer program converts the input value of measured thermal conductivity into W/m K if it is not already in these units. It also converts the dry density to metric units. It then calculates the thermal conductivity given by each method selected and determines its deviation from the measured value. This deviation is expressed as a percentage of the measured value. The program also calculates the porosity of the soil, its degree of saturation and its unfrozen water content as a fraction by volume of the total soil volume.

Input parameters

Specific gravity of the soil solids. This was generally taken as 2.70, unless a particular value was given by the experimenter. The program uses this specific gravity and the input soil dry density to calculate the soil porosity.

Temperature. It was necessary to know the temperature at which the measurement was made so that the values of the thermal conductivity of the soil components appropriate to this temperature could be input. In cases where the experimenters had not specified the temperature, a suitable value was assumed.

Thermal conductivity of soil air k_a . For dry soil k_a was taken as 0.024 W/m K at 0°C or below and 0.025 W/m K at 10°C. Where the soil was moist and unsaturated, moisture migration effectively increases k_a , depending upon the temperature. Under these conditions the values of k_a chosen at different temperatures were taken or interpolated from Table 11 which follows from the suggestions of De Vries (1963).

Table 11. Apparent thermal conductivity of moist pore air k_a (based on De Vries 1963)*.

Temperature (°C)	k_a (W/m K)
0	0.046
5	0.055
10	0.067
15	0.078
20	0.099
25	0.120
30	0.151

* Assuming air is saturated with water vapor.

Thermal conductivity of soil water k_w . The thermal conductivity of the soil water k_w was calculated at different temperatures using the following equation given in the *Thermophysical Properties of Matter* (Touloukian et al. 1970):

$$10^6 k_w = -1390.53 + 15.1937 T \\ - 0.0190398 T^2$$

which gives k_w in units of cal/cm s °C, T being the absolute temperature. The resulting values of k_w are given in Table 12. For intermediate temperatures, k_w was interpolated linearly.

Table 12. Thermal conductivity of water k_w *

Temperature (°C)	k_w (W/m K)
0	0.560
5	0.570
10	0.579
15	0.588
20	0.597
25	0.605
30	0.613

*Based on the equation given by Touloukian et al. (1970).

Unfrozen water content UWC. In some cases the unfrozen water content was known from the experimenter's data (e.g. Penner's Leda clay). In other cases a suitable value of UWC was assumed, depending on the soil texture, or the UWC was simply set at zero.

For unfrozen soils the UWC value in the computer printout tables in Appendix B represents the moisture

content of the soil expressed as a fraction of the total soil volume.

Thermal conductivity of soil ice k_i . To determine the value of k_i , the following formula quoted by Sawada (1977) was used:

$$k_i = 488.19/T + 0.4685 \text{ W/m K}$$

T being the absolute temperature. The resulting values of k_i at different temperatures are given in Table 13. At intermediate temperatures, k_i was linearly interpolated.

Table 13. Thermal conductivity of ice k_i *

Temperature (°C)	k_i (W/m K)
0	2.26
-4	2.28
-10	2.32
-15	2.36
-20	2.40
-25	2.44
-30	2.48

*Based on the formula given by Sawada (1977).

Thermal conductivity of quartz k_q . Quartz is anisotropic, having a thermal conductivity parallel to the c-axis $k_{||}$ greater than the conductivity at right angles to this axis k_{\perp} . The thermal conductivity of a polycrystalline quartz aggregate of random orientation was determined using a geometric mean equation (Farouki 1981).

$$k_q = (k_{||} k_{\perp} k_{||})^{1/3} = k_{||}^{2/3} k_{\perp}^{1/3}.$$

Table 14. Thermal conductivity of quartz k_q *

Temperature (°C)	k_q (W/m K)
30	7.28
25	7.43
20	7.58
15	7.72
10	7.86
4	8.04
0	8.16
-4	8.29
-10	8.50
-20	8.84
-30	9.18

*Based on the tabulated data for $k_{||}$ and k_{\perp} in Touloukian et al. (1970).

The values of k_o and k_i at various temperatures were taken from the *Thermophysical Properties of Matter* (Touloukian et al. 1970). Where necessary these were linearly interpolated and then used in the geometric mean equation to calculate k_g at various temperatures. The results are given in Table 14 and these again were linearly interpolated to arrive at values of k_g for intermediate temperatures.

Thermal conductivity of the soil solids other than quartz, k_o . As suggested by Johansen (1975), the value of k_o was usually taken to be 2.0 W/m K irrespective of temperature. For coarse soils, however, having a quartz content less than 20% of the solids, some calculations were also made with k_o taken as 3.0 W/m K, a value assumed by Johansen for these soils.

Owing to the uncertainty in the value of k_o , there was no point in varying its input value with temperature. It should be noted, however, that the thermal conductivity of feldspar increases as the temperature increases. Feldspar may be the chief component other than quartz. This behavior of feldspar is exceptional among crystalline materials which usually show a decrease in the thermal conductivity as the temperature increases. Variations in k_o with temperature may therefore be somewhat damped.

Thermal structural value α' for dry soils. Where the Smith method is applied to dry soils, a suitable value of α' must be input.

Some program details

When the data for a particular soil were input, the soil had to be specified as being coarse or fine and its condition as being unfrozen or frozen, and unsaturated, saturated or dry. It had also to be specified as natural or crushed so that the appropriate equation in the Johansen method could be applied.

The thermal conductivity methods to be applied in each case are specified. For dry soils, the program automatically calculates the adjusted De Vries value from the value given by De Vries' method. For saturated soils, the computer program calculates the geometric mean value assuming a two-phase material (i.e. solid and water or ice).

Gemant method. This method is inapplicable when the moisture content is below a certain value corresponding to the adsorbed film water. The solids thermal conductivity obtained from the equation suggested by Gemant* was computed by

*This equation is $k_s = 5.84 + 0.33p$ W/m K where p is the percent of clay in the soil solids (see Farouki 1981).

the program and represented as CSOLIDS 2. This was for comparison and it was not used further.

De Vries method. While De Vries' method was not originally intended to apply to frozen soils, Penner (1970) applied it to two frozen clays with good results. It was decided for this analysis to apply De Vries to different conditions of frozen soils, i.e. saturated or unsaturated, and partially or completely frozen.

For saturated frozen soils containing some unfrozen water, this water may be considered as the continuous medium, so the following equation was used for the soil thermal conductivity k :

$$k = \frac{x_w k_w + F_i x_i k_i + F_s (1-n) k_s}{x_w + F_i x_i + F_s (1-n)}$$

The equation for such soils that are partially saturated is

$$k = \frac{x_w k_w + F_i x_i k_i + F_a x_a k_a + F_s (1-n) k_s}{x_w + F_i x_i + F_a x_a + F_s (1-n)}$$

If the frozen soils can be considered to have no unfrozen water content (e.g. frozen gravels or sands), the ice is taken to be the continuous medium, giving the equations

$$k = \frac{x_i k_i + F_s (1-n) k_s}{x_i + F_s (1-n)}$$

for saturated soils where $x_i = n$, the porosity, and

$$k = \frac{x_i k_i + F_a x_a k_a + F_s (1-n) k_s}{x_i + F_a x_a + F_s (1-n)}$$

for unsaturated soils.

In the above equations x represents the volume fraction of the soil component corresponding to its subscript (w for water, i for ice, a for air and s for soil solids). The 'F' values are given by

$$F_s = \frac{1}{3} \left\{ \frac{2}{1 + [(k_s/k_f) - 1]} 0.125 \right. \\ \left. + \frac{1}{1 + [(k_s/k_f) - 1]} 0.75 \right\}$$

$$F_i = \frac{1}{3} \left\{ \frac{2}{1 + [(k_i/k_f) - 1]} 0.125 \right. \\ \left. + \frac{1}{1 + [(k_i/k_f) - 1]} 0.75 \right\}$$

and

$$F_a = \frac{1}{3} \left\{ \frac{2}{1 + [(k_a/k_f) - 1] g_a} + \frac{1}{1 + [(k_a/k_f) - 1] g_c} \right\}$$

in which k_f is the thermal conductivity of the fluid continuous medium (unfrozen water or ice), and the shape factors g_a, g_c for the pore air are given by (De Vries 1963)

$$g_a = 0.333 - x_a \frac{(0.333 - 0.035)}{n}$$

for $0.09 \leq x_f \leq n$, or

$$g_a = 0.013 + 0.944 x_f$$

for $0 \leq x_f < 0.09$, and

$$g_c = 1 - 2 g_a$$

Mickley method. Although Mickley's method was originally derived for unfrozen soils it was also applied here to frozen soils simply by considering ice to occupy the place of water in the unit cube soil model (Mickley 1951).

McGaw method. In applying McGaw's conductance equation, the interfacial efficiency ϵ was taken as unity. A value for n_c (the volume of series fluid in unit soil volume) was required and this was calculated using

$$n_c = n(1-n)(0.304 - 0.09 \log k_s/k_w)$$

as suggested by McGaw (pers. comm.). If this equation, however, gave a value for n_c which was greater than nS_f , the magnitude of n_c was limited to nS_f . (The computer printout tables in Appendix B give the n_c value in the form NC .)

Van Rooyen method. The equation of Van Rooyen and Winterkorn (1959) is:

$$1/k = A 10^{-BS_f} + s \text{ cm } ^\circ\text{C/W}.$$

Where $A = 10^{a_1 - 0.44 \gamma_d^2}$,

$$B = b_1 - 5.5 \gamma_d$$

$$s = s_1 - s_2 \gamma_d$$

S_f = the degree of saturation of the soil
(fractional)

and

γ_d is its dry density (g/cm^3).

Based on the experimental data and analysis of Van Rooyen and Winterkorn (1959), the value of a_1 was taken as 3.55 and the following values for b_1 , s_1 and s_2 were chosen for the different soil types. For cohesionless soil:

$$b_1 = 16.18$$

$$s_2 = 47.5$$

$$s_1 = 200 - 94q \quad \text{for } q > 0.75$$

$$\text{or } s_1 = 435 - 407q \quad \text{for } 0.75 > q > 0.20$$

$$\text{or } s_1 = 353.6 \quad \text{for } q < 0.20$$

where q is the quartz content (fractional). For cohesive soils:

silts

$$b_1 \approx 5.6 \times 10^{-4} p + 9.58$$

$$s_2 \approx 134.6$$

$$s_1 \approx 202$$

clays

$$b_1 \approx 5.6 \times 10^{-4} p + 9.58$$

$$s_2 \approx 155$$

$$s_1 \approx 317$$

where p is the clay content (fractional).

Thus the quartz content for cohesionless soils and the clay content for cohesive soils are required as input data. If the latter was not known, the following rough average values for b_1 were used:

$$b_1 = 11.8 \text{ for silts}$$

$$\text{or } b_1 = 9.58 \text{ for clays.}$$

Van Rooyen is the only method which differentiates cohesive soils by subdividing them into clays and silts according to the general description of the soil.

Applicability of the methods

In this section the various methods for calculating thermal conductivity are tested to see under what

conditions their predictions agree with measured values of the thermal conductivity and to determine the extent of agreement. The deviations of the predicted values from the measured values are determined at different values of dry density and moisture content. This is done for soils that are unfrozen or frozen, coarse or fine, unsaturated, saturated or dry.

Applicability to unfrozen coarse soils

Figures 31 and 32 show the deviations given by the seven* applicable methods which were tested on data for unfrozen coarse soils and crushed rocks. These data are the result of measurements made by Kersten (1949) on Fairbanks sand, Lowell sand, Northway fine sand, Northway sand, standard Ottawa sand, graded Ottawa sand, Chena River gravel, crushed trap rock, crushed feldspar, crushed granite, crushed quartz and crushed fine quartz; by Johansen (pers. comm.) on sands SA1, SA2, SA4, SA8, SA13 and gravels GR1, GR6, GR7 and GR12, on crushed rocks PU1, PU5, PU6, PU7, PU9 and PU10; and by De Vries (1963) on Wageningen sand (data on these soils are given in Appendix A).

Kersten method. Figure 31a shows the deviations given by Kersten for unfrozen sands. As Kersten himself noted his relevant equation does not apply to the sands he tested that had a low quartz composition, i.e. Northway sand and Northway fine sand. It gives values that are too high, with deviations up to 150% for most of the saturation range (see Appendix Tables B1 and B2). The Kersten equation also gives some high deviations (55%) for several Johansen sands with intermediate quartz content (Fig. 31a).

For the sands tested by Kersten that have medium or high quartz content (Fairbanks sand, Lowell sand, standard and graded Ottawa sand), the Kersten equation generally gives good agreement within $\pm 20\%$, many of the deviations being negative. Such an agreement may be expected because Kersten fit his equation to these experimental data. However the Kersten equation underpredicts when applied to the data obtained by other workers on sands having high quartz content. Thus for the sands of Johansen and De Vries having high quartz content (Johansen sands SA4 and SA13, De Vries Wageningen sand [$q > 0.65$]), the Kersten equation gives many deviations in the range -25 to -50% at varied values of S_r .

The deviations resulting from the application of the Kersten equation to unfrozen gravels and crushed

rocks are shown in Figure 32a. The predictions given by the Kersten equation show similar trends with these gravels and crushed rocks as with the sands considered above. Thus Kersten gives predictions that are much too high for the low-quartz gravel GR7 (Table B3) and the deviations remain substantial at high values of S_r . In the same manner, for all Johansen's crushed rocks which have low quartz content (PU1, PU5, PU6, PU7, PU9, PU10), Kersten gives predictions that are much too high, the deviations reaching 144% (e.g. Table B4). Kersten also gives unacceptably high predictions for the low-quartz crushed rocks tested by him, i.e. crushed trap rock, crushed feldspar and crushed granite (e.g. Table B5). This confirms that Kersten should not be applied to materials with low quartz content.

For the medium-quartz gravels (Chena River gravel and Johansen's GR1, GR6 and GR12 gravels [$0.40 \leq q \leq 0.65$]), Kersten applies well to his own Chena River gravel and to one sample of Johansen's GR12 gravel, but gives some unacceptably high deviations for Johansen's GR1 and GR6 gravels.

As with the high-quartz sands, Kersten gives predictions that are too low for the crushed quartz materials tested by Kersten himself.

To summarize, when applied to coarse soils or soil materials, Kersten overpredicts for those with a low quartz content while it underpredicts for those with a high quartz content. In either case the deviations are too large to be acceptable. Use of the Kersten method should therefore be limited to coarse soils with intermediate quartz content, say around 60% of the soil solids. The expected deviations would then generally be within $\pm 25\%$, though many may be larger and even unacceptable, particularly for gravels.

Johansen method. Johansen gives good predictions (within $\pm 25\%$) for coarse soils and crushed rocks of varied quartz content at S_r values above 0.2 (Fig. 31b and 32b). Between $S_r = 0.1$ and $S_r = 0.2$ the predictions are somewhat worse, the deviations showing a marked negative bias extending to -40%. This could be due to the effect of moisture migration, which Johansen does not take into account and which could appreciably increase the measured thermal conductivity. Below about $S_r = 0.1$, Johansen gives large deviations, chiefly negative.

De Vries method. At degrees of saturation greater than about 0.2, De Vries gives deviations within $\pm 20\%$ for sands (Fig. 31c), and within $\pm 30\%$ for gravels and crushed rocks (Fig. 32c). There is a tendency for mainly negative deviations in the range $S_r = 0.1$ to 0.3, but it is less marked than that shown by Johansen, as may be expected because De Vries attempts to take the effect of moisture migration into account.

*These are Kersten, Johansen, De Vries, Gemant, Mickley, McGaw and Van Rooyen which apply to unsaturated (or saturated) unfrozen soils. For saturated soils two additional methods apply: modified resistor and Kunii-Smith.

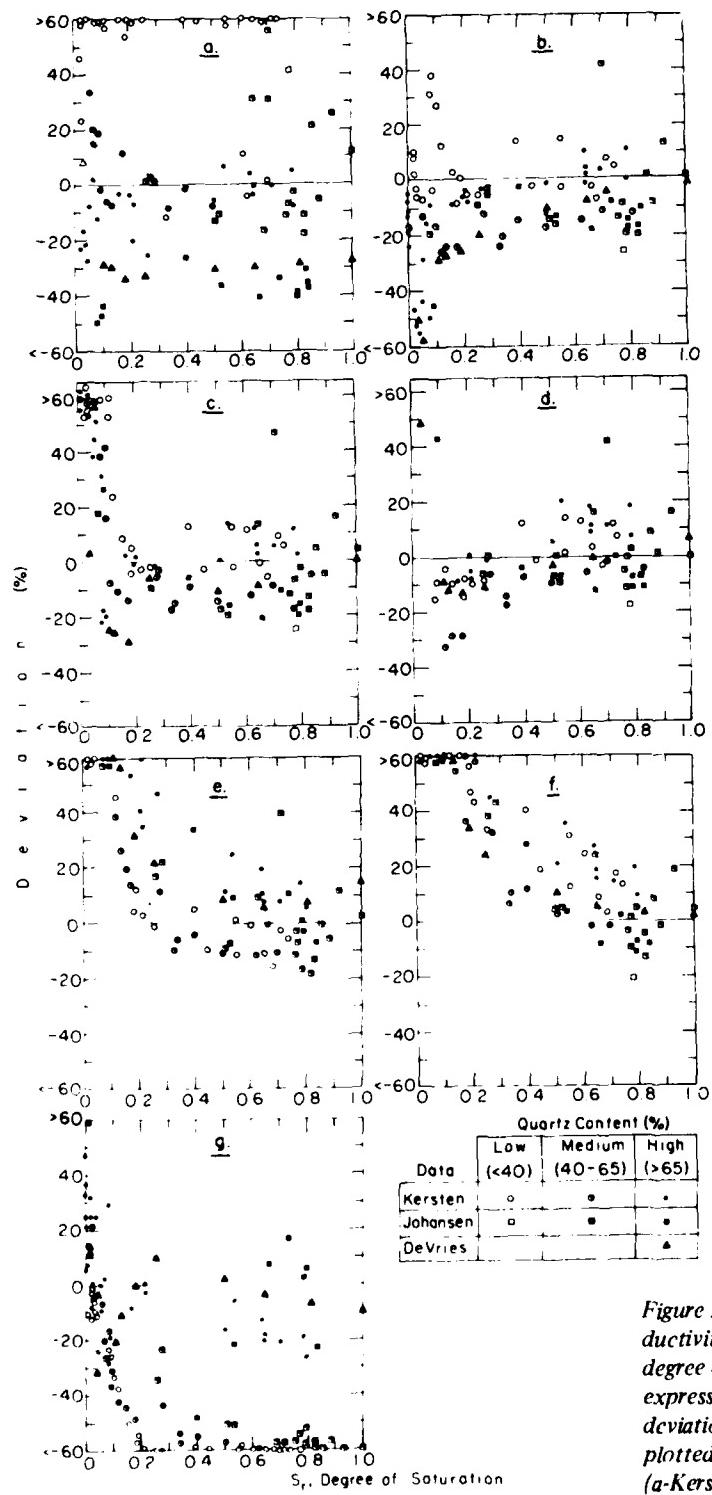


Figure 31. Deviations of calculated thermal conductivity from measured thermal conductivity vs degree of saturation of unfrozen sands. Deviations expressed as a percentage of the measured values; deviations outside the limits of $\pm 60\%$ have been plotted directly on top or bottom horizontal scales (a-Kersten, b-Johansen, c-De Vries, d-Gemant, e-Mickley, f-McGaw, g-Van Rooyen).

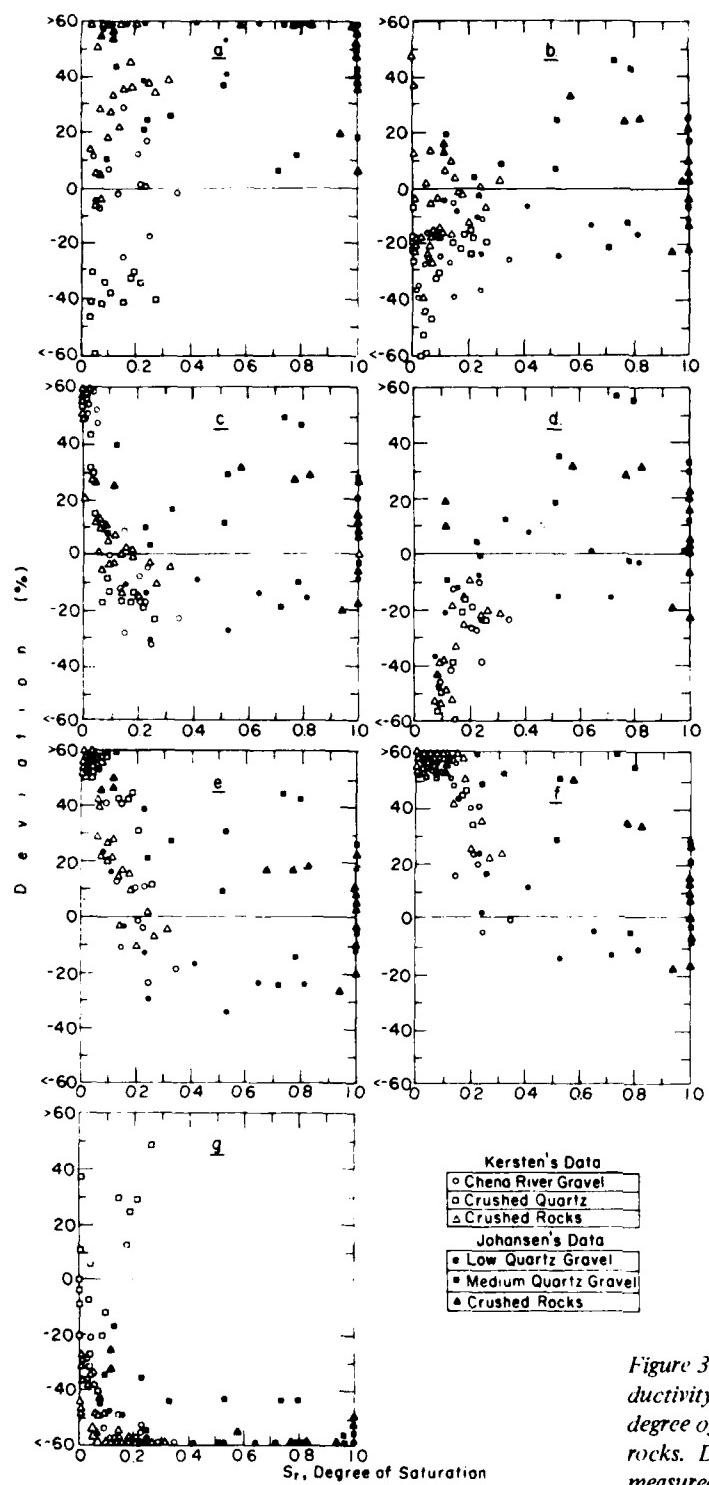


Figure 32. Deviations of calculated thermal conductivity from measured thermal conductivity vs degree of saturation of unfrozen gravels and crushed rocks. Deviations expressed as a percentage of the measured values; deviations outside the limits of $\pm 60\%$ have been plotted directly on top or bottom horizontal scales (a-Kersten, b-Johansen, c-De Vries, d-Gemant, e-Mickley, f-McGaw, g-Van Rooyen).

Gemant method. By its nature, Gemant is inapplicable below a certain moisture content. It begins to give reasonable results above an S_f value of roughly 0.2 where the deviations are generally within $\pm 20\%$ for sands (Fig. 31d) or within -25 and 35% for gravels and crushed rocks (Fig. 32d). Similarly to Johansen, Gemant shows a marked negative bias in the range $S_f = 0.1$ to 0.3, which may be attributed to the effect of moisture migration.

Mickley and McGaw methods. These two methods are generally similar in their predictions. In effect both assume good particle-to-particle contact, implying very efficient thermal transfer. Obviously this cannot be the case in the dry condition which therefore gives rise to very large positive deviations from both methods. The presence of some water soon improves the interfacial efficiency and decreases the deviations. However, these deviations remain positive, masking any contribution to heat transfer from moisture migration, which would tend to give negative deviations. For sands Figures 31e and 31f show the marked trend for the deviations to decrease as S_f increases. For gravels and crushed rocks this trend, while still evident, is less consistent at high S_f values (Fig. 32e and 32f).

In the case of sands McGaw shows a roughly linear decrease in the deviations as S_f increases (Fig. 31f). Its predictions can be improved by introducing a suitable value for the interfacial efficiency factor ϵ . Such a value would be appreciably less than unity in the nearly dry condition, increasing linearly to unity at high S_f values.

For sands Mickley gives good agreement (within $\pm 25\%$, roughly) for S_f values above 0.45, while McGaw does not do so until S_f is greater than about 0.6. With the gravels the threshold of S_f is lower, being about 0.25 for Mickley and 0.5 for McGaw. In the case of the crushed rocks, this threshold is slightly lower still, suggesting that the surface characteristics of crushed materials provide better contact efficiency (Farouki 1981).

Van Rooyen method. Van Rooyen's equation appears to be generally applicable to sands having a high quartz content at S_f values above 0.1 but not for lower S_f values (Fig. 31g). It also appears that Van Rooyen applies well (in fact better than the other six methods) below an S_f value of 0.1 for sands with low or medium quartz content. In this region, if we exclude the nearly dry condition ($S_f < 0.015$), the deviations for the high-quartz ($q > 0.65$) sands are not very large, lying within $\pm 35\%$, which is still much better than Johansen and the other methods.

For gravels and particularly for low-quartz ($q < 0.40$) crushed rocks, the Van Rooyen equation does not generally apply well. It does, however, give some reasonable predictions for crushed quartz in the range $S_f = 0.01$ to 0.2 (Fig. 32g).

Summary. Above an S_f of 0.2 Johansen generally gives the best agreement (within $\pm 25\%$), while De Vries and Gemant are close behind. Mickley and McGaw give good predictions at higher S_f values of 0.45 and 0.6 respectively. Under the stipulated conditions these five methods are applicable to coarse materials of high, intermediate or low quartz content. This is because they take the solids thermal conductivity k_s into account which Kersten's method does not.

In the range of S_f values from 0.1 to 0.2, De Vries appears to be the best method, giving deviations between 10 and -30% . Johansen gives a wider range of variation, between 20 and -40% , thus showing more extensive negative deviations.

Below an S_f value of 0.1 and extending to around $S_f = 0.015$, Van Rooyen gives the best predictions for sands, but some of the deviations are rather extensive (up to $\pm 35\%$). For the gravels and low-quartz crushed rocks, however, at such low S_f values, Van Rooyen does not apply well nor does any one of the other methods.

Effect of variation in k_o . The effect on the soil thermal conductivity of variation in k_o is greatest for materials with low quartz content, which implies a high content of the other minerals. Such materials may sometimes have a larger k_o value than the 2.0 W/m K assumed in all the calculations on which Figures 31 and 32 (except 31a and 32a) are based. Johansen (1975) suggested that for his coarse materials with quartz content less than 20%, a k_o value of 3.0 W/m K should be used. Figures 33 and 34 show the deviations resulting from such an assumption as compared with the previous choice of 2.0 W/m K for k_o . Thus these figures show the effect of uncertainties in knowledge of k_o , this effect being greatest for materials with a very low quartz content (less than 20%). These materials include Kersten's two Northway sands, 50% of which is derived from igneous rocks, and Johansen's crushed rocks PU1, PU7, PU9, and PU10. They also include Johansen's sand SA10 and gravel GR7.

Figures 33a and 34a apply to the Johansen method. As expected, the deviations corresponding to a value of 3.0 W/m K for k_o shift upwards compared to the deviations corresponding to $k_o = 2.0$ W/m K. The ranges covered by these deviations are 10 to 40% for the sands (Fig. 33a) and 20 to $\sim 10\%$ for gravels and crushed rocks corresponding to S_f values above

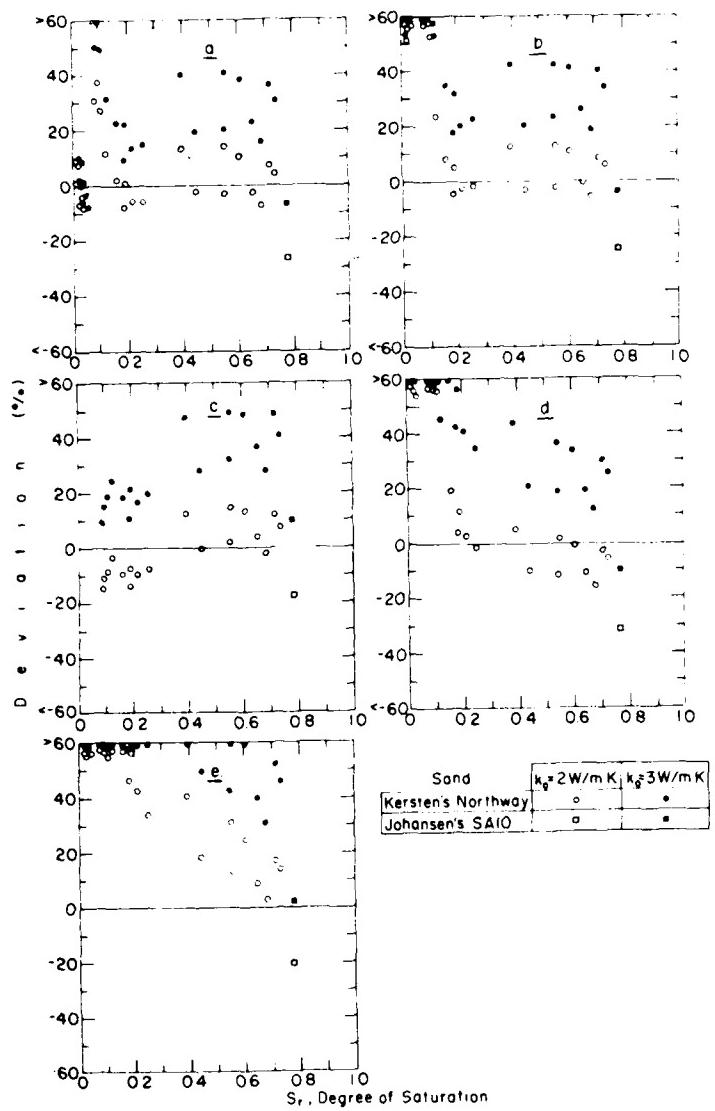


Figure 33. Effect of variation in thermal conductivity k_o of soil solids other than quartz on the deviations of calculated thermal conductivity from measured thermal conductivity of unfrozen low-quartz sands ($q < 0.20$) at various degrees of saturation. Deviations expressed as a percentage of the measured value; deviations outside the limits of $\pm 60\%$ have been plotted directly on top or bottom horizontal scales (a-Johansen, b-De Vries, c-Gemant, d-Mickley, e-McGaw).

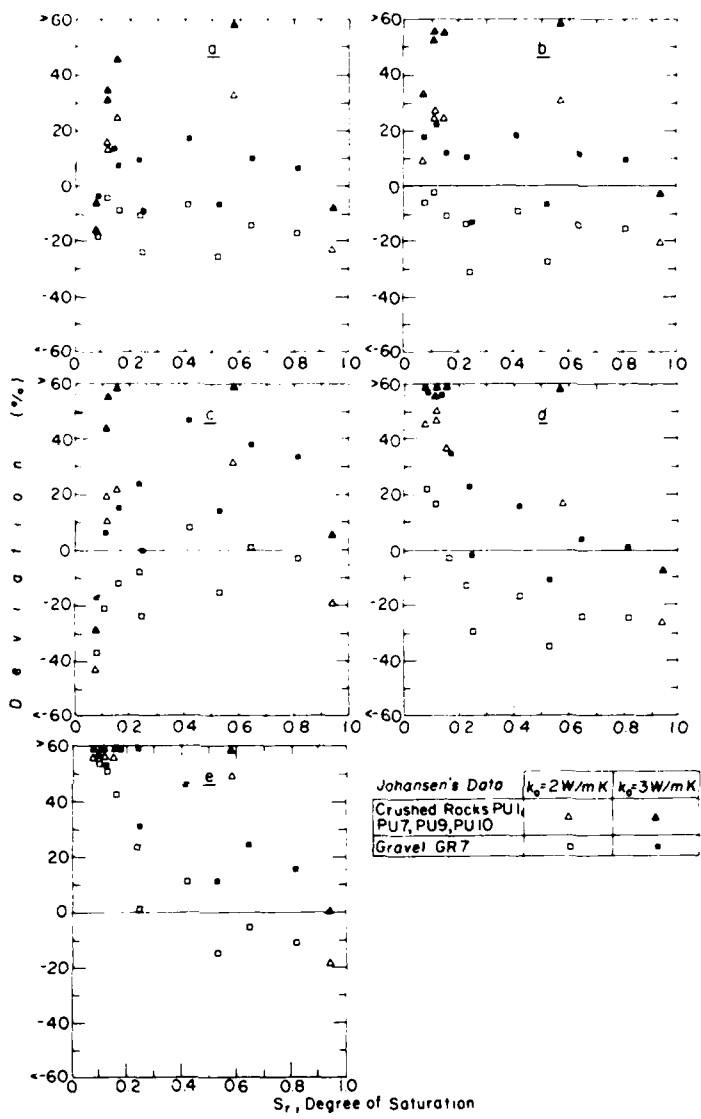


Figure 34. Effect of variation in thermal conductivity k_o of soil solids other than quartz on the deviations of calculated thermal conductivity from measured thermal conductivity of unfrozen low-quartz crushed rocks and gravels ($q < 0.20$) at various degrees of saturation. Deviations expressed as a percentage of the measured value; deviations outside the limits of $\pm 60\%$ have been plotted directly on top or bottom horizontal scales (a-Johansen, b-De Vries, c-Gemant, d-Mickley, e-McGaw).

0.2. With a k_s value of 3.0 W/m K and S_f greater than 0.2, De Vries gives deviations in the range 40 to -5% for the sands (Fig. 33b) and in the range 20 to -10% for the crushed rocks (Fig. 34b). For similar conditions, Gemant gives deviations lying in the range 50 to 10% for the sands (Fig. 33c) and 50 to 5% for the crushed rocks (Fig. 34c). Mickley produces deviations in the range 35 to -10% for the sands at S_f values above 0.5 (Fig. 33d), while for the crushed rocks the deviations lie in the range 35 to -10% at S_f values above 0.3 (Fig. 34d). McGaw gives deviations for the sands that are too high (Fig. 33e), but for the crushed rocks (Fig. 34e) the deviations are reasonable at S_f values above 0.5.

Applicability to a sandy silt-clay. Where a sand contains a large amount of silt-clay, such as Kersten's Dakota sandy loam with a silt-clay content of 31%, the Kersten equation overestimates considerably at S_f values below about 0.45, giving deviations up to 144% (see Table B6) in spite of the fact that this soil has a medium quartz content. At higher S_f values the Kersten method provides reasonable agreement as does each of the other six methods except Van Rooyen. In particular, Johansen and De Vries, while overestimating at low S_f values, give reasonable agreement from S_f values of 0.24 and 0.32 respectively. Van Rooyen gives good agreement at $S_f \geq 0.24$ and Gemant is generally applicable throughout the range of S_f values.

Applicability to saturated coarse soils. For saturated coarse soils or soil materials (e.g. crushed rocks), nine methods may be applied which, in addition to the seven methods discussed above, include modified resistor and Kunii-Smith. Representative deviations produced by all these methods are given in Tables B7-B11 and also for the seven methods in Figures 31 and 32 for the $S_f = 1.0$ points.

Kersten gives unacceptably high deviations for the saturated low-quartz crushed rocks (Johansen's PU1, PU5, PU6, PU7, PU9 and PU10) and also for the low-quartz gravel GR7. Excepting Van Rooyen, all the other methods generally give good agreement (within $\pm 25\%$) for these materials as well as for the medium-quartz saturated gravels (GR1, GR3 and GR6) and the SA2 sand of Johansen. For the GR1 and GR6 gravels, Kersten persists in giving some unacceptably high deviations which, however, are not as bad as in the case of the low-quartz materials. While for the other medium-quartz gravel (GR3) Kersten gives acceptable deviations, these are wider than those given by the other methods. Similarly, while Kersten gives good agreement for the medium-quartz SA2 sand, six other methods give even better agreement. These are Johansen, De Vries, Kunii-Smith, modified resistor, McGaw and Gemant, any

of which is preferable to Kersten for calculating the thermal conductivity of saturated coarse soils.

Applicability of methods to unfrozen fine soils

Figure 35 shows the deviations obtained with the seven applicable methods (Kersten, Johansen, De Vries, Gemant, Mickley, McGaw and Van Rooyen) for some unfrozen fine soils at various values of S_f . The evaluated data were for soils tested by Kersten, i.e. Healy clay, Fairbanks silty clay loam, Fairbanks silt loam, Northway silt loam and Ramsey sandy loam. The quartz contents of these soils were known and varied from 0.015 to 0.641* (as a fraction of the total solids content). In the computer program, k_s was set at 2.0 W/m K. In addition to Kersten's soils, data for Russian chernozem given by Kolyasev and Gupalo (1958) were evaluated as well as data given by Mickley (1951) and by Reno and Winterkorn (1967) on fine soils, k_s being taken as 2.0 W/m K.

Kersten method. Below an S_f of about 0.3, Kersten's equation for unfrozen fine soils gives deviations that are either too high, particularly for the chernozem soil, or too low for Kersten's own soils. From $S_f = 0.3$ to full saturation, the Kersten equation gives deviations that are scattered between 35 and -35%, with some of the highest deviations occurring for Kersten's own soils (Fig. 35a), particularly for his low-quartz Northway silt loam.

Johansen method. At low values of S_f , Johansen gives better agreement than Kersten. Nevertheless, Johansen still gives some excessive deviations below $S_f = 0.2$, these generally varying between 20 and -45%. At higher values of S_f Johansen gives deviations in the range 35 to -35%, roughly, which is similar behavior to Kersten (see Fig. 35b).

Above $S_f = 0.2$ Johansen gives positive deviations for most of Kersten's samples, but it tends to give negative deviations for the other soils, possibly because the assumed value of k_s for these (i.e. 2.0 W/m K) may generally be too low.

Other methods. De Vries gives deviations that are much too high at $S_f < 0.3$, but even above this S_f value the deviations for Kersten's samples continue to be unacceptably high (about 50%) and remain so at near full saturation (Fig. 35c). The range of deviations for Kersten's soils is from 50 to -3%, indicating an overprediction for these soils which is similar to that given by Johansen, but greater. For the other soils De Vries gives deviations in the range of 10 to -35% which is again similar to Johansen.

At $S_f > 0.3$ Gemant, Mickley and McGaw show trends generally similar to De Vries, but the deviation

*All these Kersten soils contain a medium amount of quartz ($0.40 < q < 0.65$) except for the low-quartz Healy clay and Northway silt loam.

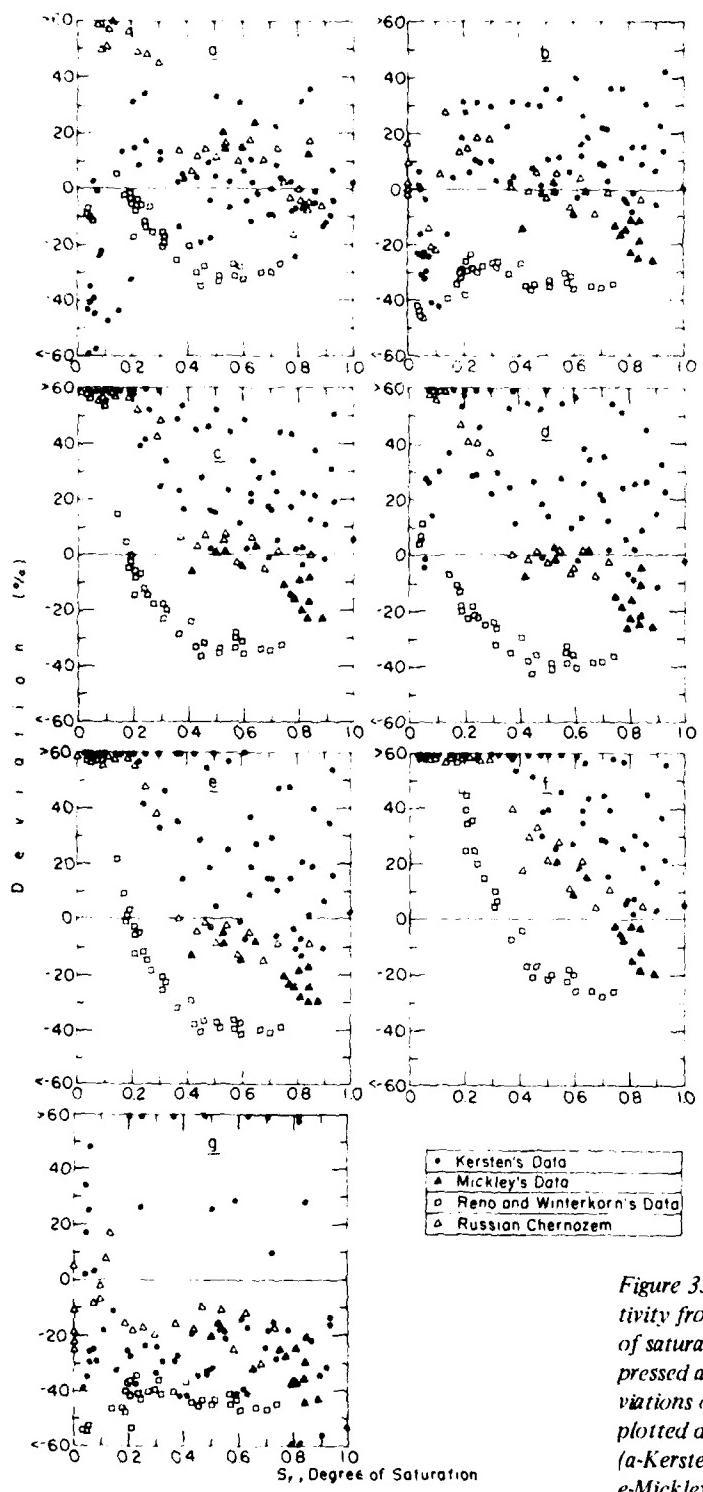


Figure 35. Deviations of calculated thermal conductivity from measured thermal conductivity vs degree of saturation of unfrozen fine soils. Deviations expressed as a percentage of the measured value; deviations outside the limits of $\pm 60\%$ have been plotted directly on top or bottom horizontal scales (a-Kersten, b-Johansen, c-De Vries, d-Gemant, e-Mickley, f-McGaw, g-Van Rooyen).

ranges are larger (see Fig. 35d-f). Geman, Mickley and McGaw give mostly positive deviations for Kersten's soils and Geman and Mickley give mostly negative deviations for the other soils.

As with coarse soils McGaw shows a particularly noticeable trend of decreasing deviation as the S_r value for a particular type of fine soil increases (Fig. 35f). This again suggests that application of a proper value of the interfacial efficiency factor ϵ would give better agreement. As proposed in an earlier section, ϵ would have a small value at low S_r values and it would increase linearly to near unity as full saturation is approached.

Van Rooyen gives a scatter of deviations that is different from all the other methods (Fig. 35g). It produces mostly negative deviations, except for some of Kersten's samples, and the range of deviations is unacceptably extensive. However, for the Russian chernozem, Van Rooyen gives good predictions for a complete range of S_r values from the dry condition to near saturation. For such a range Johansen is the only other method which produces reasonable predictions for the chernozem.

Summary. Kersten should not be used below $S_r = 0.3$ as it would give excessive deviations. Above this S_r value both Kersten and Johansen give deviations within the range 35 to -35%, so that either method is equally applicable. Below $S_r = 0.3$, only Johansen continues to give predictions in this range and does so until S_r reaches about 0.2. It is expected that the accuracy of Johansen would improve with knowledge of a proper value of k_s .

Use of the Johansen method as it stands for S_r values lower than 0.2 may give more excessive deviations. From the trends in the deviations shown in Figure 35b the following suggested scheme could be applied:

1. In the range of $0.1 < S_r < 0.2$, Johansen gives deviations between 30 and -40%, i.e. a rough "average" of -5%. The values given by Johansen could therefore be increased by 5%.
2. In the range $0 < S_r < 0.1$, Johansen gives deviations between 15 and -45%, i.e. a rough "average" of -15%. The values given by Johansen could therefore be increased by 15%.

These specific suggestions are tentative, being based on fairly limited data.

Apart from Kersten and Johansen, the other five methods generally show more extensive deviations, many of which are unacceptable so that use of these methods is not recommended.

Comparison of predictions with tabulated Soviet values. The predictions of the various methods are compared in Table B12 with the values for Soviet clay soils tabulated in the U.S.S.R. Building Code

(1960). For these calculations, the value of k_s was assumed to be 2.0 W/m K but it could be greater. Nearly all of the values given by the methods are lower than the U.S.S.R. Code values. Kersten gives differences varying between -5 and -33%, these differences tending to increase as the dry density increases at a given moisture content. Johansen and the other methods generally give lower negative deviations.

Applicability to saturated unfrozen fine soils. The applicability of nine methods was tested for Kersten's Healy clay (Table B13), Penner's Leda clay (Table B14) and the Dames & Moore (1973) clay or silt (Tables B15 and B16).

The Kersten method gave good predictions, except for Penner's Leda clay, for which it gave values that were too high. All the other methods, except Van Rooyen, gave good predictions overall. Thus seven methods are more or less equally applicable. These are Johansen, De Vries, modified resistor, Kunii-Smith, Mickley, McGaw and Geman. The Johansen method is suggested as the first choice.

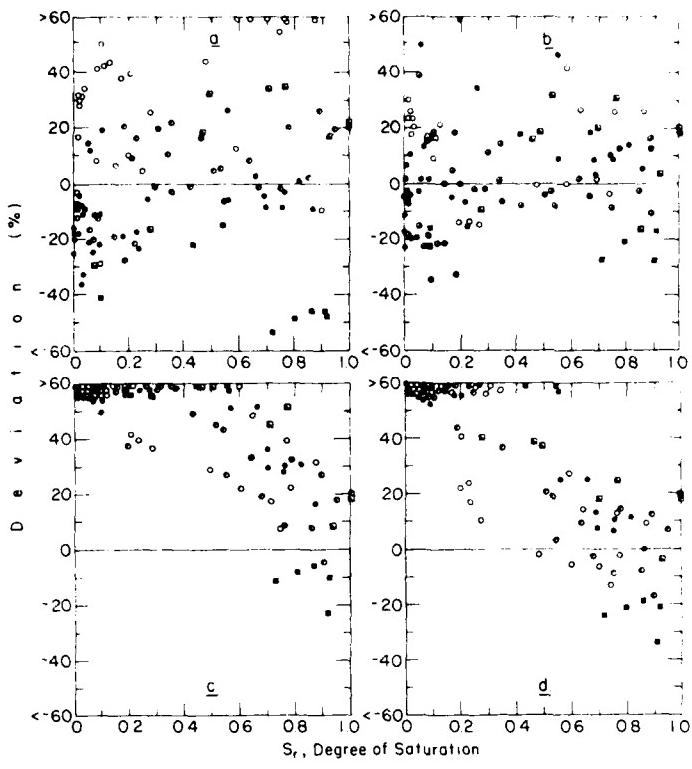
Applicability of methods to frozen coarse soils

For frozen soils only four methods could be used, i.e. Kersten, Johansen, De Vries and Mickley. The predictions of these methods were compared with the thermal conductivity measurements made by Kersten and Johansen on frozen sands, gravels and crushed rocks (Fig. 36 and 37).

Kersten method. As may be seen from Figures 36a and 37a, Kersten gives predictions that are generally much too high for frozen sands, gravels and crushed rocks having low quartz content. On the other hand it gives predictions that are too low for sands or crushed materials with high quartz content. This is in line with the trends shown by Kersten for unfrozen soils.

For frozen materials with intermediate quartz content, Kersten shows conflicting trends. For the sands it gives deviations lying between 35 and -25% at S_r values between 0.2 and 0.6. At S_r values greater than 0.6, the range of deviations is narrower, between 35 and -10%. While, as may be expected, the Kersten equation gives good agreement for Kersten's own Chena River gravel, there are inconsistencies and wide divergencies for Johansen's medium-quartz gravels GR1 and GR6. Thus, even for such materials, Kersten should be used with caution and large deviations expected.

Johansen method. Above an S_r value of 0.1 Johansen generally gives good or adequate predictions (within $\pm 35\%$) for frozen sands, gravels and crushed rocks of any quartz content (Fig. 36b and 37b). However there are some exceptions, such as



Quartz Content (%)			
	Low (<40)	Medium (40-65)	High (>65)
Kersten	○	●	●
Johansen		■	■

Figure 36. Deviations of calculated thermal conductivity from measured thermal conductivity vs degree of saturation of frozen sands. Deviations expressed as a percentage of the measured value; deviations outside the limits of $\pm 60\%$ have been plotted directly on top or bottom horizontal scales (a-Kersten, b-Johansen, c-De Vries, d-Mickley).

for several crushed rocks, where deviations that are too large occur at high S_r values.

De Vries method. De Vries generally gives good agreement above an S_r value of 0.8 but even so there are some exceptions (Fig. 36c and 37c).

Mickley method. For gravels and crushed rocks Mickley appears to give good agreement at S_r values greater than about 0.3 (Fig. 37d), while for sands such agreement is not obtained until S_r is greater than 0.6 (Fig. 36d). As with the other methods there are exceptions.

Summary. Johansen is the method giving the best agreement and is generally applicable from an S_r of 0.1 to higher values. De Vries and Mickley apply with good agreement at high S_r values, greater than about 0.8 and 0.6 respectively. Kersten should be used only with medium-quartz materials and,

even then, with caution, as it may give large deviations.

Effect of variation in k_o . As with the unfrozen coarse soils, calculations were made to determine the effect of setting k_o equal to 3.0 W/m K instead of 2.0 W/m K (2.0 W/m K was used to obtain Figures 36 and 37 [except 36a and 37a]). This procedure was carried out for the frozen sands and crushed rocks having a quartz content of less than 20% (Kersten's Northway sand and Northway fine sand; Johansen's sand SA10, gravel GR7 and crushed rocks PU1, PU7, PU9 and PU10). The effect of the variation in k_o is shown in Figures 38 and 39. The sensitivity to k_o of the thermal conductivity from each of the three relevant methods is evident from these figures.

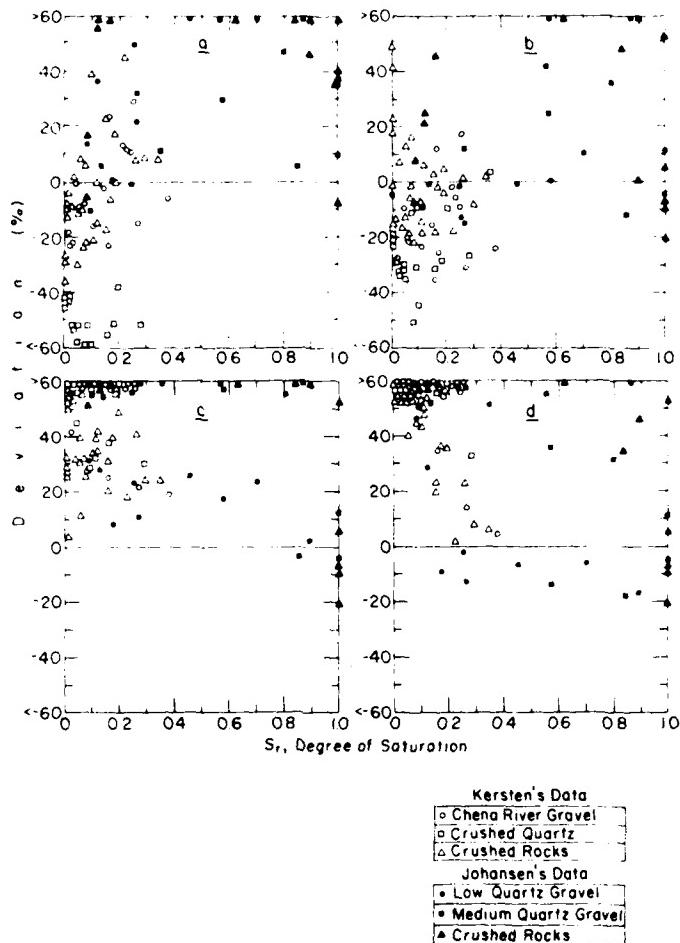


Figure 37. Deviations of calculated thermal conductivity from measured thermal conductivity vs degree of saturation of frozen gravels and crushed rocks. Deviations expressed as a percentage of the measured value; deviations outside the limits of $\pm 60\%$ have been plotted directly on top or bottom horizontal scales (a-Kersten, b-Johansen, c-De Vries, d-Mickley).

Consider the deviations in Figures 38 and 39 corresponding to a k_b value of 3.0 W/m K . For the frozen sands (Fig. 38a), Johansen gives deviations within the range 40 to 0% provided S_r is less than 0.6. At larger S_r values the deviations become too high. For the frozen crushed rocks at S_r values above 0.2, Johansen gives deviations within the range 25 to -5% (Fig. 39a). De Vries gives deviations that are too high for sands at all S_r values (Fig. 38b) but for the crushed rocks they become reasonable (25 to -5%) at S_r values exceeding 0.9 (Fig. 39b). Mickley also gives deviations that are too high for sands (Fig. 38c) but generally reasonable deviations are obtained for the crushed rocks at S_r values above 0.4 (Fig. 39c).

Applicability to a frozen sandy silt-clay.

Where the coarse soil contains a large amount of silt-clay, as in the case of Kersten's Dakota sandy loam, the above conclusions appear to hold. As Table B17 shows, Kersten gives good agreement for this medium-quartz soil, but Johansen gives even better agreement. Again De Vries and Mickley give good predictions only at high S_r values.

Applicability to saturated frozen coarse soils.

For the saturated condition, in addition to the four methods used above, modified resistor and Kunii-Smith may be used. Tables B18-B22 give a representative picture of the resulting deviations, as do Figures 36 and 37 for the previous four methods (see values for $S_r = 1.0$).

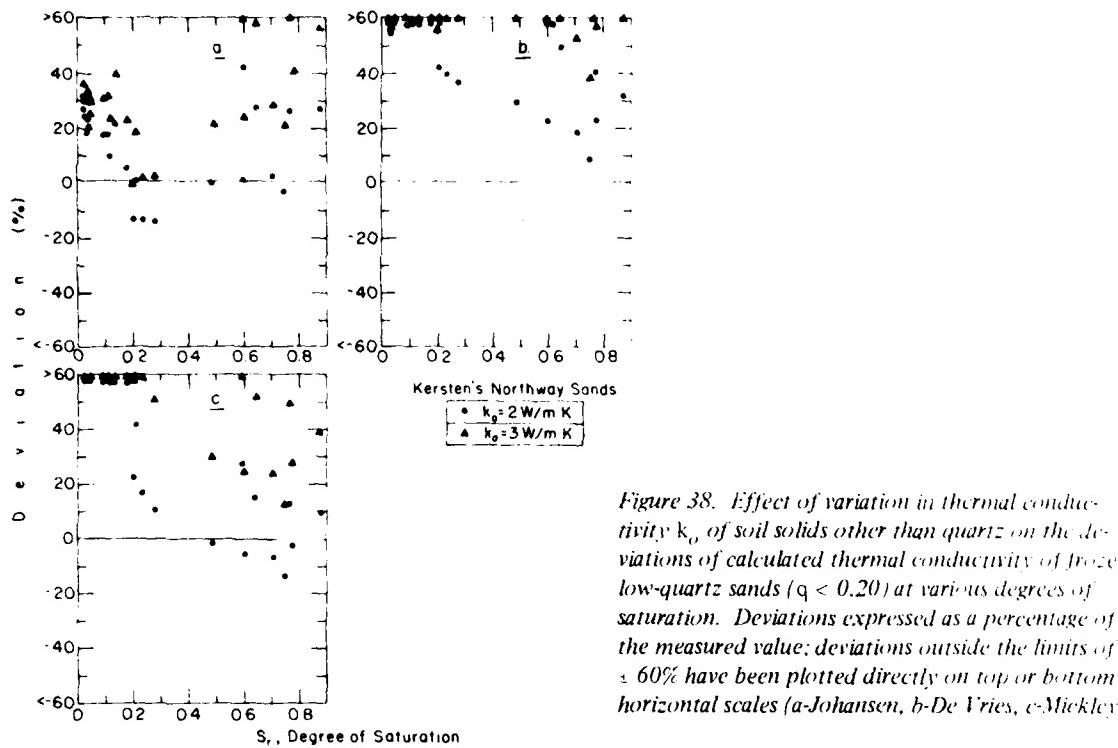


Figure 38. Effect of variation in thermal conductivity k_o of soil solids other than quartz on the deviations of calculated thermal conductivity of frozen low-quartz sands ($q < 0.20$) at various degrees of saturation. Deviations expressed as a percentage of the measured value; deviations outside the limits of $\pm 60\%$ have been plotted directly on top or bottom horizontal scales (a-Johansen, b-De Vries, c-Mickley).

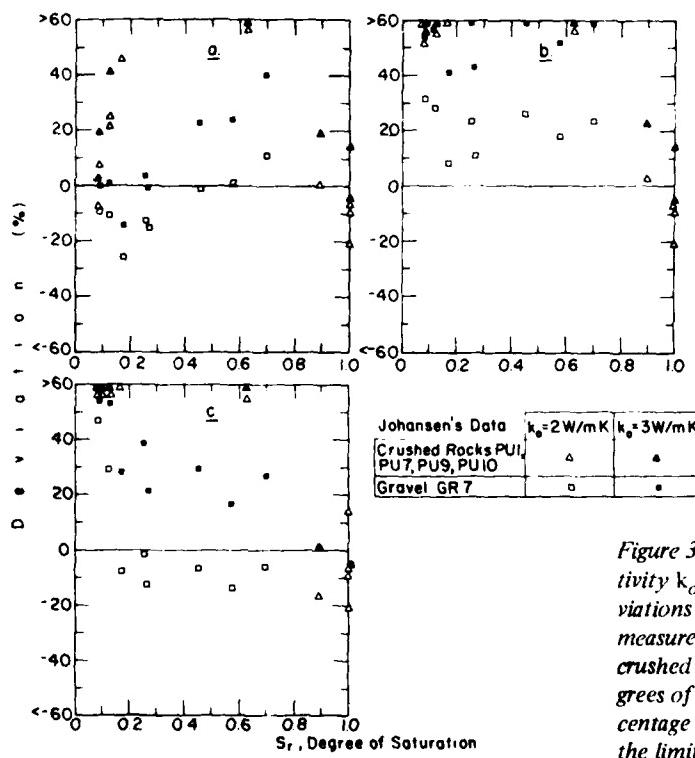


Figure 39. Effect of variation in thermal conductivity k_o of soil solids other than quartz on the deviations of calculated thermal conductivity from measured thermal conductivity of frozen low-quartz crushed rocks and gravels ($q < 0.20$) at various degrees of saturation. Deviations expressed as a percentage of the measured value; deviations outside the limits of $\pm 60\%$ have been plotted directly on top or bottom horizontal scales (a-Johansen, b-De Vries, c-Mickley).

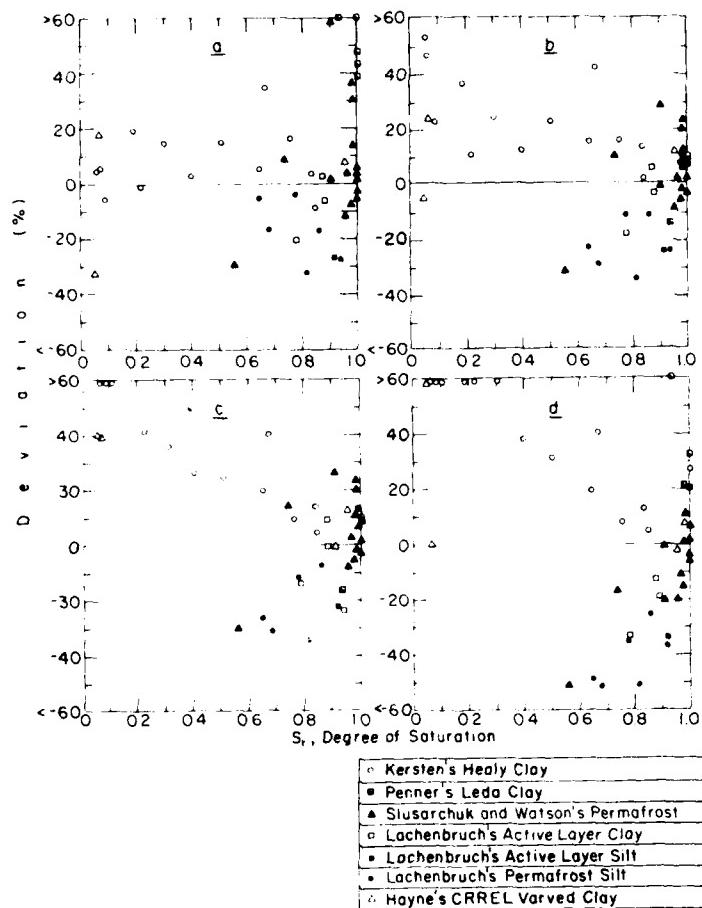


Figure 40. Deviations of calculated thermal conductivity from measured thermal conductivity vs degree of saturation of frozen fine soils. Deviations expressed as a percentage of the measured value; deviations outside the limits of $\pm 60\%$ have been plotted directly on top or bottom horizontal scales (a-Kersten, b-Johansen, c-De Vries, d-Mickley).

Kersten gives predictions that are too high for the low-quartz crushed rocks (except PU7) and high even for the medium-quartz gravel GR1. The best methods to use are Johansen, De Vries, Mickley, modified resistor or Kunit-Smith, any of which generally gives good agreement (except for crushed rock PU6 which is anomalous).

Applicability of methods to frozen fine soils

The four methods applicable to unsaturated frozen soils (Kersten, Johansen, De Vries and Mickley) were tested on Kersten's Healy clay, the CRREL varved clay (Haynes et al. 1980), the active layer silt or clay and the permafrost silt of Lachenbruch (pers. comm.), Penner's (1970) Leda clay, and the undisturbed permafrost of Slusarchuk and Watson (1975). The value of k_s was either that given by the experimenter or, if not known, taken as 2.0

W/m K. In the case of the Leda clay of Penner (1970), the unfrozen water content (UWC) was obtained from his data. For the other soils UWC was taken as zero or a suitable value assumed. The deviations resulting from the application of the four methods are shown in Figure 40.

From these limited results, it is seen that Kersten gives good predictions (generally between 20 and -35% deviations) up to an S_r value of about 0.9, above which it overpredicts considerably for Penner's Leda clay and for Slusarchuk and Watson's permafrost. While Johansen shows a tendency to overpredict below an S_r value of 0.1, it otherwise gives good or adequate agreement (generally within $\pm 35\%$) up to and including full saturation. Above an S_r value of about 0.4, De Vries gives values very similar to Johansen. Mickley is the worst predictor of all and only gives reasonable results at high values of

S_r (above 0.8 roughly) but even then shows some excessive deviations.

Calculations were made using the data available for soils 4 through 10 of Penner et al. (1975). These were silt-clay materials with little or no sand, except for soil 7 which contained about 40% sand-gravel (but was nevertheless a plastic material). The actual values of k_s are unknown so a value of 3.0 W/m K was assumed. This introduces some uncertainty into the calculated deviations which are given in Tables B23-B29. Three values of UWC were used in each set of calculations, i.e. 0.0, 0.05 and 0.10. Again there is uncertainty about the proper UWC value; UWC will vary from soil to soil depending particularly on the soil's specific surface area. However the UWC value is expected in every case to be greater than zero at -5°C, the temperature to which the measured values are applicable.

Kersten is of course independent of k_s and the UWC value. Tables B23-B29 show that Kersten gives agreement which varies from very good to barely satisfactory (up to 36% deviation). In the case of Johansen, excepting soils 8 and 10, the agreement is generally good, providing the UWC value is taken as 0.05 or 0.10. This shows the importance of assuming a proper value for UWC in the Johansen method. Soils 8 and 10 in fact contain the highest proportion of clay-sized material (~ 54%). Thus the assumption of a value of UWC greater than 0.10 for these soils should improve agreement by reducing the calculated values.

De Vries is particularly sensitive to the UWC value, as is evident from Tables B23-B29. Apart from soils 8 and 10, De Vries gives good agreement, with the assumption of a certain value of UWC. Mickley, on the other hand, shows little sensitivity to the UWC value, but gives good agreement at S_r values above 0.5, excepting again soils 8 and 10.

Tables B30-B32 show the deviations obtained by the four methods for Kersten's soils: Fairbanks silty clay loam, Northway silt loam and Ramsey sandy loam. The deviations were calculated at a UWC value of zero, but in the case of the Fairbanks silty clay loam calculations were also made at a UWC value of 0.06 to determine the resulting effect. Kersten generally gives good agreement for these loams, as may be expected because the Kersten equation was fitted to these data, but it gives some high values (around 25%) for Northway silt loam (this loam has low quartz content, whereas the others have medium quartz content). Johansen generally gives values that are too high, although predictions are acceptable for some samples of Ramsey sandy loam and for samples of Northway silt loam at high S_r values. Assumption of a proper UWC value would help to improve agreement when using Johansen, as it does to a certain

extent for Fairbanks silty clay loam (Table B30).

The effect, however, may not be sufficient, suggesting that Johansen should not generally be applied to frozen loams, i.e. silt-clays.

De Vries and Mickley also give unacceptably high values for these loams, but Mickley tends to give good agreement at high S_r values for Northway silt loam.

Summary. Kersten gives the best agreement up to an S_r value of 0.9; above this it provides many unacceptably high predictions. It also gives high (but acceptable) deviations for a low-quartz loam (Northway silt loam) but not for a low-quartz clay (Healy clay).

Johansen generally gives good predictions from an S_r value of 0.1 up to full saturation, provided a suitable UWC value is used. However this method gives some unacceptably high predictions for the frozen loams. De Vries gives similar results to Johansen for S_r values above 0.4, while Mickley generally gives good predictions at higher S_r values, preferably greater than 0.8.

Comparison of predictions with tabulated Soviet values. The predictions of the four methods were compared with the data for frozen clay soils tabulated in the U.S.S.R. Building Code (1960). The k_s value was taken as 2.0 W/m K and the UWC as zero, the actual values being unknown. The comparative results are shown in Table B33 which gives the differences between the calculated values and the Soviet tabulated data. Below an S_r value of 0.36 (moisture content of 18%), Kersten gives deviations that are too low (extending to -40%), but it agrees well at higher S_r values. Johansen gives somewhat better agreement than Kersten, while De Vries provides the best agreement throughout the saturation range from an S_r value of 0.14 to full saturation. Mickley gives good agreement at high S_r values and also, surprisingly, at S_r values below 0.2.

Applicability to saturated frozen fine soils. The usable data that are available for saturated frozen fine soils are rather limited. However they are sufficient to show some trends.

Table B34 shows the results of calculations on Penner's data for Leda clay at temperatures from -2.5 to -22°C. The values for k_s and UWC were taken from Penner's data (Penner 1970). Kersten's equation gives deviations that are too high, apart from the fact that it cannot allow for the effect of UWC. Johansen and De Vries give very good agreement throughout the temperature range. The other methods, modified resistor, Kunitz-Smith and Mickley give adequate or borderline agreement (up to a deviation of 37%).

In the case of Penner's Sudbury silty clay, appropriate values of k_s and UWC were again obtained from

Penner's (1970) data. The calculated values for the various methods are compared with the measured values at temperatures varying from 0 to -20°C in Table B35. Keistens gives a high deviation (43%) at 0°C, but because Keistens prediction does not change with temperature it agrees well with the larger measured values at -2.5°C and lower. Johansen and De Vries give excellent agreement throughout the temperature range whereas modified resistor, Kunit-Smith and Mickley are satisfactory at -2.5°C and below. These three latter methods show little difference in their predictions from the values given by a geometric mean equation assuming two phases only, i.e. solids and ice.

All the methods give excellent or good agreement for soil 7 (at -5°C) of Penner et al. (1975) and for the undisturbed permafrost (at -6°C) of Slisarchuk and Watson (1975), as is evident from Tables B36 and B37 in which UWC is assumed zero.

Also there is generally good agreement with the U.S.S.R. Building Code (1960) data for saturated frozen fine soils, especially at low dry densities and high moisture (ice) content (again UWC is assumed zero), as shown in Table B38.

It may be concluded that Johansen and De Vries give the best agreement and are capable of taking into account the effect of the unfrozen water content. Modified resistor, Kunit-Smith and Mickley are not as good and may not be quite adequate. Keistens gives some predictions that are too high and it cannot allow for UWC or its variation with temperature.

Applicability of methods to dry soils

The Keistens equations are not applicable to dry soils; nor is the Geman method. The Mickley equation gives values that are much too high and so does the McGaw equation when the interfacial efficiency ϵ is taken as unity. The other methods that were evaluated on data for dry soils were Johansen, De Vries (and adjusted De Vries), modified resistor, Kunit-Smith, Smith and Van Rooyen.

Applicability to dry coarse soils. The various methods were evaluated by comparing their predictions with the measured values obtained by Johansen (pers. comm.) in his experiments on certain natural gravel and sands as well as on some crushed rocks. These particular data were used because the quartz composition of these materials was known so k_s could be calculated and used in the various methods. Johansen is the only method which specifies different equations for natural soils and for crushed materials. In the case of Smith, an appropriate value for α' was used.

Johansen gives the best agreement for the gravels, mostly within +25%, followed closely by De Vries

(e.g. Table B39). Adjusted De Vries, Smith, Van Rooyen and modified resistor generally give good agreement but there are some inconsistencies, deviations in excess of 25% occur. Kunit-Smith gives many deviations that are too high and should therefore not be applied to gravels.

Analysis of the data for the sands shows that Johansen is the best method, providing good to excellent agreement, with all the deviations being negative (but not below -25%). The adjusted De Vries method also gives good to excellent predictions. Modified resistor and De Vries follow closely behind. The remaining three methods, Kunit-Smith, Smith and Van Rooyen, while showing some good agreement, also give deviations that are excessive (e.g. Table B40 for one of the sands).

For the natural coarse soils, the gravels and sands, Johansen gives the best results (within +25%). De Vries, adjusted De Vries and modified resistor generally give good or adequate predictions within +25%, with some values slightly beyond. While Smith and Van Rooyen give some good predictions, they also give some excessive deviations. Kunit-Smith gives unacceptable high deviations for the gravels and should therefore be considered inapplicable.

For the crushed rocks, the modified resistor is the best method, followed by adjusted De Vries. Although Johansen gives some good predictions it also provides a number of excessive deviations, positive and negative. Also Johansen, because it accounts for the porosity only, is insensitive to temperature changes which affect the value of k_s . The other four methods, Kunit-Smith, Smith, Van Rooyen and De Vries, show unacceptable deviations and should therefore be rejected (e.g. Table B41).

Applicability to dry fine soils. Comparisons were made between calculated values and values measured on fine soils by Smith and Byers (1938). Adjusted De Vries and modified resistor give the best agreement, generally well within +15%. The agreement with Johansen is not as good. It provides negative deviations extending to -25%.

A similar picture holds for Johansen's (pers. comm.) dry silts, i.e. adjusted De Vries and modified resistor are best while Johansen gives negative deviations as low as -33%.

The results for Johansen's (pers. comm.) dry clay are interesting. In this case Smith gives excellent agreement while all the other methods give values that are too low. This may be explained by the formation of secondary aggregations in this clay which would correspond to Smith's derivation and give rise to more effective heat transfer mechanisms. The result would be a greater effective thermal conductivity for the soil.

DISCUSSION AND CONCLUSIONS

The *Analysis of Methods for Calculating Thermal Conductivity* section described the basis of each of the methods for calculating the thermal conductivity of soils. The effect on this calculation of variations in the soil moisture content and the dry density was determined. The sensitivity to changes in the solids thermal conductivity k_s was also found. In order to determine the predictions of the methods it was necessary to know the soil mineral composition, particularly the quartz content from which k_s could be calculated. These predictions were compared with experimental data in the *Evaluation of Methods for Calculating Thermal Conductivity* section to determine the applicability of the methods to various types and conditions of soils.

For most practical applications it is sufficient to know the thermal conductivity to within about $\pm 25\%$ of its true value. Variation in soil properties from point to point in the field because of a lack of homogeneity could mean variations in the thermal conductivity to a similar extent. It is pointless to attempt to calculate thermal conductivity values to a higher degree of accuracy. Reasonable predictions are therefore considered to be those that do not deviate more than about $\pm 25\%$ from measured values.

Major errors are caused by not taking the soil mineralogical composition into account. According to Johansen (1975) this error could introduce uncertainties of about $\pm 30\%$ into the thermal conductivity value. It is obviously important to use a suitable k_s value in the methods that are based on the mineral composition.

Applicability to unfrozen soils

For unfrozen soils Kersten, Gemant, De Vries and Johansen show roughly similar trends with respect to variation of thermal conductivity with moisture content w at constant dry density γ_d or with γ_d at constant w . Mickley also gives similar trends for fine soil, but for coarse soil the curve of thermal conductivity against w (at constant γ_d) has an opposite curvature, indicating an increasing rate of change of thermal conductivity with increasing w which is contrary to what may be expected. Similar to Mickley, McGaw gives values that are too high in the dry or nearly dry condition, requiring an interfacial efficiency factor of less than unity to be applied. This method also gives a very low sensitivity of the thermal conductivity to variations in w . Van Rooyen shows rather odd thermal conductivity behavior, particularly at high values of w and γ_d where it becomes obviously inapplicable.

A comparison of Tables 1 and 2 reveals that all the methods, except Kersten, show an absolute sensi-

tivity s_w of the thermal conductivity to w (at constant γ_d) that is smaller for the fine soil than for the coarse soil. With respect to the absolute sensitivity s_g of the thermal conductivity to γ_d (at constant w), Kersten gives the lowest value for unfrozen coarse soils but the highest value for unfrozen fine soils.

Figures 21 and 25, which apply to coarse soils, show that Kersten gives the lowest curve for the thermal conductivity as compared with the other methods. Kersten's equation for coarse soil implies a k_s of about 5 W/m K (Farouki 1981). Kersten should therefore not be applied to unfrozen coarse soils having a high quartz content, since it seriously underpredicts for these soils. On the other hand it overpredicts for soils having a low quartz content. It should therefore be applied only to those unfrozen coarse soils with intermediate quartz content, say about 60% of the soil solids.

For degrees of saturation S_f above 0.2, Johansen provides the best agreement with the data for unfrozen coarse soils, giving deviations which are generally in the range $\pm 25\%$, while De Vries and Gemant generally deviate a little more. As can be seen from Figure 25, which applies to $S_f = 0.5$, De Vries gives a curve parallel and very close to Johansen's, while Gemant's curve is somewhat higher. In the range of S_f values from 0.1 to 0.2, De Vries in fact gives the best agreement, with deviations between 10 and -30% , while Johansen is next, covering a wider range between 20 and -40% . Below an S_f value of about 0.1, none of the methods gives good predictions, except Van Rooyen. This method gives reasonable values for sands and gravels down to S_f values about 0.015. However it underpredicts excessively for the crushed rocks of low quartz content which may be because Van Rooyen's empirical equation is not based on data for such materials.

Figures 22 and 26 show that, for unfrozen fine soils, Kersten gives the highest curve for the thermal conductivity (McGaw gives slightly higher values for unfrozen fine soil [Fig. 26]). The Kersten equation for fine soil implies a k_s of around 3.0 W/m K (Farouki 1981). Johansen, De Vries, Gemant and Mickley all give curves that are quite close together at $S_f = 0.5$ and $S_f = 1.0$ over the whole dry density range. Johansen generally gives the best predictions over the whole range of S_f values. Above $S_f = 0.2$ it gives deviations lying within the range $\pm 35\%$, but below $S_f = 0.2$ they may be as low as -45% . Kersten may be applied above $S_f = 0.3$ where it gives deviations within the range $\pm 35\%$ as does Johansen. However Kersten should not be applied below $S_f = 0.3$ because it then gives excessive deviations.

Applicability to frozen soils

In the case of frozen soils, the four applicable

methods (Kersten, Mickley, De Vries and Johansen) show generally similar trends with respect to variation of the thermal conductivity with γ_d at constant w (Fig. 11 and 12). With regard to variation of the thermal conductivity with w at constant γ_d , both Kersten and Johansen give a linear relationship. De Vries shows a lower rate of increase in thermal conductivity with increasing w , while Mickley gives a faster rate. The latter behavior is contrary to expectation (see Fig. 3 and 4).

Above a moisture content of 5%, s_w is greater for frozen soil than for unfrozen soil (compare Tables 3 and 4 with Tables 1 and 2 respectively). This may be expected because the thermal conductivity of ice is considerably higher than that of water. Comparing Table 4 to Table 3 one can see that s_w is greater for frozen coarse soil than for frozen fine soil. At a dry density of 1.4 g/cm³, Johansen gives a value of s_w for the coarse soil which is slightly more than twice the value for fine soil, while Kersten gives a value only about 40% larger. Both Johansen and Kersten give values of s_w that remain constant with changes in w at a given γ_d . The value of s_w increases with γ_d , particularly so for the coarse soil.

As with unfrozen coarse soils, Kersten overpredicts for frozen coarse soils having low quartz content, while it underpredicts when they have a high quartz content. Figures 23 and 27 show the low values given by Kersten as compared with the other methods in which a k_s value of 8.0 W/m K was used. Also Table 3 shows that Kersten gives a considerably lower s_w value than Johansen.

For unsaturated frozen coarse soils, Johansen gives the best predictions. These are reasonable for S_r values above 0.1 (approximately), while below this value Johansen gives some excessive deviations (though remaining the best predictor). While Mickley and De Vries may be applied with good results at high S_r values (above 0.6 for Mickley and above 0.8 for De Vries), computations can be more easily carried out with Johansen so that its general use is suggested for these soils.

With regard to frozen fine soils at $S_r = 0.5$, Kersten provides a curve of thermal conductivity against γ_d which differs little from the curves provided by Johansen and Mickley, while De Vries' curve is appreciably higher (Fig. 28). Also Table 4 shows that Kersten and Johansen give s_w values that are nearly the same.

While the predictions of the methods were compared with only a limited amount of available data for unsaturated frozen fine soils, certain trends can be seen. Kersten provides good agreement (generally within + 30%) up to an S_r of 0.9. Beyond this it gives deviations that are too high for naturally occurring frozen soils such as Slusarchuk and Watson's

undisturbed permafrost and Penner's Leda clay (Fig. 40a). On the other hand, while Johansen gives a few high deviations at values of S_r below 0.1, it otherwise generally gives good predictions (within $\pm 35\%$) up to and including $S_r = 1$ (saturation). Thus, while Kersten may be applied for values of S_r below 0.9, Johansen should be used for higher values of S_r . De Vries gives values close to Johansen at S_r values above 0.4. Both these methods can allow in their equations for the presence of unfrozen water, while Kersten cannot do so.

Applicability to saturated soils

In a saturated soil the ratio of the thermal conductivities of the phases is low. It varies from nearly 15:1 for quartz-water to about 1:1 for clay-ice. Such a low ratio means that application of a geometric mean equation, as in the Johansen method, should give good agreement with measured values (Farouki 1981). In fact, for the unfrozen soils all the applicable methods (Johansen, De Vries, Gemant, Mickley, McGaw, Kunii-Smith and modified resistor, but not Kersten and Van Rooyen) gave good agreement. The resulting deviations were within the range $\pm 25\%$. The easiest method to use is Johansen as it reduces to a simple geometric mean equation.

The situation is similar for saturated frozen coarse soils where any of the applicable methods (Johansen, De Vries, Mickley, modified resistor and Kunii-Smith) may be used except Kersten. The Kersten method overpredicts for low-quartz coarse materials, frozen or unfrozen, in the saturated condition just as in the unsaturated condition.

For the saturated frozen fine soils, Johansen and De Vries give the best agreement. Moreover they are capable of taking the unfrozen water content into account which Kersten and the other methods cannot. It is important to know the unfrozen water content present in a given fine-grained soil at temperatures below 0°C. This depends in particular on the specific surface area of the soil.

It should be noted that Johansen assumes that the thermal conductivity of the unfrozen water is the same as that of ordinary water. While this may be true for a large part of the unfrozen water, it has been proposed that the strongly adsorbed unfrozen water (the boundary phase) may have a relatively high thermal conductivity, perhaps even higher than that of ice (Farouki 1981).

Effect of soil mineral composition

In calculating soil thermal conductivity it is important to know the soil mineral composition, particularly its quartz content, as has been stressed. It has been shown that the Kersten method should not be applied to coarse soils having high or low quartz

content. For the other methods, the quartz content is required so that the thermal conductivity of the soil solids k_s can be calculated.

There is some uncertainty regarding the "true" value of the thermal conductivity of quartz. In conformity with Lachenbruch (among others) the value used in the *Analysis of Methods for Calculating Thermal Conductivity* section was obtained from a geometric mean equation applied to the thermal conductivity values of the quartz crystals measured along and at right angles to the principal axis (k_1 and k_2 , respectively). Recently Linnik (pers. comm.) reported a quartz thermal conductivity of 10 W/m K which was inferred from measurements on saturated materials. However, it is difficult to see how such a high value can result from the tabulated values of k_1 and k_2 (Table 14). Further research on this matter is necessary.

Similarly more data are required on the thermal conductivity of minerals, other than quartz, that may be present in soils. In particular measurements of the values for the various clay minerals would be useful. There are indications that the thermal properties of kaolinite, illite, montmorillonite, etc. may be different.

If a coarse soil has a high quartz content, the thermal conductivity of the other minerals present k_o makes little difference to the soil's thermal conductivity. On the other hand, when the quartz content is low, k_o and its variation has a considerable influence on k_s and therefore on the soil thermal conductivity as may be seen from Figures 33, 34, 38 and 39. With Johansen the effect of a change of k_o from 2.0 to 3.0 W/m K is to increase the deviations by 20 or 30% at intermediate S_f values for the unfrozen condition (Fig. 33a) and similarly for the frozen condition (Fig. 38a).

For fine soils a k_s value of 2.0 W/m K has generally been used in the calculations in earlier sections. It would be more accurate to use a value dependent on the type of clay minerals present but further information on these would be required.

Figures 17 and 18 can be used to determine the effect of changes in k_s on the soil thermal conductivity according to the various applicable methods (except Kerssen) and for unfrozen or frozen soils. The sensitivity of the thermal conductivity to k_s resulting from Johansen, De Vries and Gemant increases markedly as S_f increases, more so for the frozen condition. This implies that knowledge of a more accurate value of k_s is more important for soils having higher values of S_f .

Assumption of a suitable k_s value in Gemant gives better agreement between its predictions and measured

values. The nomogram of Makowski and Mochalski (1956), which is based on Gemant's equations, should therefore be redeveloped on the basis of a more suitable value for k_s . This could be calculated from the quartz content and its thermal conductivity rather than from Gemant's subsidiary equation for k_s .

The effect of temperature on the value of k_s should be taken into account. The thermal conductivity of quartz increases as the temperature decreases as shown in Table 14. This variation was allowed for in the calculations of the *Evaluation of Methods for Calculating Thermal Conductivity* section. Similarly the thermal conductivity of the other soil minerals, except feldspar, decreases with increasing temperature. Feldspar may be present in coarse or fine soils and, if minerals other than quartz are also present, the effect of the feldspar would be to provide some counterbalance to the variation of k_o with temperature (i.e. k_o would have less sensitivity to temperature variation). Because of the uncertainty in the actual magnitude of k_o , this value was not varied with temperature in the calculations using the various methods. Such an allowance may be made if the amounts and properties of the minerals other than quartz are known.

Applicability to dry soils

Dry soils have a high ratio of thermal conductivity of the two components (solids and air). As a result the region encompassed by the Hashin-Shtrikman bounds is wide and a geometric mean equation does not give good results. The thermal conductivity is highly sensitive to variations in microstructure (see Johansen 1975). This is taken into account by Johansen's method which allows for two different empirical equations, one for natural and the other for crushed materials. The former is a function only of the soil's dry density while the latter is a function of its porosity alone. The implication is that the solids thermal conductivity has little effect.

The analysis of the predictions of the various methods showed that Johansen applies well to dry natural coarse soils (within ± 25%) but not to dry crushed rocks. The thermal conductivity of these was better predicted by the modified resistor or adjusted De Vries methods (the Kersten method is inapplicable). These two methods also applied well for most of the dry fine-grained soils. However, Smith gave the best results for a dry clay. This method allows for structural effects by means of a thermal structure factor. Considerations of structure and contact effects are particularly important in the case of dry soils, fine or coarse.

Summary of applicability of methods

The best methods to apply for different types and conditions of soils are as follows. In most cases these methods are predictions within about $\pm 25\%$, which is acceptable for practical application.

Unfrozen coarse soils

$0.015 < S_f < 0.1$	Van Rooyen for sands and gravels (not for low quartz crushed rocks)
$0.1 < S_f < 0.2$	De Vries
$S_f > 0.2$	Johansen
Saturated	Gemant

Note: Kersten's method should not be applied to coarse soils with low or high quartz content.

Unfrozen fine soils

$0 < S_f < 0.1$	Johansen (increase prediction by 15%)
$0.1 < S_f < 0.2$	Johansen (increase prediction by 5%)
$S_f > 0.2$	Johansen
Saturated	Johansen, De Vries, modified resistor, Kunii-Smith, Mickley, Gemant, or McGaw

Frozen coarse soils

$S_f > 0.1$	Johansen
Saturated	Johansen, De Vries, Mickley, modified resistor or Kunii-Smith

Note: Kersten should not be applied to frozen coarse soils with low or high quartz content.

Frozen fine soils

$S_f < 0.9$	Kersten
$0.1 < S_f < 1$	Johansen (with suitable unfrozen water content)
Saturated	Johansen and De Vries (Kersten should not be used where unfrozen water content is appreciable)

Dry coarse soils

Natural	Johansen
Crushed rocks	Modified resistor, adjusted De Vries

Dry fine soils

General	Modified resistor, adjusted De Vries
Clay	Smith

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APPENDIX A: PROPERTIES OF SOME TEST SOILS

Soils of Kersten (1949)

Table A1. General physical properties of soils.

Soil no.	Soil designation	Mechanical analysis					Physical constants					Textural class		
		Sand	Silt	Modified optimum moisture content			max. density	Specific gravity	Absorption (%)	U.S. Bur. of Chem. and soils	Unified soil classn.			
		Gravel 0.05 over 2.00 ¹	0.005 to 2.00	Clay 0.05 to 0.005	Liquid ² limit	Plasticity ³ index	N.P. ³	—	—	2.70	0.75	Gravel GP		
P4601	Chena River gravel	80.0	19.4	— 0.6 —	—	N.P. ³	—	—	2.70	0.75	Gravel GP			
P4703	Crushed quartz	15.5	79.0	— 5.5 —	—	N.P.	—	—	2.65	0.26	Coarse sand SW			
P4704	Crushed trap rock	27.0	63.0	— 10.0 —	—	N.P.	—	—	2.97	0.20	Coarse sand SM			
P4705	Crushed feldspar	25.5	70.3	— 4.2 —	—	N.P.	—	—	2.56	0.75	Coarse sand SW			
P4706	Crushed granite	16.2	77.0	— 6.8 —	—	N.P.	—	—	2.67	0.56	Coarse sand SW			
P4702	20-30 Ottawa sand	0.0	100.0	0.0 0.0	—	N.P.	—	—	2.65	0.17	Coarse sand SP			
P4701	Graded Ottawa sand	0.0	99.9	— 0.1 —	—	N.P.	—	—	2.65	0.19	Medium sand SP			
P4714	Fine crushed quartz	0.0	100.0	0.0 0.0	—	N.P.	—	—	2.65	—	Medium sand SP			
P4709	Fairbanks sand	27.5	70.0	— 2.5 —	—	N.P.	12.0	122.5	2.72	—	Medium sand SW			
P4604	Lowell sand	0.0	100.0	0.0 0.0	—	N.P.	12.2	119.0	2.67	—	Medium sand SW			
P4503	Northway sand	3.0	97.0	0.0 0.0	—	N.P.	14.0	112.8	2.74	—	Medium sand SW			
P4502	Northway fine sand	0.0	97.0	3.0 0.0	—	N.P.	11.4	116.0	2.76	—	Fine sand SP			
P4711	Dakota sandy loam	10.9	57.9	21.2 10.0	17.1	4.9	6.5	138.5	2.71	—	Sandy loam SM			
P4713	Ramsey sandy loam	0.4	53.6	27.5 18.5	24.6	9.3	9.0	127.5	2.68	—	Sandy loam CL			
P4505	Northway silt loam	1.0	21.0	64.4 13.6	27.3	N.P.	15.7	112.0	2.70	—	Silt loam ML			
P4602	Fairbanks silt loam	0.0	7.6	80.9 11.5	34.0	N.P.	15.5	110.0	2.70	—	Silt loam ML			
P4710	Fairbanks silty clay loam	0.0	9.2	63.8 27.0	39.2	12.4	18.0	102.0	2.71	—	Silty clay loam			
P4708	Healy clay	0.0	1.9	20.1 78.0	39.4	15.0	17.0	108.0	2.59	—	Clay CL			
P4707	Fairbanks peat	—	—	—	—	N.P.	—	—	—	—	Peat Pt			

¹ Size in millimeters.

² Minus no. 40 mesh fraction.

³ N.P. = non-plastic.

Table A2. Mineral and rock composition of soils (percentage by weight).

Soil no.	Soil designation	Unified soil classn.	Quartz			Ortho-feldspar	Plagioclase feldspar	Pyroxene, amphibole, and olivine	Basic igneous rock	Kaolinite clay min. and clay coat, min.	Hematite and magnetite	Mica	Coal	Others
			By petrogr. exam.	By X-ray analysis	Felsite									
P4601	Chena River gravel	GP	43.1		11.6		12.9	27.0			2.1		3.3	
P4703	Crushed quartz	SW	95+ ¹								2.0		1.0	
P4704	Crushed trap rock	SM	3.0		10.0		50.0 ²	34.0						
P4705	Crushed feldspar	SW	15.0		55.0		30.0							10.0
P4706	Crushed granite	SW	20.0		30.0		40.0							
P4702	20-30 Ottawa sand	SP	99+ ³											
P4701	Graded Ottawa sand	SP	99+ ³											
P4714	Fine crushed quartz	SP	95+ ¹											
P4709	Fairbanks sand	SW	59.4		3.6	5.0	6.3	8.0	10.0		2.5	0.1	5.1	
P4604	Lowell sand	SW	72.2		20.5			3.0			1.3		3.0	
P4503	Northway sand	SW	7.5			11.5	9.0	7.5	51.0				13.5	
P4502	Northway fine sand	SP	12.0			7.0	18.0	12.0	40.0				11.0	
P4711	Dakota sandy loam	SM	59.1		12.9		1.0	12.1		12.4			2.5	
P4713	Ramsey sandy loam	CL	51.3		11.8		5.6	12.6		15.9			2.8	
P4505	Northway silt loam	ML	1.5				31.5	19.5	4.5	27.5	10.0		5.5	
P4602	Fairbanks silt loam	ML	13.3	40.3						28.3		18.1		
P4710	Fairbanks silty clay loam	ML	4.6	59.5				2.2		28.9	1.6	3.2		
P4708	Healy clay	CL	22.5							55.0		22.0	0.5	

¹ By visual inspection; impurities less than 5%.

² Andesine feldspar.

³ By visual inspection; impurities less than 1%.

Soils of Penner et al. (1975)

The grain size distribution curves for soils no. 4 to 10 tested by Penner et al. (1975) are given in Figure A1. The sieve and hydrometer analyses followed ASTM procedures.

Table A3. Atterberg limits.

Soil no.	Liquid limit at 25 blows W_L (%)	Plastic limit, W_p (%)	Plasticity index I_p (%)
4	37	21	16
5	25	18	7
6	30	21	9
7	28	14	14
8	43	24	18
9	33	22	11
10	48	23	25

The Atterberg limits for soils no. 4 to 10 are given in Table A3. Table A4 shows the relative proportions of minerals in the size fraction smaller than μm .

Table A4. Relative proportions of minerals in the <0.002 mm size fraction.

Soil	Quartz	Illite	Chlorite	Kaolinite	Vermiculite
4	+++	++	++	++	-
5	+	+	-	-	-
6	+	+	+	+	-
7	++	++	+	++	-
8	+++	+++	++	+	-
9	+++	+++	+	+	-
10	+++	+++	++	++	-

+ = small amount present; ++ = moderate amount present;
+++ = large amount present.

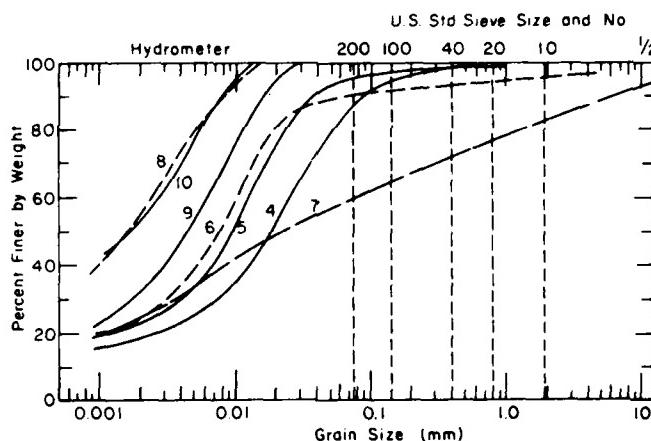


Figure A1. Grain size analysis of soils tested by Penner et al. (1975).

Soils of Johansen

Table A5. Soil materials of Johansen (pers. comm.).

<i>Material</i>	<i>Quartz content (%)</i>	<i>Specific weight (kg/m³)</i>	<i>d₆₀ (mm)</i>	<i>d₁₀ (mm)</i>	<i>Uniformity coefficient d₆₀/d₁₀</i>
Sand					
SA1	50	2700	0.69	0.12	5.7
SA2	48	2600	0.27	0.067	4.0
SA3	58	2670	0.15	0.082	1.8
SA4	80	2670	0.16	0.125	1.3
SA5	45	2720	0.26	0.11	2.4
SA7	39	2700	1.5	0.10	15.0
SA8	61	2730	0.20	0.075	2.7
SA10	10	2850	0.52	0.125	4.2
SA13	100	2650	0.70	0.60	1.2
Gravel					
GR1	49	2740	2.0	0.21	9.5
GR3	47	2700	2.0	0.20	10.0
GR6	57	2700	6.1	0.27	22.6
GR7	2	3000	1.4	0.06	23.3
GR12	41	2700	1.8	0.15	12.0
Crushed rock					
PU1	2	3000	31.0	24.0	1.3
PU5	33	2680	6.1	0.25	24.4
PU7	9	2750	35.0	26.0	1.3
PU9	3	2730	8.5	1.0	8.5
PU10	9	3100	17.0	10.5	1.6
Clay					
LE1	22	2800	0.040	0.0035	—

**APPENDIX B: COMPARISON OF THERMAL CONDUCTIVITY
VALUES COMPUTED BY THE VARIOUS METHODS AND OF
THEIR DEVIATIONS FROM THE VALUES MEASURED**

Table B1.

KEKSTEN MIRTH JAY SAND UNSATURATED UNFREEZEN

TYPE OF SOIL	SAND	NATURAL	UNSATURATED	UNFREEZEN	RHO = 2.660	TEMP = 6.307 K	0.273 CL %	0.000 CSOLIDS = *	6.220 ALPH = -----	KA = 0.056	
KW = 0.569	KICK = 2.200	KI = 8.020	KO = 2.000	CSOLID2 = *							
SAMPLE	W, %	DRY DENSITY	K MEASURED	METHOD K COMPUTER	DEVIATION (PERCENT)	METHOD K COMPUTER	DEVIATION (PERCENT)	PUR SAT	NC	WC	
0.00	1.477	J,202	KERSTEN	J,204	-4.1	MICKLEY	0.688	240.6	0.481	0.009	
			JOHANSEN	J,203	+12.9	MC GAW	1.1418	+52.6	0.019	0.009	
			ADJ DE VRIES	-----		GEMANT	-----				
			RESISTOR	-----		SMITH	-----				
			KUUTI-SMITH	-----		VAN ROOYEN	0.180	+10.9			
						GEOM, MEAN	-----				
SAMPLE	0.10	1.600	J,253	KERSTEN	6.3	MICKLEY	0.756	199.4	0.616	0.011	
			JOHANSEN	J,260	+6.1	MC GAW	1.125	199.4	0.627	0.011	
			DE VRIES	J,253	+10.9	SMITH	-----				
			ADJ RESISTOR	-----		VAN ROOYEN	0.223	+12.0			
			KUUTI-SMITH	-----		GEOM, MEAN	-----				
SAMPLE	0.50	1.618	J,254	KERSTEN	0.690	93.1	MICKLEY	0.777	206.0	0.410	0.013
			JOHANSEN	J,257	+13.0	MC GAW	1.142	+52.4	0.032	0.013	
			ADJ DE VRIES	-----		GEMANT	-----				
			RESISTOR	-----		SMITH	-----				
			KUUTI-SMITH	-----		VAN ROOYEN	0.236	+7.1			
SAMPLE	0.60	1.693	J,280	KERSTEN	0.566	69.4	MICKLEY	0.826	186.4	0.393	0.012
			JOHANSEN	J,271	+6.2	MC GAW	1.139	186.4	0.393	0.012	
			DE VRIES	J,580	+104.2	SMITH	-----				
			ADJ RESISTOR	-----		VAN ROOYEN	0.260	+6.7			
			KUUTI-SMITH	-----		GEOM, MEAN	-----				
SAMPLE	5.80	1.475	0,495	KERSTEN	0.964	45.9	MICKLEY	0.720	+5.5	0.462	0.056
			JOHANSEN	J,895	+13.8	MC GAW	1.134	+13.4	0.121	0.056	
			ADJ DE VRIES	-----		GEMANT	0.677	+3.5			
			RESISTOR	-----		SMITH	-----				
			KUUTI-SMITH	-----		VAN ROOYEN	0.308	+3.7			
SAMPLE	4.10	1.611	0,581	KERSTEN	1.613	78.1	MICKLEY	0.815	+1.7	0.412	0.066
			JOHANSEN	J,697	+2.5	MC GAW	1.154	+8.6	0.160	0.066	
			DE VRIES	J,735	+8.0	SMITH	-----				
			ADJ RESISTOR	-----		VAN ROOYEN	0.916	+30.7			
			KUUTI-SMITH	-----		GEOM, MEAN	-----				
SAMPLE	4.50	1.691	0,782	KERSTEN	1.584	77.0	MICKLEY	0.876	+14.9	0.393	0.059
			JOHANSEN	J,698	+1.9	MC GAW	1.122	+14.4	0.190	0.059	
			ADJ DE VRIES	-----		GEMANT	0.724	+7.4			
			RESISTOR	-----		SMITH	-----				
			KUUTI-SMITH	-----		VAN ROOYEN	0.353	+5.6			
SAMPLE	4.20	1.669	0,384	KERSTEN	1.707	92.2	MICKLEY	0.934	+5.2	0.391	0.060
			JOHANSEN	J,906	+13.5	MC GAW	1.245	+40.2	0.393	0.154	
			DE VRIES	J,100	+14.3	SMITH	-----				
			ADJ RESISTOR	-----		VAN ROOYEN	0.366	+16.0			
			KUUTI-SMITH	-----		GEOM, MEAN	-----				
SAMPLE	14.00	1.587	0,934	KERSTEN	1.618	92.6	MICKLEY	0.929	+1.8	0.931	0.092
			JOHANSEN	J,050	+12.8	MC GAW	1.128	-----			
			DE VRIES	-----		GEMANT	1.068	+14.6			
			ADJ RESISTOR	-----		SMITH	-----				
			KUUTI-SMITH	-----		VAN ROOYEN	0.359	+8.14			
SAMPLE	15.00	1.696	1,044	KERSTEN	1.711	88.1	MICKLEY	1.041	+0.7	0.381	0.059
			JOHANSEN	J,162	+10.8	MC GAW	1.307	+24.7	0.606	0.231	
			DE VRIES	J,162	+10.8	SMITH	-----				
			ADJ RESISTOR	-----		VAN ROOYEN	0.378	+13.0			
			KUUTI-SMITH	-----		GEOM, MEAN	-----				
SAMPLE	15.50	1.615	1,204	KERSTEN	2.286	94.9	MICKLEY	1.269	+7.6	0.933	0.284
			JOHANSEN	J,309	+8.7	MC GAW	1.270	+7.6			
			DE VRIES	-----		SMITH	-----				
			ADJ DE VRIES	-----		GEMANT	1.342	+11.5			
			RESISTOR	-----		SMITH	-----				
			KUUTI-SMITH	-----		VAN ROOYEN	0.374	+6.9			
SAMPLE	16.50	1.693	1,171	KERSTEN	2.064	76.3	MICKLEY	1.105	+3.7	0.381	0.059
			JOHANSEN	J,421	+4.3	MC GAW	1.331	+13.7	0.733	0.480	
			DE VRIES	J,424	+5.4	SMITH	-----				
			ADJ RESISTOR	-----		VAN ROOYEN	0.368	+6.7			
			KUUTI-SMITH	-----		GEOM, MEAN	-----				

NOTES:

Moisture content in %.

Dry density in g/cm³.

All other units W/m K except where noted.

DEVIATION refers to difference between value COMPUTED by respective METHOD and the MEASURED value.

TEMP = test temperature (°C).

ALPHA = parameter in Smith method.

O = Quartz content in soil solids (fractional).

CL = Clay content in soil solids (fractional).

RHO = Specific gravity of soil solids.

CSOLID2 = Thermal conductivity of solids (calculated or assumed).

CSOLID2 = Value from Gemant's equation.

KA, KW, KICL, KO = conductivities of air, water, ice, quartz and other minerals respectively (values found according to test TEMPERATURE).

POR = Porosity.

SAT = Degree of saturation.

SC = Parameter in McGaw's method.

IWC = Unfrozen water content (fraction by volume of total soil volume).

Table B2.

KERSTEN NURTHWAY FINE SAND UNSATURATED UNFRUZEN

TYPE OF SULFUR NATURAL UNSATURATED UNFROZEN
RHO = 2.760 TEMP = 44.000 E = 0.120 CL = 0.000 C30LUS = 2.363 ALPHA = ----- KA = 2.25
K = 2.000 TFE = 2.400 KU = 2.000 CSOLUS = 1.820

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PW, SAT	NC	
SAMPLE	0.50	1.564	0.208	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	0.257 0.260 0.260 0.260 -----	23.7 147.1	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	0.712 1.321 1.321 1.321 0.228 0.228	222.6 536.3 536.3 536.3 -3.6 -3.6	R.035 R.018
SAMPLE	0.50	1.645	0.234	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	0.298 0.298 0.298 0.298 -----	24.3 87.7 134.6	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	0.828 1.399 1.399 1.399 0.228 0.228	234.7 455.4 455.4 455.4 -2.0 -2.0	R.024 R.020
SAMPLE	0.50	1.707	0.297	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	0.348 0.348 0.348 0.348 -----	16.0 111.0	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	0.914 1.489 1.489 1.489 0.207 0.207	227.7 421.3 421.3 421.3 -14.0 -14.0	R.309 R.025
SAMPLE	2.00	1.651	0.392	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	0.466 0.466 0.466 0.466 -----	141.1 50.1	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	0.853 1.369 1.369 1.369 0.324 0.324	117.4 231.6 231.6 231.6 -33.6 -33.6	R.462 R.022
SAMPLE	1.00	1.702	0.438	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	1.088 1.088 1.088 1.088 -----	145.5 79.4	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	0.932 1.396 1.396 1.396 0.324 0.324	112.7 217.8 217.8 217.8 -26.0 -26.0	R.362 R.023
SAMPLE	2.10	1.759	0.497	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	1.138 1.138 1.138 1.138 -----	127.0 52.7	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	0.973 1.476 1.476 1.476 0.320 0.320	147.6 176.6 176.6 176.6 -13.8 -13.8	R.363 R.022
SAMPLE	5.10	1.563	0.796	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	1.222 1.222 1.222 1.222 -----	53.5 44.4	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	0.820 1.164 1.164 1.164 0.343 0.343	4.2 46.3 46.3 46.3 -3.4 -3.4	R.434 R.024
SAMPLE	5.20	1.645	0.867	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	1.385 1.385 1.385 1.385 -----	59.6 50.7	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	0.892 1.235 1.235 1.235 0.354 0.354	42.6 59.3 59.3 59.3 -39.2 -39.2	R.464 R.212
SAMPLE	5.20	1.759	0.999	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	1.528 1.528 1.528 1.528 -----	63.8 55.9	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	0.985 1.336 1.336 1.336 0.366 0.366	71.5 133.7 133.7 133.7 -7.8 -7.8	R.363 R.252
SAMPLE	10.00	1.645	1.087	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	1.728 1.664 1.664 1.664 -----	58.0 52.0	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	0.989 1.284 1.284 1.284 0.363 0.363	79.9 181.1 181.1 181.1 -66.6 -66.6	R.464 R.024
SAMPLE	11.20	1.707	1.249	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	2.074 1.212 1.212 1.212 -----	66.1 55.9 55.9	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	1.168 1.598 1.598 1.598 0.371 0.371	-11.3 11.9 11.9 11.9 -76.3 -76.3	R.368 R.058
SAMPLE	11.40	1.660	1.368	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	2.381 1.359 1.359 1.359 -----	74.6 67.7 67.7	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	1.221 1.488 1.488 1.488 0.377 0.377	-16.8 8.7 8.7 8.7 -72.5 -72.5	R.326 R.050
SAMPLE	13.00	1.765	1.380	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KUNITZ-SMITH	2.189 1.284 1.284 1.284 -----	58.6 55.9 55.9	MICKLEY MEYER SMITH GERMANI VAN RODEN GEOM. MEAN	1.169 1.421 1.421 1.421 0.371 0.371	-16.3 3.0 3.0 3.0 -73.1 -73.1	R.368 R.081

Table B3.

JUHANSEN GRAVEL GR7 MUIST UNFROZEN

TYPE OF SOIL: COARSE NATURAL UNSATURATED UNFROZEN

RHO = 3,000 TEMP = 2,100 U = 0,020 CL = 0,000 CSOLIUS = T 2,057 ALPHA = ----- KA = 0,044
K = 0,504 ICE = T 2,000 KU = 0,100 KD = 2,000 CSOLIUS2 = 0,848

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PUR. SAT.	NC URL
SAMPLE	2,50	1,537	0,490	KERSTEN	0,091	-81,9	MICKLEY	0,018	24,5
				JOHANSEN	0,420	-14,3	MC GEE	0,076	99,2
				DE VRIES	0,464	-6,4	SMITH	-----	-----
				ADJ DE VRIES	-----	-----	GEMANT	0,205	25,9
				KUNIT-SMITH	-----	-----	VAN RUYEN	0,278	43,4
							GEOM. MEAN	-----	-----

RHO = 3,000 TEMP = 3,100 U = 0,020 CL = 0,000 CSOLIUS = T 2,057 ALPHA = ----- KA = 0,048
K = 0,508 ICE = T 2,000 KU = 0,040 CSOLIUS2 = 0,848

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PUR. SAT.	NC URL
SAMPLE	2,50	1,537	0,505	KERSTEN	0,091	-76,5	MICKLEY	0,015	31,8
				JOHANSEN	0,421	-16,7	MC GEE	0,070	91,9
				DE VRIES	0,470	-5,7	SMITH	-----	-----
				ADJ DE VRIES	-----	-----	GEMANT	0,210	36,9
				KUNIT-SMITH	-----	-----	VAN RUYEN	0,278	45,0
							GEOM. MEAN	-----	-----

RHO = 3,000 TEMP = 2,500 U = 0,020 CL = 0,000 CSOLIUS = T 2,053 ALPHA = ----- KA = 0,104
K = 0,505 ICE = T 2,000 KU = 0,430 CSOLIUS2 = 0,840

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PUR. SAT.	NC URL
SAMPLE	2,50	1,537	0,049	KERSTEN	0,091	-37,4	MICKLEY	0,002	5,6
				JOHANSEN	0,420	-34,0	MC GEE	0,016	56,5
				DE VRIES	0,018	-4,7	SMITH	-----	-----
				ADJ DE VRIES	-----	-----	GEMANT	0,531	88,2
				KUNIT-SMITH	-----	-----	VAN RUYEN	0,278	57,2
							GEOM. MEAN	-----	-----

RHO = 3,000 TEMP = 3,500 U = 0,020 CL = 0,000 CSOLIUS = T 2,050 ALPHA = ----- KA = 0,048
K = 0,508 ICE = T 2,000 KU = 0,040 CSOLIUS2 = 0,848

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PUR. SAT.	NC URL
SAMPLE	2,50	1,537	0,053	KERSTEN	1,397	100,1	MICKLEY	0,708	34,4
				JOHANSEN	0,423	-44,6	MC GEE	1,140	-----
				DE VRIES	0,023	-2,3	SMITH	-----	-----
				ADJ DE VRIES	-----	-----	GEMANT	0,515	73,1
				KUNIT-SMITH	-----	-----	VAN RUYEN	0,340	53,9
							GEOM. MEAN	-----	-----

RHO = 3,000 TEMP = 2,500 U = 0,020 CL = 0,000 CSOLIUS = T 2,053 ALPHA = ----- KA = 0,104
K = 0,505 ICE = T 2,000 KU = 0,430 CSOLIUS2 = 0,840

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PUR. SAT.	NC URL
SAMPLE	2,50	1,503	0,478	KERSTEN	1,397	100,1	MICKLEY	0,708	34,4
				JOHANSEN	0,423	-44,6	MC GEE	1,140	-----
				DE VRIES	0,780	-1,4	SMITH	-----	-----
				ADJ DE VRIES	-----	-----	GEMANT	0,770	71,0
				KUNIT-SMITH	-----	-----	VAN RUYEN	0,340	50,2
							GEOM. MEAN	-----	-----

RHO = 3,000 TEMP = 3,500 U = 0,020 CL = 0,000 CSOLIUS = T 2,050 ALPHA = ----- KA = 0,053
K = 0,508 ICE = T 2,000 KU = 0,040 CSOLIUS2 = 0,840

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PUR. SAT.	NC URL
SAMPLE	2,50	1,503	0,478	KERSTEN	1,397	100,1	MICKLEY	0,820	6,5
				JOHANSEN	0,423	-44,6	MC GEE	1,173	40,8
				DE VRIES	0,780	-1,4	SMITH	-----	-----
				ADJ DE VRIES	-----	-----	GEMANT	0,770	71,0
				KUNIT-SMITH	-----	-----	VAN RUYEN	0,340	50,2
							GEOM. MEAN	-----	-----

RHO = 3,000 TEMP = 2,500 U = 0,020 CL = 0,000 CSOLIUS = T 2,053 ALPHA = ----- KA = 0,126
K = 0,508 ICE = T 2,000 KU = 0,430 CSOLIUS2 = 0,840

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PUR. SAT.	NC URL
SAMPLE	2,50	1,488	0,465	KERSTEN	1,117	71,3	MICKLEY	0,624	35,2
				JOHANSEN	0,591	-38,3	MC GEE	0,919	42,5
				DE VRIES	0,591	-1,7	SMITH	-----	-----
				ADJ DE VRIES	-----	-----	GEMANT	0,567	12,1
				KUNIT-SMITH	-----	-----	VAN RUYEN	0,330	44,0
							GEOM. MEAN	-----	-----

RHO = 3,000 TEMP = 3,500 U = 0,020 CL = 0,000 CSOLIUS = T 2,053 ALPHA = ----- KA = 0,126
K = 0,508 ICE = T 2,000 KU = 0,040 CSOLIUS2 = 0,840

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PUR. SAT.	NC URL
SAMPLE	2,50	1,488	0,465	KERSTEN	1,117	71,3	MICKLEY	0,710	11,9
				JOHANSEN	0,591	-24,8	MC GEE	1,107	19,9
				DE VRIES	0,799	-12,4	SMITH	-----	-----
				ADJ DE VRIES	-----	-----	GEMANT	0,712	16,0
				KUNIT-SMITH	-----	-----	VAN RUYEN	0,330	59,0
							GEOM. MEAN	-----	-----

RHO = 3,000 TEMP = 2,500 U = 0,020 CL = 0,000 CSOLIUS = T 2,053 ALPHA = ----- KA = 0,050
K = 0,508 ICE = T 2,000 KU = 0,430 CSOLIUS2 = 0,840

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PUR. SAT.	NC URL
SAMPLE	2,50	1,488	0,465	KERSTEN	1,117	71,3	MICKLEY	0,704	12,9
				JOHANSEN	0,596	-15,4	MC GEE	1,107	21,0
				DE VRIES	0,777	-13,7	SMITH	-----	-----
				ADJ DE VRIES	-----	-----	GEMANT	0,431	27,7
				KUNIT-SMITH	-----	-----	VAN RUYEN	0,365	59,5
							GEOM. MEAN	-----	-----

RHO = 3,000 TEMP = 3,500 U = 0,020 CL = 0,000 CSOLIUS = T 2,053 ALPHA = ----- KA = 0,138
K = 0,508 ICE = T 2,000 KU = 0,040 CSOLIUS2 = 0,840

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PUR. SAT.	NC URL
SAMPLE	2,50	1,489	0,466	KERSTEN	1,074	89,8	MICKLEY	0,704	12,9
				JOHANSEN	0,596	-15,4	MC GEE	1,107	21,0
				DE VRIES	0,777	-13,7	SMITH	-----	-----
				ADJ DE VRIES	-----	-----	GEMANT	0,431	27,7
				KUNIT-SMITH	-----	-----	VAN RUYEN	0,365	59,5
							GEOM. MEAN	-----	-----

NOTE:
In plotting the figures the values corresponding to the temperature nearest 40°C were used.

Table B3 (cont'd).

SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	DUM, SAT, NC	
SAMPLE	5.4%	1.607	1.607	KERSTEN	1.674	-5.9	MILKLEY	1.687	-16.9
				JOHANSEN	1.621	-2.5	MC GEE	1.624	-0.2
				DE VRIES	1.691	+15.0	SMITH	1.624	-
				ADJ DE VRIES	1.691	+15.0	GEMANI	1.673	+10.8
				RESISTANCE	1.691	+15.0	VAN RUYEN	1.636	+70.0
				KUNIJS-SMITH	1.691	+15.0	GLOM, MEAN	1.636	-
RHO =	0.6800 TEMP =	20.000 u =	0.6800 CL =	0.6800 CSOLIUS =	+0.000	2.600 ALPHA =	-0.101		
KA =	0.608 X ICE =	2.608 KU =	0.608 KU	2.608 CSOLIUS2 =	+0.000				
SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	DUM, SAT, NC	
SAMPLE	8.3%	1.469	0.920	KERSTEN	1.571	-5.9	MILKLEY	1.605	-21.2
				JOHANSEN	1.594	-2.3	MC GEE	1.591	-0.2
				DE VRIES	0.640	+51.6	SMITH	1.591	-
				ADJ DE VRIES	1.594	-2.3	GEMANI	1.607	+23.0
				RESISTANCE	1.594	-2.3	VAN RUYEN	1.544	+62.4
				KUNIJS-SMITH	1.594	-2.3	GLOM, MEAN	1.544	+62.4
RHO =	0.6800 TEMP =	20.000 u =	0.6800 CL =	0.6800 CSOLIUS =	+0.000	2.600 ALPHA =	-0.102		
KA =	0.608 X ICE =	2.608 KU =	0.608 KU	2.608 CSOLIUS2 =	+0.000				
SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	DUM, SAT, NC	
SAMPLE	8.3%	1.469	1.469	KERSTEN	1.571	-5.9	MILKLEY	1.747	-10.2
				JOHANSEN	1.593	-2.1	MC GEE	1.745	-0.8
				DE VRIES	0.661	+51.6	SMITH	1.745	-
				ADJ DE VRIES	1.593	-2.1	GEMANI	1.758	+23.0
				RESISTANCE	1.593	-2.1	VAN RUYEN	1.544	+62.4
				KUNIJS-SMITH	1.593	-2.1	GLOM, MEAN	1.544	+62.4
RHO =	0.6800 TEMP =	20.000 u =	0.6800 CL =	0.6800 CSOLIUS =	+0.000	2.600 ALPHA =	-0.103		
KA =	0.608 X ICE =	2.608 KU =	0.608 KU	2.608 CSOLIUS2 =	+0.000				
SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	DUM, SAT, NC	
SAMPLE	8.3%	1.469	1.469	KERSTEN	1.571	-5.9	MILKLEY	1.747	-10.2
				JOHANSEN	1.593	-2.1	MC GEE	1.745	-0.8
				DE VRIES	0.661	+51.6	SMITH	1.745	-
				ADJ DE VRIES	1.593	-2.1	GEMANI	1.757	+23.0
				RESISTANCE	1.593	-2.1	VAN RUYEN	1.544	+62.4
				KUNIJS-SMITH	1.593	-2.1	GLOM, MEAN	1.544	+62.4
RHO =	0.6800 TEMP =	20.000 u =	0.6800 CL =	0.6800 CSOLIUS =	+0.000	2.600 ALPHA =	-0.103		
KA =	0.608 X ICE =	2.608 KU =	0.608 KU	2.608 CSOLIUS2 =	+0.000				
SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	DUM, SAT, NC	
SAMPLE	12.5%	1.757	1.365	KERSTEN	2.420	-53.1	MILKLEY	1.890	-17.2
				JOHANSEN	1.549	+10.1	MC GEE	1.252	+2.7
				DE VRIES	1.509	+1.7	SMITH	1.216	-
				ADJ DE VRIES	1.549	+10.1	GEMANI	1.169	+5.1
				RESISTANCE	1.549	+10.1	VAN RUYEN	1.377	+60.1
				KUNIJS-SMITH	1.549	+10.1	GLOM, MEAN	1.377	+60.1
RHO =	0.6800 TEMP =	20.000 u =	0.6800 CL =	0.6800 CSOLIUS =	+0.000	2.600 ALPHA =	-0.103		
KA =	0.608 X ICE =	2.608 KU =	0.608 KU	2.608 CSOLIUS2 =	+0.000				
SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	DUM, SAT, NC	
SAMPLE	12.5%	1.757	1.365	KERSTEN	2.420	-53.1	MILKLEY	1.890	-17.2
				JOHANSEN	1.549	+10.1	MC GEE	1.252	+2.7
				DE VRIES	1.509	+1.7	SMITH	1.216	-
				ADJ DE VRIES	1.549	+10.1	GEMANI	1.169	+5.1
				RESISTANCE	1.549	+10.1	VAN RUYEN	1.377	+60.1
				KUNIJS-SMITH	1.549	+10.1	GLOM, MEAN	1.377	+60.1
RHO =	0.6800 TEMP =	20.000 u =	0.6800 CL =	0.6800 CSOLIUS =	+0.000	2.600 ALPHA =	-0.103		
KA =	0.608 X ICE =	2.608 KU =	0.608 KU	2.608 CSOLIUS2 =	+0.000				
SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	DUM, SAT, NC	
SAMPLE	12.5%	1.757	1.365	KERSTEN	2.420	-53.1	MILKLEY	1.903	-12.3
				JOHANSEN	1.553	+27.5	MC GEE	1.102	+14.9
				DE VRIES	1.568	+26.4	SMITH	1.260	-
				ADJ DE VRIES	1.553	+27.5	GEMANI	1.216	+12.1
				RESISTANCE	1.553	+27.5	VAN RUYEN	1.374	+72.9
				KUNIJS-SMITH	1.553	+27.5	GLOM, MEAN	1.374	+72.9
RHO =	0.6800 TEMP =	20.000 u =	0.6800 CL =	0.6800 CSOLIUS =	+0.000	2.600 ALPHA =	-0.103		
KA =	0.608 X ICE =	2.608 KU =	0.608 KU	2.608 CSOLIUS2 =	+0.000				
SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	DUM, SAT, NC	
SAMPLE	12.5%	1.757	1.365	KERSTEN	2.420	-53.1	MILKLEY	1.903	-12.3
				JOHANSEN	1.553	+27.5	MC GEE	1.102	+14.9
				DE VRIES	1.568	+26.4	SMITH	1.260	-
				ADJ DE VRIES	1.553	+27.5	GEMANI	1.216	+12.1
				RESISTANCE	1.553	+27.5	VAN RUYEN	1.374	+72.9
				KUNIJS-SMITH	1.553	+27.5	GLOM, MEAN	1.374	+72.9
RHO =	0.6800 TEMP =	20.000 u =	0.6800 CL =	0.6800 CSOLIUS =	+0.000	2.600 ALPHA =	-0.103		
KA =	0.608 X ICE =	2.608 KU =	0.608 KU	2.608 CSOLIUS2 =	+0.000				
SAMPLE	MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	DUM, SAT, NC	
SAMPLE	12.5%	1.757	1.365	KERSTEN	2.420	-53.1	MILKLEY	1.903	-12.3
				JOHANSEN	1.553	+27.5	MC GEE	1.102	+14.9
				DE VRIES	1.568	+26.4	SMITH	1.260	-
				ADJ DE VRIES	1.553	+27.5	GEMANI	1.216	+12.1
				RESISTANCE	1.553	+27.5	VAN RUYEN	1.374	+72.9
				KUNIJS-SMITH	1.553	+27.5	GLOM, MEAN	1.374	+72.9
RHO =	0.6800 TEMP =	20.000 u =	0.6800 CL =	0.6800 CSOLIUS =	+0.000	2.600 ALPHA =	-0.103		
KA =	0.608 X ICE =	2.608 KU =	0.608 KU	2.608 CSOLIUS2 =	+0.000				

Table B3 (cont'd).

SAMPLE	MOISTURE CONTENT	DENSITY K MEASURED	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	POR. SAT.	NC VAL
SAMPLE	15.67	1.894	1.468	-6.1%	1.169	-24.4	2.369	2.079
			KERSTEN	2.11%	0.5%	-11.5	2.812	2.209
			JOHANSEN	1.22%	-1.4%	-		
			DE VRIES	1.35%	-1.4%	-		
ADJ	DE VRIES	-----	-----	-----	SMITH	-----		
			RESISTOR	-----	GEMANT	1.42%	-	
			KUNIT-SMITH	-----	VAN ROOYEN	0.1%	-7.4%	
					GEOM. MEAN	-----		
RHO =	3.000 TEMP =	2.000 K =	0.680 CL =	2.000 CSOLIDS =	2.053 ALPHA =	----- KA =	0.152	
KU =	0.680 KICE =	2.000 K =	0.579 KU =	2.000 CSOLIDS2 =	2.042 ALPHA =	----- KA =	0.152	
SAMPLE	MOISTURE CONTENT	DENSITY K MEASURED	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	POR. SAT.	NC VAL
SAMPLE	15.67	1.894	1.082	-2.1%	1.188	-24.4	2.369	2.079
			KERSTEN	2.11%	0.1%	-24.6	2.812	2.249
			JOHANSEN	1.22%	-2.5%	-		
			DE VRIES	1.27%	-2.5%	-		
ADJ	DE VRIES	-----	-----	-----	SMITH	-----		
			RESISTOR	-----	GEMANT	1.44%	-12.3	
			KUNIT-SMITH	-----	VAN ROOYEN	0.1%	-7.4%	
					GEOM. MEAN	-----		

Table B4.

JOHANSEN CRUSHED ROCK UNFROZEN

TYPE OF SOIL: COARSE CRUSHED UNSATURATED UNFROZEN								
RHO =	3.100 TEMP =	2.511 K =	3.223 RT =	2.000 CSOLIDS =	2.827 ALPHA =	----- KA =	0.053	
KU =	0.680 KICE =	2.000 K =	0.579 KU =	2.000 CSOLIDS2 =	2.840			
SAMPLE	MOISTURE CONTENT	DENSITY K MEASURED	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	POR. SAT.	NC VAL
SAMPLE	3.50	1.894	1.693	-1.0%	0.718	-122.2	0.145	0.052
			KERSTEN	1.15%	15.0	-		
			JOHANSEN	1.35%	13.3	-		
			DE VRIES	1.39%	26.3	-		
ADJ	DE VRIES	-----	-----	-----	SMITH	-----		
			RESISTOR	-----	GEMANT	0.567	19.3	
			KUNIT-SMITH	-----	VAN ROOYEN	0.321	-32.3	
					GEOM. MEAN	-----		
RHO =	3.100 TEMP =	2.911 K =	3.223 RT =	2.000 CSOLIDS =	2.260 ALPHA =	----- KA =	0.053	
KU =	0.670 KICE =	2.000 K =	0.571 KU =	2.000 CSOLIDS2 =	2.840			
SAMPLE	MOISTURE CONTENT	DENSITY K MEASURED	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	POR. SAT.	NC VAL
SAMPLE	3.50	1.894	1.693	-1.0%	0.716	6.5	0.142	0.054
			KERSTEN	1.15%	17.8	-		
			JOHANSEN	1.35%	15.6	-		
			DE VRIES	1.39%	39.9	-		
ADJ	DE VRIES	-----	-----	-----	SMITH	-----		
			RESISTOR	-----	GEMANT	0.575	33.3	
			KUNIT-SMITH	-----	VAN ROOYEN	0.321	-24.0	
					GEOM. MEAN	-----		
RHO =	3.100 TEMP =	2.911 K =	3.223 RT =	2.000 CSOLIDS =	2.259 ALPHA =	----- KA =	0.053	
KU =	0.670 KICE =	2.000 K =	0.571 KU =	2.000 CSOLIDS2 =	2.840			
SAMPLE	MOISTURE CONTENT	DENSITY K MEASURED	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	POR. SAT.	NC VAL
SAMPLE	3.50	1.894	1.693	-1.0%	0.716	6.5	0.142	0.054
			KERSTEN	1.15%	17.8	-		
			JOHANSEN	1.35%	15.6	-		
			DE VRIES	1.39%	39.9	-		
ADJ	DE VRIES	-----	-----	-----	SMITH	-----		
			RESISTOR	-----	GEMANT	0.575	33.3	
			KUNIT-SMITH	-----	VAN ROOYEN	0.321	-24.0	
					GEOM. MEAN	-----		
RHO =	3.100 TEMP =	2.911 K =	3.223 RT =	2.000 CSOLIDS =	2.257 ALPHA =	----- KA =	0.053	
KU =	0.670 KICE =	2.000 K =	0.571 KU =	2.000 CSOLIDS2 =	2.840			
SAMPLE	MOISTURE CONTENT	DENSITY K MEASURED	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	POR. SAT.	NC VAL
SAMPLE	3.50	1.894	1.693	-1.0%	0.726	59.3	0.145	0.054
			KERSTEN	1.15%	15.3	-		
			JOHANSEN	1.35%	19.2	-		
			DE VRIES	1.39%	41.1	-		
ADJ	DE VRIES	-----	-----	-----	SMITH	-----		
			RESISTOR	-----	GEMANT	0.668	66.3	
			KUNIT-SMITH	-----	VAN ROOYEN	0.321	-30.2	
					GEOM. MEAN	-----		
RHO =	3.100 TEMP =	2.911 K =	3.223 RT =	2.000 CSOLIDS =	2.257 ALPHA =	----- KA =	0.053	
KU =	0.670 KICE =	2.000 K =	0.571 KU =	2.000 CSOLIDS2 =	2.840			
SAMPLE	MOISTURE CONTENT	DENSITY K MEASURED	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	POR. SAT.	NC VAL
SAMPLE	3.50	1.894	1.693	-1.0%	0.726	59.3	0.145	0.054
			KERSTEN	1.15%	15.3	-		
			JOHANSEN	1.35%	19.2	-		
			DE VRIES	1.39%	41.1	-		
ADJ	DE VRIES	-----	-----	-----	SMITH	-----		
			RESISTOR	-----	GEMANT	0.728	-16.3	
			KUNIT-SMITH	-----	VAN ROOYEN	0.321	-30.0	
					GEOM. MEAN	-----		

Table B5.

KERSTEN CRUSHED TRAP ROCK UNSATURATED UNFROZEN

TYPE OF SOIL: COARSE CRUSHED UNSATURATED UNFROZEN								
RHO =	2.970 TEMP =	4.500 K =	0.830 CL =	0.100 CSOLIDS =	2.085 ALPHA =	----- KA =	0.054	
KU =	0.670 KICE =	2.000 K =	0.820 KU =	2.000 CSOLIDS2 =	2.800			
SAMPLE	MOISTURE CONTENT	DENSITY K MEASURED	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	METHOD K COMPUTED (PERCENT)	DEVIATION (PERCENT)	POR. SAT.	NC VAL
SAMPLE	0.28	1.843	0.280	-6.137	0.668	131.3	0.447	0.053
			KERSTEN	0.137	MC GAI	0.167	0.445	
			JOHANSEN	0.230	SMITH	-----		
			DE VRIES	0.430	GEMANT	-----		
ADJ	DE VRIES	-----	-----	-----	VAN ROOYEN	0.269	-27.7	
			RESISTOR	-----	GEOM. MEAN	-----		
			KUNIT-SMITH	-----				
SAMPLE	0.28	1.922	0.472	-0.494	NICKLEY	0.824	74.0	0.353
			KERSTEN	0.386	MC GAI	1.355	187.3	0.811
			JOHANSEN	0.586	SMITH	-----		
			DE VRIES	0.588	GEMANT	-----		
ADJ	DE VRIES	-----	-----	-----	VAN ROOYEN	0.298	-36.7	
			RESISTOR	-----	GEOM. MEAN	-----		
			KUNIT-SMITH	-----				

Table B5 (cont'd).

SAMPLE	1.00	1.000	0.382	KERSTEN JOHANSEN DE VRIES RESISTOR KUNIT-SMITH	0.316	-0.17	MICKLEY MC GAW SMITH GEOM. MEAN	1.031	20.8	0.339 + 0.018
ADJ	DE VRIES RESISTOR KUNIT-SMITH	-----	-----	VAN ROOVEN GEOM. MEAN	0.381	20.8				
SAMPLE	1.10	1.019	0.647	KERSTEN JOHANSEN DE VRIES RESISTOR KUNIT-SMITH	0.379	-0.17	MICKLEY MC GAW SMITH GEOM. MEAN	1.037	20.8	0.339 + 0.018
ADJ	DE VRIES RESISTOR KUNIT-SMITH	-----	-----	VAN ROOVEN GEOM. MEAN	0.331	20.8				
SAMPLE	1.00	1.050	0.567	KERSTEN JOHANSEN DE VRIES RESISTOR KUNIT-SMITH	0.328	-0.13	MICKLEY MC GAW SMITH GEOM. MEAN	1.033	20.8	0.339 + 0.018
ADJ	DE VRIES RESISTOR KUNIT-SMITH	-----	-----	VAN ROOVEN GEOM. MEAN	0.328	20.8				
SAMPLE	1.70	1.927	0.672	KERSTEN JOHANSEN DE VRIES RESISTOR KUNIT-SMITH	0.393	-0.13	MICKLEY MC GAW SMITH GEOM. MEAN	1.033	20.8	0.339 + 0.018
ADJ	DE VRIES RESISTOR KUNIT-SMITH	-----	-----	VAN ROOVEN GEOM. MEAN	0.330	20.8				
SAMPLE	3.00	1.653	0.737	KERSTEN JOHANSEN DE VRIES RESISTOR KUNIT-SMITH	0.291	-0.08	MICKLEY MC GAW SMITH GEOM. MEAN	1.038	20.8	0.342 + 0.018
ADJ	DE VRIES RESISTOR KUNIT-SMITH	-----	-----	VAN ROOVEN GEOM. MEAN	0.337	20.8				
SAMPLE	3.70	1.927	0.992	KERSTEN JOHANSEN DE VRIES RESISTOR KUNIT-SMITH	1.070	-0.08	MICKLEY MC GAW SMITH GEOM. MEAN	1.033	20.8	0.339 + 0.018
ADJ	DE VRIES RESISTOR KUNIT-SMITH	-----	-----	VAN ROOVEN GEOM. MEAN	0.330	20.8				

Table B6.

KERSTEN DAKUTA SANDY LOAM UNSATURATED UNFROZEN

TYPE OF SAMPLE	DENSITY	K MEASURED	METHOD	COMPUTED DEVIATION (PERCENT)	METHOD	COMPUTED DEVIATION (PERCENT)	PER.	NC	RHO = 2,710 TDR = 4.500 K = 0.501 CL = 0.100 CROLIDE = 4.545 ALPHA = ----- KA = 0.094	
									RHO = 0.549 KICF = + 2,200 K4 = 6,020 KO = 2,000 CROLIDP = 5.508	
SAMPLE	WET SURFACE	DRY	KERSTEN	0.499	MICKLEY	1.218	304.3	0.501	0.026	
SAMPLE	1.00	1.352	0.245	JOHANSEN	0.314	MC GAW	1.802	639.0	0.051	0.026
			DE VRIES	0.460	SMITH	1.802				
			ADJ	DE VRIES	-----	-----				
			RESISTOR	-----	-----					
			KUNIT-SMITH	-----	VAN ROOVEN	0.304	24.1			
SAMPLE	2.10	1.600	0.460	KERSTEN	0.900	95.6	MICKLEY	1.520	232.2	0.410 + 0.034
			JOHANSEN	1.048	MC GAW	2.081	359.2	0.082	0.034	
			DE VRIES	0.937	SMITH	0.938	18.8			
			ADJ	DE VRIES	-----	-----				
			RESISTOR	-----	VAN ROOVEN	0.343				
			KUNIT-SMITH	-----	GEOM. MEAN					
SAMPLE	1.00	1.764	0.554	KERSTEN	1.083	95.6	MICKLEY	1.754	316.8	0.369 + 0.034
			JOHANSEN	1.046	MC GAW	2.334	321.5	0.086	0.034	
			DE VRIES	1.139	SMITH	0.623	33.3			
			ADJ	DE VRIES	-----	-----				
			RESISTOR	-----	VAN ROOVEN	0.680				
			KUNIT-SMITH	-----	GEOM. MEAN					
SAMPLE	2.10	1.971	0.887	KERSTEN	1.225	43.6	MICKLEY	3.083	181.8	0.738 + 0.028
			JOHANSEN	1.243	MC GAW	3.083	181.8	0.738	0.028	
			DE VRIES	1.443	SMITH	1.019	19.3			
			ADJ	DE VRIES	-----	-----				
			RESISTOR	-----	VAN ROOVEN	0.852				
			KUNIT-SMITH	-----	GEOM. MEAN					
SAMPLE	3.40	1.747	0.789	KERSTEN	1.611	78.0	MICKLEY	1.788	136.7	0.348 + 0.051
			JOHANSEN	1.187	MC GAW	2.145	132.7	0.173	0.060	
			DE VRIES	1.526	SMITH	1.169	48.3			
			ADJ	DE VRIES	-----	-----				
			RESISTOR	-----	VAN ROOVEN	0.852				
			KUNIT-SMITH	-----	GEOM. MEAN					
SAMPLE	3.60	1.927	1.289	KERSTEN	1.817	52.0	MICKLEY	2.055	89.4	0.340 + 0.048
			JOHANSEN	1.575	MC GAW	2.443	132.7	0.173	0.060	
			DE VRIES	1.670	SMITH	1.607	24.6			
			ADJ	DE VRIES	-----	-----				
			RESISTOR	-----	VAN ROOVEN	0.937				
			KUNIT-SMITH	-----	GEOM. MEAN					
SAMPLE	3.50	2.084	1.051	KERSTEN	2.088	76.6	MICKLEY	2.328	43.4	0.331 + 0.039
			JOHANSEN	2.074	MC GAW	2.328	43.4	0.331	0.039	
			DE VRIES	2.074	SMITH	2.036	39.3			
			ADJ	DE VRIES	-----	-----				
			RESISTOR	-----	VAN ROOVEN	1.032				
			KUNIT-SMITH	-----	GEOM. MEAN					
SAMPLE	4.60	1.969	1.518	KERSTEN	2.007	92.7	MICKLEY	2.143	67.3	0.274 + 0.076
			JOHANSEN	1.781	MC GAW	2.326	67.3	0.274	0.076	
			DE VRIES	1.452	SMITH	1.813	37.7			
			ADJ	DE VRIES	-----	-----				
			RESISTOR	-----	VAN ROOVEN	0.971				
			KUNIT-SMITH	-----	GEOM. MEAN					
SAMPLE	4.60	1.924	1.478	KERSTEN	2.023	78.0	MICKLEY	2.053	83.3	0.329 + 0.026
			JOHANSEN	1.911	MC GAW	2.450	83.3	0.329	0.026	
			DE VRIES	1.911	SMITH	1.873	35.9			
			ADJ	DE VRIES	-----	-----				
			RESISTOR	-----	VAN ROOVEN	1.058				
			KUNIT-SMITH	-----	GEOM. MEAN					

Table B6 (cont'd).

SAMPLE	4.90	2.111	4.271	KERSTEN JOHANSEN DE VRIES RESISTOR KUNI-SMITH	2.948 2.926 2.926 2.926 2.926	16.4 2.1 2.1 2.1 2.1	NICKLEY MC GAW SMITH VAN ROOYEN GEOM. MEAN	2.169 2.041 2.041 2.041 2.041	-10.7 -10.7 -10.7 -10.7 -10.7	0.321 + 0.038
SAMPLE	4.30	2.194	4.722	KERSTEN JOHANSEN DE VRIES RESISTOR KUNI-SMITH	2.956 2.920 2.920 2.920 2.920	19.1 2.0 2.0 2.0 2.0	NICKLEY MC GAW SMITH VAN ROOYEN GEOM. MEAN	2.169 2.041 2.041 2.041 2.041	-12.5 -12.5 -12.5 -12.5 -12.5	0.329 + 0.038
SAMPLE	8.90	1.924	4.164	KERSTEN JOHANSEN DE VRIES RESISTOR KUNI-SMITH	2.439 2.148 2.221 2.221 2.221	12.7 20.4 19.4 19.4 19.4	NICKLEY MC GAW SMITH VAN ROOYEN GEOM. MEAN	2.209 2.149 2.149 2.149 2.149	-15.0 -15.0 -15.0 -15.0 -15.0	0.390 + 0.046
SAMPLE	9.30	4.076	4.040	KERSTEN JOHANSEN DE VRIES RESISTOR KUNI-SMITH	3.073 2.450 2.718 2.718 2.718	16.5 2.0 2.0 2.0 2.0	NICKLEY MC GAW SMITH VAN ROOYEN GEOM. MEAN	2.169 2.041 2.041 2.041 2.041	-11.2 -11.2 -11.2 -11.2 -11.2	0.325 + 0.038
SAMPLE	12.60	1.914	4.384	KERSTEN JOHANSEN DE VRIES RESISTOR KUNI-SMITH	2.696 2.392 2.452 2.452 2.452	14.1 1.9 3.7 3.7 3.7	NICKLEY MC GAW SMITH VAN ROOYEN GEOM. MEAN	2.169 2.041 2.041 2.041 2.041	-1.5 -1.5 -10.9 -10.9 -10.9	0.366 + 0.046

Table B7.

JUHANSEN SAND Saturated UNFROZEN

TYPE OF SOIL: SAND NATURAL SATURATED UNFROZEN											
RHO =	2.600	TEMP =	5.50	CL =	0.630	CSOLIDS =	+	5.895	ALPHA =	----- KA = + 0.025	
KJ =	0.369	KICF =	+	2.200	KJ =	3.773	XJ =	2.000	CSOLID82 =	5.840	
SAMPLE	WETSTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	POR.	NC		
	18.00	1.771	2.104	KERSTEN JOHANSEN DE VRIES RESISTOR KUNI-SMITH	2.352 2.104 2.104 2.387 2.089	11.1 0.0 0.0 13.0 -1.0	NICKLEY MC GAW SMITH VAN ROOYEN GEOM. MEAN	2.165 2.041 2.041 2.110 2.110	-3.0 -3.0 -3.0 -6.4 0.0	0.149 0.000 + 0.150	
RHO =	2.600	TEMP =	22.877	CL =	0.630	CSOLIDS =	+	5.772	ALPHA =	----- KA = + 0.025	
KJ =	0.369	KICF =	+	2.200	KJ =	7.910	XJ =	2.000	CSOLID82 =	5.840	
SAMPLE	WETSTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	POR.	NC		
	18.00	1.771	2.145	KERSTEN JOHANSEN DE VRIES RESISTOR KUNI-SMITH	2.352 2.104 2.104 2.368 2.089	9.0 -2.0 -2.0 10.4 -2.8	NICKLEY MC GAW SMITH VAN ROOYEN GEOM. MEAN	2.146 2.041 2.041 2.110 2.110	-0.1 -0.1 -0.1 -7.0 -2.1	0.149 0.000 + 0.150	

Table B8.

JOHANSEN GRAVEL GR3 SATURATED UNFROZEN

TYPE OF SOIL: COARSE NATURAL SATURATED UNFROZEN											
RHO =	2.700	TEMP =	5.100	CL =	0.570	CSOLIDS =	+	5.839	ALPHA =	----- KA = + 0.025	
KJ =	0.369	KICF =	+	2.200	KJ =	8.578	XJ =	2.000	CSOLID82 =	5.840	
SAMPLE	WETSTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	POR.	NC		
	14.50	1.944	2.600	KERSTEN JOHANSEN DE VRIES RESISTOR KUNI-SMITH	2.331 2.334 2.334 2.361 2.361	10.0 -2.8 -2.8 -7.8 -7.8	NICKLEY MC GAW SMITH VAN ROOYEN GEOM. MEAN	2.328 2.328 2.328 2.328 2.328	-0.1 -0.1 -0.1 -7.8 -7.8	0.060 + 0.268	
RHO =	2.700	TEMP =	22.600	CL =	0.570	CSOLIDS =	+	5.722	ALPHA =	----- KA = + 0.025	
KJ =	0.369	KICF =	+	2.200	KJ =	7.910	XJ =	2.000	CSOLID82 =	5.840	
SAMPLE	WETSTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	POR.	NC		
	14.50	1.944	2.615	KERSTEN JOHANSEN DE VRIES RESISTOR KUNI-SMITH	2.333 2.321 2.321 2.361 2.361	10.0 -3.8 -3.8 -8.8 -8.8	NICKLEY MC GAW SMITH VAN ROOYEN GEOM. MEAN	2.312 2.312 2.312 2.312 2.312	-0.1 -0.1 -0.1 -7.8 -7.8	0.060 + 0.268	

Table B9.

JOHANSEN GRAVEL Saturated J-FROZEN

TYPE OF SOIL: COARSE NATURAL SATURATED UNFROZEN							
RHO =	3,000 TEMP =	2,700 K =	3,273 C =	2,000 CSOLIDS =	2,056 ALPHA = ----- KA = + 0,025	POR.	NC
KU =	0,373 KICE =	2,200 KQ =	3,273 KJ =	2,000 CSOLIDS2 =	3,840	SAT.	0,234
SAMPLE MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTER	DEVIATION (PERCENT)	METHOD K COMPUTER	DEVIATION (PERCENT)	POR. NC
SAMPLE 10.20	2,297	1,722	KERSTEN	-51.6	HICKLEY	-19.7	0,234
			JOHANSEN	-11.4	MC GAV	-10.1	
			DE VRIES	-79.1	SMITH	-72.2	
			ADJ RESISTOR	-1.600	VAN ROOYEN	-1.200	
			KUNIYASUITH	-1.624	GEOM. MEAN	-1.525	
RHO =	3,000 TEMP =	2,700 K =	3,273 C =	2,000 CSOLIDS =	2,056 ALPHA = ----- KA = + 0,025		
KU =	0,373 KICE =	2,200 KQ =	3,273 KJ =	2,000 CSOLIDS2 =	3,840		
SAMPLE MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTER	DEVIATION (PERCENT)	METHOD K COMPUTER	DEVIATION (PERCENT)	POR. NC
SAMPLE 10.20	2,297	1,743	KERSTEN	-4.333	HICKLEY	-16.7	0,234
			JOHANSEN	-1.334	MC GAV	-1.332	
			DE VRIES	-1.273	SMITH	-1.271	
			ADJ RESISTOR	-1.613	VAN ROOYEN	-1.200	
			KUNIYASUITH	-1.633	GEOM. MEAN	-1.534	
RHO =	3,000 TEMP =	2,700 K =	3,273 C =	2,000 CSOLIDS =	2,056 ALPHA = ----- KA = + 0,025		
KU =	0,373 KICE =	2,200 KQ =	3,273 KJ =	2,000 CSOLIDS2 =	3,840		
SAMPLE MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTER	DEVIATION (PERCENT)	METHOD K COMPUTER	DEVIATION (PERCENT)	POR. NC
SAMPLE 10.20	2,297	1,727	KERSTEN	-4.333	HICKLEY	-11.6	0,234
			JOHANSEN	-1.336	MC GAV	-1.330	
			DE VRIES	-1.574	SMITH	-1.276	
			ADJ RESISTOR	-1.733	VAN ROOYEN	-1.900	
			KUNIYASUITH	-1.659	GEOM. MEAN	-1.330	

Table B10.

JOHANSEN CRUSHED ROCK PJD SATURATED UNFROZEN

TYPE OF SOIL: COARSE CRUSHED SATURATED UNFROZEN							
RHO =	2,700 TEMP =	2,900 K =	3,330 C =	2,000 CSOLIDS =	3,082 ALPHA = ----- KA = + 0,025	POR.	NC
KU =	0,366 KICE =	2,200 KQ =	3,330 KJ =	2,000 CSOLIDS2 =	3,840	SAT.	0,277
SAMPLE MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTER	DEVIATION (PERCENT)	METHOD K COMPUTER	DEVIATION (PERCENT)	POR. NC
SAMPLE 10.20	1,931	1,739	KERSTEN	-89.4	HICKLEY	-9.0	0,277
			JOHANSEN	-9.6	MC GAV	-13.3	
			DE VRIES	-13.5	SMITH	-13.5	
			ADJ RESISTOR	-2.077	VAN ROOYEN	-3.6	
			KUNIYASUITH	-2.016	GEOM. MEAN	-9.6	
RHO =	2,700 TEMP =	2,900 K =	3,330 C =	2,000 CSOLIDS =	3,079 ALPHA = ----- KA = + 0,025		
KU =	0,366 KICE =	2,200 KQ =	3,330 KJ =	2,000 CSOLIDS2 =	3,840		
SAMPLE MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTER	DEVIATION (PERCENT)	METHOD K COMPUTER	DEVIATION (PERCENT)	POR. NC
SAMPLE 10.20	1,931	1,807	KERSTEN	-89.4	HICKLEY	-10.4	0,277
			JOHANSEN	-6.7	MC GAV	-1.0	
			DE VRIES	-10.6	SMITH	-10.6	
			ADJ RESISTOR	-2.177	VAN ROOYEN	-2.000	
			KUNIYASUITH	-2.088	GEOM. MEAN	-9.0	

Table B11.

JOHANSEN CRUSHED ROCK PUD SATURATED UNFROZEN

TYPE OF SOIL: COARSE CRUSHED SATURATED UNFROZEN							
RHO =	3,120 TEMP =	2,900 K =	3,330 C =	2,000 CSOLIDS =	2,265 ALPHA = ----- KA = + 0,025	POR.	NC
KU =	0,372 KICE =	2,200 KQ =	3,330 KJ =	2,000 CSOLIDS2 =	3,840	SAT.	0,492
SAMPLE MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTER	DEVIATION (PERCENT)	METHOD K COMPUTER	DEVIATION (PERCENT)	POR. NC
SAMPLE 10.20	1,938	1,737	KERSTEN	-89.3	HICKLEY	-8.3	0,492
			JOHANSEN	-8.7	MC GAV	-8.3	
			DE VRIES	-13.7	SMITH	-13.7	
			ADJ RESISTOR	-2.188	VAN ROOYEN	-2.083	
			KUNIYASUITH	-2.098	GEOM. MEAN	-10.2	
RHO =	3,120 TEMP =	2,900 K =	3,330 C =	2,000 CSOLIDS =	2,262 ALPHA = ----- KA = + 0,025		
KU =	0,372 KICE =	2,200 KQ =	3,330 KJ =	2,000 CSOLIDS2 =	3,840		
SAMPLE MOISTURE CONTENT	DRY DENSITY	K MEASURED	METHOD K COMPUTER	DEVIATION (PERCENT)	METHOD K COMPUTER	DEVIATION (PERCENT)	POR. NC
SAMPLE 10.20	1,938	1,930	KERSTEN	-89.3	HICKLEY	-10.4	0,492
			JOHANSEN	-8.7	MC GAV	-8.3	
			DE VRIES	-13.7	SMITH	-13.7	
			ADJ RESISTOR	-2.178	VAN ROOYEN	-2.083	
			KUNIYASUITH	-2.198	GEOM. MEAN	-10.4	

Table B12.

“SSS BUILDING CODE (1999) MAY 2011” XSAFRAIPE (NERSI)

Table B12 (cont'd).

S-112	27.10	1,111	1,142	KERSTEN	1.667	318.8	MICKLEY	0.819	-3.3	81203 + 81203
				ADJ. DE VRIES	1.676	324.4	SMITH	0.811	-3.8	
				KUNITSCHMIDT	1.678	324.4	GERMANY	0.811	-3.8	
				KUNITSCHMIDT	1.678	324.4	VAN ROOYEN	0.828	-3.4	
S-113	27.10	1,121	1,131	KERSTEN	1.684	313.5	MICKLEY	0.849	-3.2	0.366 + 0.063
				ADJ. DE VRIES	1.686	323.8	MC GAW	0.849	-3.0	
				KUNITSCHMIDT	1.686	323.8	GERMANY	0.849	-3.2	
				KUNITSCHMIDT	1.686	323.8	VAN ROOYEN	0.849	-3.1	
S-114	27.10	1,125	1,132	KERSTEN	1.694	323.3	MICKLEY	0.733	-6.3	0.365 + 0.063
				ADJ. DE VRIES	1.695	322.8	MC GAW	0.733	-6.2	
				KUNITSCHMIDT	1.695	322.8	SMITH	0.733	-6.2	
				KUNITSCHMIDT	1.695	322.8	VAN ROOYEN	0.729	-6.3	
S-115	27.10	1,131	1,132	KERSTEN	1.694	321.7	MICKLEY	0.830	-4.5	0.366 + 0.063
				ADJ. DE VRIES	1.695	321.9	MC GAW	0.830	-4.5	
				KUNITSCHMIDT	1.695	321.9	GERMANY	0.830	-4.8	
				KUNITSCHMIDT	1.695	321.9	VAN ROOYEN	0.830	-4.7	
S-116	27.10	1,131	1,133	KERSTEN	1.714	276.6	MICKLEY	0.885	-43.2	0.691 + 0.066
				ADJ. DE VRIES	1.717	277.3	MC GAW	0.879	-50.7	
				KUNITSCHMIDT	1.717	276.6	GERMANY	0.873	-53.8	
				KUNITSCHMIDT	1.717	276.6	VAN ROOYEN	0.852	-45.3	
S-117	27.10	1,131	1,144	KERSTEN	1.734	259.3	MICKLEY	1.012	-23.7	0.855 + 0.063
				ADJ. DE VRIES	1.734	259.3	MC GAW	1.012	-23.7	
				KUNITSCHMIDT	1.734	259.3	SMITH	1.030	-25.0	
				KUNITSCHMIDT	1.734	259.3	VAN ROOYEN	0.926	-24.0	
S-118	27.10	1,131	1,144	KERSTEN	1.734	255.9	MICKLEY	0.723	-35.6	0.733 + 0.063
				ADJ. DE VRIES	1.734	256.8	MC GAW	0.723	-35.6	
				KUNITSCHMIDT	1.734	256.8	SMITH	0.731	-36.1	
				KUNITSCHMIDT	1.734	256.8	VAN ROOYEN	0.670	-42.9	
S-119	27.10	1,131	1,144	KERSTEN	1.744	249.4	MICKLEY	0.780	-41.2	0.582 + 0.062
				ADJ. DE VRIES	1.747	251.8	MC GAW	0.957	-27.8	
				KUNITSCHMIDT	1.747	251.8	SMITH	0.823	-33.8	
				KUNITSCHMIDT	1.747	251.8	VAN ROOYEN	0.823	-33.8	
S-120	27.10	1,131	1,151	KERSTEN	1.677	335.9	MICKLEY	0.863	-33.9	81202 + 81203
				ADJ. DE VRIES	1.682	336.8	MC GAW	0.853	-34.4	
				KUNITSCHMIDT	1.682	336.8	SMITH	0.751	-36.1	
				KUNITSCHMIDT	1.682	336.8	VAN ROOYEN	0.670	-42.9	
S-121	27.10	1,131	1,151	KERSTEN	1.914	179.4	MICKLEY	0.780	-41.2	0.791 + 0.060
				ADJ. DE VRIES	1.917	179.4	MC GAW	0.957	-27.8	
				KUNITSCHMIDT	1.917	179.4	SMITH	0.823	-33.8	
				KUNITSCHMIDT	1.917	179.4	VAN ROOYEN	0.823	-33.8	
S-122	27.10	1,131	1,161	KERSTEN	1.652	333.9	MICKLEY	0.863	-33.9	81202 + 81203
				ADJ. DE VRIES	1.652	334.2	MC GAW	0.858	-32.8	
				KUNITSCHMIDT	1.652	334.2	SMITH	0.750	-36.1	
				KUNITSCHMIDT	1.652	334.2	VAN ROOYEN	0.670	-42.9	
S-123	27.10	1,131	1,162	KERSTEN	1.161	310.2	MICKLEY	1.022	-36.5	0.986 + 0.060
				ADJ. DE VRIES	1.162	318.4	MC GAW	1.046	-34.7	
				KUNITSCHMIDT	1.162	318.4	SMITH	0.887	-36.9	
				KUNITSCHMIDT	1.162	318.4	VAN ROOYEN	0.887	-36.9	
S-124	27.10	1,131	1,167	KERSTEN	1.247	252.1	MICKLEY	1.073	-23.7	0.868 + 0.063
				ADJ. DE VRIES	1.247	242.1	MC GAW	1.073	-23.7	
				KUNITSCHMIDT	1.247	242.1	SMITH	1.080	-23.8	
				KUNITSCHMIDT	1.247	242.1	VAN ROOYEN	1.080	-23.8	

Table B13.

KUNTEK HEAVY CLAY SATURATED UNFROZEN											
TYPE OF SAMPLE			CLAY NATURE SATURATED UNFROZEN			COMPUTED					
NAME	MOISTURE	DRY	MEASURED	METHOD	COMPUTED	DEVIATION	METHOD	COMPUTED	DEVIATION	POR.	NC
SAMPLE 1 100% H2O	1.243	1.243	KERSTEN	1.576	2.1%	MICKLEY	1.572	2.1%	0.350	0.330	
SAMPLE 2 100% H2O	1.243	1.243	JOHANSEN	1.548	0.3%	MC GAW	1.626	-5.5	1.000	0.330	
SAMPLE 3 100% H2O	1.243	1.243	DE VRIES	1.622	5.1%	SMITH	-	-	-	-	
SAMPLE 4 100% H2O	1.243	1.243	RESISTOR	-	-	GERMANY	0.510	-53.1	-	-	
SAMPLE 5 100% H2O	1.243	1.243	KUNITSCHMIDT	1.601	33.6	VAN ROOYEN	1.548	-5.0	-	-	

Table B14.

MENNER LETA CLAY SATURATED UNFROZEN

TYPE OF SOIL: CLAY NATURAL SATURATED UNFROZEN				ALPHA = ----- KA = + 0.025									
RHO =	1.210	TEMP =	5.000	RHO =	1.000	TEMP =	5.000						
SAMPLE	25.00	1.300	0.900	KERSTEN	1.060	16.3	MICKLEY	0.916	12.1	POR.	0.391	WC	0.391
				JOHANSEN	0.903	16.0	MC GAW	0.948	18.5		1.000	+ 0.371	
				DE VRIES	0.941	17.5	SMITH	0.938	13.2				
				ADJ RESISTOR	0.993	24.1	VAN ROOYEN	0.938	13.2				
				KUNII-SMITH	0.818	2.2	SEOM. MEAN	-----	-----				
SAMPLE	35.00	1.300	0.900	KERSTEN	1.092	49.9	MICKLEY	0.988	8.7	POR.	0.305	WC	0.305
				JOHANSEN	1.011	11.8	MC GAW	0.918	12.6		1.000	+ 0.371	
				DE VRIES	1.021	11.8	SMITH	0.938	13.2				
				ADJ RESISTOR	1.092	18.7	VAN ROOYEN	0.938	13.2				
				KUNII-SMITH	0.882	19.0	SEOM. MEAN	-----	-----				
SAMPLE	55.00	1.300	0.900	KERSTEN	1.063	40.5	MICKLEY	1.023	9.4	POR.	0.608	WC	0.608
				JOHANSEN	1.036	12.6	MC GAW	0.930	13.1		1.000	+ 0.608	
				DE VRIES	1.036	12.6	SMITH	0.930	13.1				
				ADJ RESISTOR	1.092	18.7	VAN ROOYEN	0.948	13.0				
				KUNII-SMITH	0.921	0.1	SEOM. MEAN	-----	-----				
SAMPLE	54.00	1.400	0.900	KERSTEN	1.063	43.8	MICKLEY	0.990	6.5	POR.	0.498	WC	0.498
				JOHANSEN	0.980	3.6	MC GAW	1.026	10.3		1.000	+ 0.498	
				DE VRIES	1.019	2.5	SMITH	0.955	2.7				
				ADJ RESISTOR	1.074	12.5	VAN ROOYEN	0.955	2.7				
				KUNII-SMITH	0.901	15.6	SEOM. MEAN	0.962	3.4				
SAMPLE	25.10	1.500	0.900	KERSTEN	1.064	47.8	MICKLEY	1.045	3.5	POR.	0.667	WC	0.667
				JOHANSEN	1.057	4.7	MC GAW	1.042	3.2		1.000	+ 0.667	
				DE VRIES	1.076	8.6	SMITH	-----	-----				
				ADJ RESISTOR	1.132	14.3	VAN ROOYEN	0.939	3.8				
				KUNII-SMITH	0.969	2.2	SEOM. MEAN	-----	-----				

Table B15.

GAMES AND PIGUE SAMPLES SATURATED UNFROZEN CLAY

TYPE OF SOIL: CLAY NATURAL SATURATED UNFROZEN				ALPHA = ----- KA = + 0.025									
RHO =	1.210	TEMP =	5.000	RHO =	1.000	TEMP =	5.000						
SAMPLE	25.30	1.303	0.907	KERSTEN	1.517	10.872	MICKLEY	1.230	13.7	POR.	0.409	WC	0.409
				JOHANSEN	1.220	26.2	MC GAW	1.275	22.0		1.000	+ 0.409	
				DE VRIES	1.260	23.2	SMITH	-----	-----				
				ADJ RESISTOR	1.562	18.6	VAN ROOYEN	1.089	27.1				
				KUNII-SMITH	1.153	30.2	SEOM. MEAN	1.220	26.2				
SAMPLE	25.90	1.403	1.175	KERSTEN	1.391	18.8	MICKLEY	1.179	0.8	POR.	0.888	WC	0.888
				JOHANSEN	1.166	4.8	MC GAW	1.225	4.8		1.000	+ 0.888	
				DE VRIES	1.216	3.5	SMITH	-----	-----				
				ADJ RESISTOR	1.390	12.8	VAN ROOYEN	1.089	27.1				
				KUNII-SMITH	1.081	27.8	SEOM. MEAN	1.166	20.8				
SAMPLE	27.20	1.564	1.486	KERSTEN	1.495	0.9	MICKLEY	1.298	13.8	POR.	0.887	WC	0.887
				JOHANSEN	1.194	10.6	MC GAW	1.349	13.8		1.000	+ 0.887	
				DE VRIES	1.264	16.3	SMITH	-----	-----				
				ADJ RESISTOR	1.377	21.3	VAN ROOYEN	0.939	28.8				
				KUNII-SMITH	1.120	22.0	SEOM. MEAN	1.194	20.8				
SAMPLE	30.70	1.458	1.429	KERSTEN	1.324	-7.4	MICKLEY	1.191	-16.7	POR.	0.437	WC	0.437
				JOHANSEN	1.170	-17.5	MC GAW	1.235	-13.6		1.000	+ 0.437	
				DE VRIES	1.228	-14.0	SMITH	-----	-----				
				ADJ RESISTOR	1.303	-8.8	VAN ROOYEN	0.879	-27.4				
				KUNII-SMITH	1.100	-23.0	SEOM. MEAN	1.170	-17.5				
SAMPLE	25.00	1.556	1.538	KERSTEN	1.424	-7.9	MICKLEY	1.338	-19.3	POR.	0.886	WC	0.886
				JOHANSEN	1.250	-13.3	MC GAW	1.379	-13.3		1.000	+ 0.886	
				DE VRIES	1.290	-13.3	SMITH	1.153	-25.1				
				ADJ RESISTOR	1.382	-23.2	VAN ROOYEN	0.839	-27.9				

Table B16.

DAMES AND MOORE SAMPLES SATURATED UNFROZEN SILT

TYPE OF SOIL		SILT	NATURAL	SATURATED	UNFROZEN	RHO =	2780	TEMP =	24.000 °C	K =	0.000 CL =	0.000 CSOLID5 =	1,000 CSOLID52 =	2,000 CSOLID52 =	2,000 ALPHA =	-----	KA =	+	0.026	
SAMPLE	LOCATION	WATER	DENSITY	K MEASURED	METHOD	K COMPUTED	DEVIATION	(PERCENT)	MICKLEY	MC GAV	SMITH	GEMANT	VAN ROOYEN	GEOM. MEAN	MICKLEY	MC GAV	SMITH	GEMANT	VAN ROOYEN	GEOM. MEAN
SAMPLE	25.92	1.591	1.491	KERSTEN	1.521	-2.3	MICKLEY	1.235	-17.3	MC GAV	1.291	-14.2	GEMANT	1.218	-16.8	MC GAV	1.212	-16.8	GEMANT	1.212
				JOHANSEN	1.215	-17.6														
				DE VRIES	1.265	-14.6	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	1.298	-18.6	GEOM. MEAN	1.218	-16.8								
				KUNTZ-SMITH	1.147	-22.6														
SAMPLE	25.92	1.591	1.491	KERSTEN	1.521	-2.3	MICKLEY	1.235	-17.3	MC GAV	1.291	-14.2	GEMANT	1.218	-16.8	MC GAV	1.212	-16.8	GEMANT	1.212
				JOHANSEN	1.215	-17.6														
				DE VRIES	1.265	-14.6	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	1.298	-18.6	GEOM. MEAN	1.218	-16.8								
				KUNTZ-SMITH	1.147	-22.6														
SAMPLE	29.42	1.592	1.592	KERSTEN	1.441	-7.8	MICKLEY	1.177	-24.8	MC GAV	1.221	-21.2	GEMANT	1.221	-21.2	MC GAV	1.247	-20.8	GEMANT	1.247
				JOHANSEN	1.184	-24.9														
				DE VRIES	1.214	-21.4	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	1.166	-27.5	GEOM. MEAN	1.194	-24.9								
				KUNTZ-SMITH	1.082	-32.2														

Table B17.

EFFECTEN, PAPUA SANDY LOAM UNSATURATED FROZEN

TYPE OF SOIL		SILT	NATURAL	UNSATURATED	FROZEN	RHO =	56.00	TEMP =	56.000 °C	K =	0.501 CL =	0.100 CSOLID5 =	2,000 CSOLID52 =	4,034 ALPHA =	-----	KA =	0.046	0.508		
SAMPLE	LOCATION	WATER	DENSITY	K MEASURED	METHOD	K COMPUTED	DEVIATION	(PERCENT)	MICKLEY	MC GAV	SMITH	GEMANT	VAN ROOYEN	GEOM. MEAN	MICKLEY	MC GAV	SMITH	GEMANT	VAN ROOYEN	GEOM. MEAN
SAMPLE	25.92	1.592	1.490	KERSTEN	0.287	-16.0	MICKLEY	1.271	-269.4	MC GAV	1.291	-269.4	SMITH	0.051	-	0.056	-	0.000	-	
				JOHANSEN	0.344	-19.4														
				DE VRIES	0.226	-211.4	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	-	-	GEOM. MEAN	-	-								
				KUNTZ-SMITH	-	-														
SAMPLE	25.92	1.592	1.490	KERSTEN	0.498	-8.2	MICKLEY	1.408	-250.7	MC GAV	1.489	-250.7	SMITH	0.610	-	0.689	-	0.000	-	
				JOHANSEN	0.528	-15.2														
				DE VRIES	1.224	-147.4	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	-	-	GEOM. MEAN	-	-								
				KUNTZ-SMITH	-	-														
SAMPLE	25.92	1.592	1.490	KERSTEN	0.451	-19.4	MICKLEY	1.445	-238.4	MC GAV	1.405	-238.4	SMITH	0.340	-	0.340	-	0.000	-	
				JOHANSEN	0.425	-18.1														
				DE VRIES	1.276	-190.4	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	-	-	GEOM. MEAN	-	-								
				KUNTZ-SMITH	-	-														
SAMPLE	25.92	1.592	1.490	KERSTEN	0.462	-6.7	MICKLEY	2.126	-140.0	MC GAV	2.151	-140.0	SMITH	0.261	-	0.261	-	0.000	-	
				JOHANSEN	0.485	-0.2														
				DE VRIES	1.400	-115.1	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	-	-	GEOM. MEAN	-	-								
				KUNTZ-SMITH	-	-														
SAMPLE	25.92	1.592	1.490	KERSTEN	0.418	-20.4	MICKLEY	1.932	-153.2	MC GAV	1.943	-153.2	SMITH	0.348	-	0.348	-	0.000	-	
				JOHANSEN	0.405	-18.7														
				DE VRIES	1.467	-144.4	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	-	-	GEOM. MEAN	-	-								
				KUNTZ-SMITH	-	-														
SAMPLE	25.92	1.592	1.490	KERSTEN	1.419	-10.4	MICKLEY	2.244	-77.4	MC GAV	2.254	-77.4	SMITH	0.249	-	0.249	-	0.000	-	
				JOHANSEN	1.470	-11.4														
				DE VRIES	2.140	-85.1	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	-	-	GEOM. MEAN	-	-								
				KUNTZ-SMITH	-	-														
SAMPLE	25.92	1.592	1.490	KERSTEN	1.454	-20.1	MICKLEY	2.344	-82.8	MC GAV	2.306	-82.8	SMITH	0.276	-	0.276	-	0.000	-	
				JOHANSEN	1.465	-11.7														
				DE VRIES	2.165	-88.4	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	-	-	GEOM. MEAN	-	-								
				KUNTZ-SMITH	-	-														
SAMPLE	25.92	1.592	1.490	KERSTEN	1.480	-92.0	MICKLEY	2.350	-70.4	MC GAV	2.354	-70.4	SMITH	0.290	-	0.290	-	0.000	-	
				JOHANSEN	1.577	-16.4														
				DE VRIES	2.770	-101.4	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	-	-	GEOM. MEAN	-	-								
				KUNTZ-SMITH	-	-														
SAMPLE	25.92	1.592	1.490	KERSTEN	2.464	-6.4	MICKLEY	2.829	-28.4	MC GAV	2.820	-28.4	SMITH	0.310	-	0.310	-	0.000	-	
				JOHANSEN	2.269	-23.5														
				DE VRIES	3.414	-65.2	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	-	-	GEOM. MEAN	-	-								
				KUNTZ-SMITH	-	-														
SAMPLE	25.92	1.592	1.490	KERSTEN	2.464	-6.4	MICKLEY	2.829	-28.4	MC GAV	2.820	-28.4	SMITH	0.310	-	0.310	-	0.000	-	
				JOHANSEN	2.269	-23.5														
				DE VRIES	3.414	-65.2	SMITH	-	-											
				ADJ DE VRIES	-	-	GEMANT	-	-											
				RESISTER	-	-	VAN ROOYEN	-	-	GEOM. MEAN	-	-								
				KUNTZ-SMITH	-	-														
SAMPLE	25.92	1.592	1.490	KERSTEN	2.464	-6.4	MICKLEY	2.829	-28.4	MC GAV	2.820	-28.4	SMITH	0.310	-	0.310	-	0.000	-	
				JOHANSEN																

Table B17 (cont'd).

SAMPLE	1.172	4.428	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	2.571 2.571 3.486 3.486 ----- -----	5.0 2.4 43.5 42.4 ----- -----	MICKLEY MC GAW SMITH GERMANT VAN ROOYEN GEOM. MEAN	2.892 ----- ----- ----- ----- -----	19.1 ----- ----- ----- ----- -----	0.213 0.334 ----- ----- ----- -----
SAMPLE	1.174	4.424	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	2.522 2.522 3.479 3.479 ----- -----	20.4 12.4 42.4 42.4 ----- -----	MICKLEY MC GAW SMITH GERMANT VAN ROOYEN GEOM. MEAN	3.346 ----- ----- ----- ----- -----	14.4 ----- ----- ----- ----- -----	0.190 0.270 ----- ----- ----- -----
SAMPLE	1.174	4.443	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	2.499 2.499 3.407 3.407 ----- -----	8.7 2.1 43.2 43.2 ----- -----	MICKLEY MC GAW SMITH GERMANT VAN ROOYEN GEOM. MEAN	2.721 ----- ----- ----- ----- -----	9.4 ----- ----- ----- ----- -----	0.290 0.636 ----- ----- ----- -----
SAMPLE	1.176	4.461	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	2.459 2.459 3.443 3.443 ----- -----	20.4 16.4 26.4 26.4 ----- -----	MICKLEY MC GAW SMITH GERMANT VAN ROOYEN GEOM. MEAN	3.410 ----- ----- ----- ----- -----	12.4 ----- ----- ----- ----- -----	0.274 0.860 ----- ----- ----- -----
SAMPLE	1.176	5.174	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	2.798 2.798 3.237 3.237 ----- -----	19.7 12.7 17.8 17.8 ----- -----	MICKLEY MC GAW SMITH GERMANT VAN ROOYEN GEOM. MEAN	3.390 ----- ----- ----- ----- -----	6.8 ----- ----- ----- ----- -----	0.296 0.645 ----- ----- ----- -----
SAMPLE	1.178	5.153	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	2.470 2.470 3.460 3.460 ----- -----	20.2 18.2 20.1 20.1 ----- -----	MICKLEY MC GAW SMITH GERMANT VAN ROOYEN GEOM. MEAN	3.399 ----- ----- ----- ----- -----	17.0 ----- ----- ----- ----- -----	0.352 1.000 ----- ----- ----- -----

Table B18.

JUHANSEN SAND SA2 SATURATED FROZEN

TYPE OF SOIL: SAND NATURAL SATURATED FROZEN										
RHO =	2,600	TEMP =	-20,000	KU =	8,448	CL =	8,800	CSOLID =	+	3,900
KR =	+ 0.570	KICE =	-2,310	KU =	8,400	CL =	8,800	CSOLID =	+	3,848
SAMPLE MOISTURE	DRY	X	MEASURED	METHOD K	COMPUTED	DEVIATION	METHOD K	COMPUTED	DEVIATION	POR.
SAMPLE	18.00	+ 1.722	3.111	KERSTEN	3.351	7.7	MICKLEY	3.318	6.7	0.338
				JOHANSEN	3.351	6.7	MC GAW	-----	1.000	0.000
				DE VRIES	3.353	6.8	SMITH	-----	-----	-----
				ADJ DE VRIES	-----	-----	GERMANT	-----	-----	-----
				RESISTOR	3.393	9.8	VAN ROOYEN	-----	-----	-----
				KUNIT-SMITH	3.264	4.9	GEOM. MEAN	3.321	6.7	-----
RHO =	+ 4,900	TEMP =	-20,000	KU =	8,448	CL =	8,800	CSOLID =	+	4,145
KR =	+ 0.570	KICE =	-2,310	KU =	8,138	CL =	8,800	CSOLID =	+	3,848
SAMPLE MOISTURE	DRY	X	MEASURED	METHOD K	COMPUTED	DEVIATION	METHOD K	COMPUTED	DEVIATION	POR.
SAMPLE	18.00	+ 1.722	2.678	KERSTEN	3.351	25.7	MICKLEY	3.478	29.0	0.338
				JOHANSEN	3.351	26.0	MC GAW	-----	1.000	0.000
				DE VRIES	3.316	31.1	SMITH	-----	-----	-----
				ADJ DE VRIES	-----	-----	GERMANT	-----	-----	-----
				RESISTOR	3.548	32.5	VAN ROOYEN	-----	-----	-----
				KUNIT-SMITH	3.421	27.8	GEOM. MEAN	3.488	30.8	-----

Table B19.

JUHANSEN GRAVEL GR1 SATURATED FROZEN

TYPE OF SOIL: COARSE NATURAL SATURATED FROZEN										
RHO =	2,740	TEMP =	-7,000	KU =	8,448	CL =	8,800	CSOLID =	+	4,047
KR =	+ 0.570	KICE =	-2,310	KU =	8,400	CL =	8,800	CSOLID =	+	3,848
SAMPLE MOISTURE	DRY	X	MEASURED	METHOD K	COMPUTED	DEVIATION	METHOD K	COMPUTED	DEVIATION	POR.
SAMPLE	13.20	+ 1.967	3.103	KERSTEN	4.218	35.7	MICKLEY	3.448	31.1	0.263
				JOHANSEN	3.303	12.3	MC GAW	-----	1.000	0.000
				DE VRIES	3.264	12.3	SMITH	-----	-----	-----
				ADJ DE VRIES	-----	-----	GERMANT	-----	-----	-----
				RESISTOR	3.219	3.6	VAN ROOYEN	-----	-----	-----
				KUNIT-SMITH	3.245	11.0	GEOM. MEAN	3.453	11.3	-----
RHO =	+ 2,740	TEMP =	-20,000	KU =	8,448	CL =	8,800	CSOLID =	+	4,168
KR =	+ 0.570	KICE =	-2,310	KU =	8,048	CL =	8,800	CSOLID =	+	3,848
SAMPLE MOISTURE	DRY	X	MEASURED	METHOD K	COMPUTED	DEVIATION	METHOD K	COMPUTED	DEVIATION	POR.
SAMPLE	13.20	+ 1.967	3.584	KERSTEN	4.216	17.8	MICKLEY	3.594	6.3	0.263
				JOHANSEN	3.599	8.4	MC GAW	-----	1.000	0.000
				DE VRIES	3.628	1.2	SMITH	-----	-----	-----
				ADJ DE VRIES	-----	-----	GERMANT	-----	-----	-----
				RESISTOR	3.662	2.2	VAN ROOYEN	-----	-----	-----
				KUNIT-SMITH	3.587	8.1	GEOM. MEAN	3.599	8.4	-----

Table B20.

JOHANSEN GRAVEL GR3 SATURATED FROZEN

TYPE OF SOIL:		COARSE NATURAL SATURATED FROZEN									
RHO =	2,700 TEMP = -12,000 G =	0.470 CL =	0.000 CSOLIDS =	3,868 ALPHA = ----- KA = + 0.025							PUR.
KW = +	0.570 KICE = 2,400 KU =	0.570 KU =	2,000 CSOLID52 =	3,840							SAT.
SAMPLE	MOISTURE	DRY	K MEASURED	METHOD K COMPUTED	DEVIATION	(PERCENT)	METHOD K COMPUTED	DEVIATION	(PERCENT)	PUR.	NC
	CONTENT	DENSITY								SAT.	UNC
SAMPLE	14.38	+ 1.097	3.572	KERSTEN	3.989	-9.4	MICKLEY	3.394	-5.1	0.296	-----
				JOHANSEN	3.594	-5.8	MC GAN	-----	-----	1.000	0.000
				DE VRIES	3.422	-14.2	SMITH	-----	-----		
				ADJ DE VRIES	-----	-----	GEMANT	-----	-----		
				KESISTER	3.455	-3.3	VAN ROOYEN	-----	-----		
				KUNIT-SMITH	3.372	-9.6	GEOM. MEAN	3.394	-5.8		
RHO =	2,700 TEMP = -12,000 G =	0.470 CL =	0.000 CSOLIDS =	3,868 ALPHA = ----- KA = + 0.025							
KW = +	0.570 KICE = 2,400 KU =	0.570 KU =	2,000 CSOLID52 =	3,840							
SAMPLE	MOISTURE	DRY	K MEASURED	METHOD K COMPUTED	DEVIATION	(PERCENT)	METHOD K COMPUTED	DEVIATION	(PERCENT)	PUR.	NC
	CONTENT	DENSITY								SAT.	UNC
SAMPLE	14.38	+ 1.097	3.865	KERSTEN	3.989	-1.1	MICKLEY	3.529	-8.3	0.296	-----
				JOHANSEN	3.532	-8.6	MC GAN	-----	-----	1.000	0.000
				DE VRIES	3.559	-7.9	SMITH	-----	-----		
				ADJ DE VRIES	-----	-----	GEMANT	-----	-----		
				KESISTER	3.590	-9.1	VAN ROOYEN	-----	-----		
				KUNIT-SMITH	3.508	-9.2	GEOM. MEAN	3.532	-8.6		

Table B21.

JOHANSEN CRUSHED ROCK PUG SATURATED FROZEN

TYPE OF SOIL:		COARSE CRUSHED SATURATED FROZEN									
RHO =	2,600 TEMP = -24,000 G =	0.350 CL =	0.000 CSOLID52 =	3,222 ALPHA = ----- KA = + 0.025							PUR.
KW = +	0.570 KICE = 2,320 KU =	0.480 KU =	2,000 CSOLID52 =	3,840							SAT.
SAMPLE	MOISTURE	DRY	K MEASURED	METHOD K COMPUTED	DEVIATION	(PERCENT)	METHOD K COMPUTED	DEVIATION	(PERCENT)	PUR.	NC
	CONTENT	DENSITY								SAT.	UNC
SAMPLE	15.10	+ 1.061	2.772	KERSTEN	3.812	-37.5	MICKLEY	2.912	-5.1	0.386	-----
				JOHANSEN	2.913	-5.1	MC GAN	-----	-----	1.000	0.000
				DE VRIES	2.923	-5.3	SMITH	-----	-----		
				ADJ DE VRIES	-----	-----	GEMANT	-----	-----		
				KESISTER	2.930	-5.9	VAN ROOYEN	-----	-----		
				KUNIT-SMITH	2.886	-4.1	GEOM. MEAN	2.913	-5.1		
RHO =	2,600 TEMP = -24,000 G =	0.340 CL =	0.000 CSOLID52 =	3,381 ALPHA = ----- KA = + 0.025							
KW = +	0.570 KICE = 2,470 KU =	0.340 KU =	2,000 CSOLID52 =	3,840							
SAMPLE	MOISTURE	DRY	K MEASURED	METHOD K COMPUTED	DEVIATION	(PERCENT)	METHOD K COMPUTED	DEVIATION	(PERCENT)	PUR.	NC
	CONTENT	DENSITY								SAT.	UNC
SAMPLE	15.10	+ 1.061	2.888	KERSTEN	3.812	-32.4	MICKLEY	3.020	-4.6	0.386	-----
				JOHANSEN	3.626	-4.6	MC GAN	-----	-----	1.000	0.000
				DE VRIES	3.629	-4.9	SMITH	-----	-----		
				ADJ DE VRIES	-----	-----	GEMANT	-----	-----		
				KESISTER	3.639	-5.2	VAN ROOYEN	-----	-----		
				KUNIT-SMITH	3.594	-3.7	GEOM. MEAN	3.620	-4.6		

Table B22.

JOHANSEN CRUSHED ROCK PUG SATURATED FROZEN

TYPE OF SOIL:		COARSE CRUSHED SATURATED FROZEN									
RHO =	2,734 TEMP = -31,000 G =	0.350 CL =	0.000 CSOLID52 =	2,989 ALPHA = ----- KA = + 0.025							PUR.
KW = +	0.570 KICE = 2,320 KU =	0.520 KU =	2,000 CSOLID52 =	3,840							SAT.
SAMPLE	MOISTURE	DRY	K MEASURED	METHOD K COMPUTED	DEVIATION	(PERCENT)	METHOD K COMPUTED	DEVIATION	(PERCENT)	PUR.	NC
	CONTENT	DENSITY								SAT.	UNC
SAMPLE	23.06	+ 1.003	2.408	KERSTEN	3.658	-30.6	MICKLEY	2.181	-9.4	0.412	-----
				JOHANSEN	2.181	-9.4	MC GAN	-----	-----	1.000	0.000
				DE VRIES	2.182	-9.4	SMITH	-----	-----		
				ADJ DE VRIES	-----	-----	GEMANT	-----	-----		
				KESISTER	2.184	-9.3	VAN ROOYEN	-----	-----		
				KUNIT-SMITH	2.203	-8.3	GEOM. MEAN	2.181	-9.4		
RHO =	2,734 TEMP = -31,000 G =	0.350 CL =	0.000 CSOLID52 =	2,984 ALPHA = ----- KA = + 0.025							
KW = +	0.570 KICE = 2,470 KU =	0.350 KU =	2,000 CSOLID52 =	3,840							
SAMPLE	MOISTURE	DRY	K MEASURED	METHOD K COMPUTED	DEVIATION	(PERCENT)	METHOD K COMPUTED	DEVIATION	(PERCENT)	PUR.	NC
	CONTENT	DENSITY								SAT.	UNC
SAMPLE	23.06	+ 1.003	2.494	KERSTEN	3.658	-30.6	MICKLEY	2.249	-9.8	0.412	-----
				JOHANSEN	3.650	-30.8	MC GAN	-----	-----	1.000	0.000
				DE VRIES	3.652	-9.7	SMITH	-----	-----		
				ADJ DE VRIES	-----	-----	GEMANT	-----	-----		
				KESISTER	2.255	-9.6	VAN ROOYEN	-----	-----		
				KUNIT-SMITH	2.287	-8.3	GEOM. MEAN	2.249	-9.8		

Table B23.

WILHELM ERNST & HENRICH FISCHER

TYPE II SULFIDE VENUS MATERIAL IN SATELLITE E FROZEN										0.024
RHO *	Z *	T.D. *	-S.C.	W.E. *	G.P.C. *	1,000 GEMMERS *	1,000 ALPHA E GEMMERS *	1,000	ALPHA E GEMMERS *	0.024
V.D. *	SAT. VITRES *	Z, Z-200	V.E. *	X, X-200	K.E. *	Z, Z-200	GEMMERS *	KA *	KA-200	
SAMPLE	1, 2	1.7 A	0.745	0.740	KERSTEN	1.207	173.7	MICKEY	1.224	
					JOHANSEN	1.528	115.5	MC GOW		
					DE VRIES	2.030	11.0	SUTTU		
					DU DE VRIES	-----	-----	GERHART		
					DESTYSON	-----	-----	VAN ROOYEN		
					KUNITSCHMIT	-----	-----	GEORGE	MEAL	
SAMPLE	1, 2	1.7 A	0.745	0.740	KERSTEN	1.207	173.7	MICKEY	1.546	-13 P
					JOHANSEN	1.528	115.5	MC GOW		0.368
					DE VRIES	2.030	11.0	SUTTU		0.513
					DU DE VRIES	-----	-----	GERHART		-0.800
					DESTYSON	-----	-----	VAN ROOYEN		
					KUNITSCHMIT	-----	-----	GEORGE	MEAL	
SAMPLE	1, 2	1.7 A	0.745	0.740	KERSTEN	1.207	173.7	MICKEY	1.526	-15 L
					JOHANSEN	1.528	115.5	MC GOW		0.368
					DE VRIES	2.030	11.0	SUTTU		0.480
					DU DE VRIES	-----	-----	GERHART		
					DESTYSON	-----	-----	VAN ROOYEN		
					KUNITSCHMIT	-----	-----	GEORGE	MEAL	
SAMPLE	1, 2	1.7 A	0.745	0.740	KERSTEN	1.206	174.0	MICKEY	2.526	-15.2
					JOHANSEN	2.560	116.4	MC GOW		0.330
					DE VRIES	2.451	11.4	SUTTU		0.027
					DU DE VRIES	-----	-----	GERHART		
					DESTYSON	-----	-----	VAN ROOYEN		
					KUNITSCHMIT	-----	-----	GEORGE	MEAL	
SAMPLE	1, 2	1.7 A	0.745	0.740	KERSTEN	1.240	178.0	MICKEY	2.249	-17 C
					JOHANSEN	2.172	120.8	MC GOW		0.330
					DE VRIES	2.216	11.2	SUTTU		0.000
					DU DE VRIES	-----	-----	GERHART		-0.700
					DESTYSON	-----	-----	VAN ROOYEN		
					KUNITSCHMIT	-----	-----	GEORGE	MEAL	

Table B24.

POLY(4-ETYL-1,4-PHENYLENE SULFONIC ACID) MONOSATURATED FROG ERYTHROCYTE MEMBRANE

Table B25.

PENNGR ET AL. SOIL & UNSATURATED FROZEN

TYPE OF SITE: CLAY NATURAL UNSATURATED FROZEN				T HOM ALPHA = ----- VA = 0.024	
RHO = 1.70	TEMP = -5 COLD 4 E + 0.000 CL = + 1,000 GENTIDE = 2,000	KU = 0.560 KICE = 2,290 KU = + 0.470 VA = + 2,000 GENTIDE = 2,000			
SAMPLE	MOISTURE	DAY	% MEASURED	METHOD	COMPUTED
SAMPLE	CO TENSILE	CO TENSILE	DEVIATION	METHOD	COMPUTED
	1.442	1.442	(PERCENT)	HICKLEY	1.281
	0.840	0.840		MC GAN	1.282
				SMITH	1.282
				GERANT	1.282
				VAN ROOYEN	1.282
				GEMM HEAL	1.282
SAMPLE	6.50	1.642	0.840	KERSTEN	1.285
				JOHANSEN	1.270
				DE VRIES	1.267
				ADJ DE VRIES	1.267
				RESISTOR	1.267
				KUNIT-SMITH	1.267
SAMPLE	11.20	1.767	1.530	KERSTEN	1.275
				JOHANSEN	1.270
				DE VRIES	1.267
				ADJ DE VRIES	1.267
				RESISTOR	1.267
				KUNIT-SMITH	1.267
SAMPLE	11.20	1.768	1.530	KERSTEN	1.275
				JOHANSEN	1.272
				DE VRIES	1.269
				ADJ DE VRIES	1.269
				RESISTOR	1.269
				KUNIT-SMITH	1.269
SAMPLE	14.46	1.826	1.248	KERSTEN	1.285
				JOHANSEN	1.277
				DE VRIES	1.274
				ADJ DE VRIES	1.274
				RESISTOR	1.274
				KUNIT-SMITH	1.274
SAMPLE	14.46	1.826	1.248	KERSTEN	1.285
				JOHANSEN	1.277
				DE VRIES	1.274
				ADJ DE VRIES	1.274
				RESISTOR	1.274
				KUNIT-SMITH	1.274

Table B26.

PENNGR ET AL. SITE 7 UNSATURATED FROZEN

TYPE OF SITE: CLAY NATURAL UNSATURATED FROZEN				T HOM ALPHA = ----- VA = 0.024	
RHO = 1.70	TEMP = -5 COLD 4 E + 0.000 CL = + 1,000 GENTIDE = 2,000	KU = 0.560 KICE = 2,290 KU = + 0.470 VA = + 2,000 GENTIDE = 2,000			
SAMPLE	MOISTURE	DAY	% MEASURED	METHOD	COMPUTED
SAMPLE	CO TENSILE	CO TENSILE	DEVIATION	METHOD	COMPUTED
	1.073	1.148	(PERCENT)	HICKLEY	1.420
	1.073	1.148		MC GAN	1.420
				SMITH	1.420
				GERANT	1.420
				VAN ROOYEN	1.420
				GEMM HEAL	1.420
SAMPLE	4.45	1.973	1.168	KERSTEN	1.240
				JOHANSEN	1.187
				DE VRIES	1.200
				ADJ DE VRIES	1.200
				RESISTOR	1.200
				KUNIT-SMITH	1.200
SAMPLE	8.52	2.074	1.877	KERSTEN	1.241
				JOHANSEN	1.247
				DE VRIES	1.272
				ADJ DE VRIES	1.272
				RESISTOR	1.272
				KUNIT-SMITH	1.272
SAMPLE	8.52	2.074	1.877	KERSTEN	1.241
				JOHANSEN	1.247
				DE VRIES	1.272
				ADJ DE VRIES	1.272
				RESISTOR	1.272
				KUNIT-SMITH	1.272
SAMPLE	13.50	2.040	1.581	KERSTEN	1.241
				JOHANSEN	1.240
				DE VRIES	1.242
				ADJ DE VRIES	1.242
				RESISTOR	1.242
				KUNIT-SMITH	1.242
SAMPLE	13.50	2.058	1.581	KERSTEN	1.251
				JOHANSEN	2.110
				DE VRIES	2.447
				ADJ DE VRIES	2.447
				RESISTOR	2.447
				KUNIT-SMITH	2.447

Table B27.

PENICILLIUM SP. IN SATURATED FROZEN									
BUOY	TYPE	TEST	1000 CL	1000 FCO	1000 FCO TBS	1000 ALPHA	KA	0.024	
BUOY 1	TEST 1	1,517	1,517	1,517	1,517	1,517	1,517	1,517	
SAMPLE	TEST	TYPE	MEASURED	METHOD	COMPUTED	DEVIATION	METHOD	COMPUTED	DEVIATION
SAMPLE 1	TEST 1	1,517	1,517	KERSTEN	1,454	25.8	MICKEY	0.947	147.0
			JOHANSEN	1,410	48.0	MC GAI			
			DE VRIES	1,058	102.4	SMITH			
			NJ DE VRIES	1,058	102.4	GERMANT			
			RESISTOR	1,058	102.4	VAN ROOYEN			
			KUNFT-SMITH	1,058	102.4	GEOR. MEAN			
SAMPLE 2	TEST 1	1,517	1,517	KERSTEN	1,454	25.8	MICKEY	0.958	145.5
			JOHANSEN	1,459	56.5	MC GAI			
			DE VRIES	1,098	95.7	SMITH			
			NJ DE VRIES	1,098	95.7	GERMANT			
			RESISTOR	1,098	95.7	VAN ROOYEN			
			KUNFT-SMITH	1,098	95.7	GEOR. MEAN			
SAMPLE 3	TEST 1	1,517	1,517	KERSTEN	1,493	36.4	MICKEY	1.573	45.0
			JOHANSEN	1,497	28.4	MC GAI			
			DE VRIES	1,057	117.0	SMITH			
			NJ DE VRIES	1,057	117.0	GERMANT			
			RESISTOR	1,057	117.0	VAN ROOYEN			
			KUNFT-SMITH	1,057	117.0	GEOR. MEAN			
SAMPLE 4	TEST 1	1,517	1,517	KERSTEN	1,493	36.4	MICKEY	1.546	42.0
			JOHANSEN	1,497	51.4	MC GAI			
			DE VRIES	1,094	57.4	SMITH			
			NJ DE VRIES	1,094	57.4	GERMANT			
			RESISTOR	1,094	57.4	VAN ROOYEN			
			KUNFT-SMITH	1,094	57.4	GEOR. MEAN			
SAMPLE 5	TEST 1	1,518	1,518	KERSTEN	1,425	-2.9	MICKEY	1.842	12.1
			JOHANSEN	1,435	28.4	MC GAI			
			DE VRIES	1,043	41.4	SMITH			
			NJ DE VRIES	1,043	41.4	GERMANT			
			RESISTOR	1,043	41.4	VAN ROOYEN			
			KUNFT-SMITH	1,043	41.4	GEOR. MEAN			
SAMPLE 6	TEST 1	1,518	1,518	KERSTEN	1,425	-2.9	MICKEY	1.822	0.9
			JOHANSEN	1,414	9.4	MC GAI			
			DE VRIES	1,020	10.2	SMITH			
			NJ DE VRIES	1,020	10.2	GERMANT			
			RESISTOR	1,020	10.2	VAN ROOYEN			
			KUNFT-SMITH	1,020	10.2	GEOR. MEAN			

Table B28.

PENICILLIUM SP. IN UNSATURATED FROZEN									
BUOY	TYPE	TEST	1000 CL	1000 FCO	1000 FCO TBS	1000 ALPHA	KA	0.024	
BUOY 1	TEST 1	1,518	1,518	1,518	1,518	1,518	1,518	1,518	
SAMPLE	TEST	TYPE	MEASURED	METHOD	COMPUTED	DEVIATION	METHOD	COMPUTED	DEVIATION
SAMPLE 1	TEST 1	1,518	1,518	KERSTEN	1,433	4.2	MICKEY	1.013	102.2
			JOHANSEN	1,206	78.2	MC GAI			
			DE VRIES	1,157	126.4	SMITH			
			NJ DE VRIES	1,157	126.4	GERMANT			
			RESISTOR	1,157	126.4	VAN ROOYEN			
			KUNFT-SMITH	1,157	126.4	GEOR. MEAN			
SAMPLE 2	TEST 1	1,518	1,518	KERSTEN	1,433	4.2	MICKEY	1.026	100.3
			JOHANSEN	1,447	26.4	MC GAI			
			DE VRIES	1,054	28.4	SMITH			
			NJ DE VRIES	1,054	28.4	GERMANT			
			RESISTOR	1,054	28.4	VAN ROOYEN			
			KUNFT-SMITH	1,054	28.4	GEOR. MEAN			
SAMPLE 3	TEST 1	1,518	1,518	KERSTEN	1,405	3.4	MICKEY	1.740	74.3
			JOHANSEN	1,208	32.4	MC GAI			
			DE VRIES	1,241	22.4	SMITH			
			NJ DE VRIES	1,241	22.4	GERMANT			
			RESISTOR	1,241	22.4	VAN ROOYEN			
			KUNFT-SMITH	1,241	22.4	GEOR. MEAN			
SAMPLE 4	TEST 1	1,518	1,518	KERSTEN	1,405	3.4	MICKEY	1.717	74.1
			JOHANSEN	1,400	12.4	MC GAI			
			DE VRIES	1,181	20.2	SMITH			
			NJ DE VRIES	1,181	20.2	GERMANT			
			RESISTOR	1,181	20.2	VAN ROOYEN			
			KUNFT-SMITH	1,181	20.2	GEOR. MEAN			
SAMPLE 5	TEST 1	1,518	1,518	KERSTEN	1,405	3.4	MICKEY	1.717	74.1
			JOHANSEN	1,400	12.4	MC GAI			
			DE VRIES	1,242	22.4	SMITH			
			NJ DE VRIES	1,242	22.4	GERMANT			
			RESISTOR	1,242	22.4	VAN ROOYEN			
			KUNFT-SMITH	1,242	22.4	GEOR. MEAN			
SAMPLE 6	TEST 1	1,518	1,518	KERSTEN	1,408	6.1	MICKEY	1.057	25.0
			JOHANSEN	2,224	49.4	MC GAI			
			DE VRIES	2,422	42.4	SMITH			
			NJ DE VRIES	2,422	42.4	GERMANT			
			RESISTOR	2,422	42.4	VAN ROOYEN			
			KUNFT-SMITH	2,422	42.4	GEOR. MEAN			
SAMPLE 7	TEST 1	1,518	1,518	KERSTEN	1,408	6.1	MICKEY	1.049	20.3
			JOHANSEN	1,802	18.4	MC GAI			
			DE VRIES	1,021	20.4	SMITH			
			NJ DE VRIES	1,021	20.4	GERMANT			
			RESISTOR	1,021	20.4	VAN ROOYEN			
			KUNFT-SMITH	1,021	20.4	GEOR. MEAN			

Table B29.

PARKER ET AL. 2001 7. UNSATURATED FATTY ACIDS

Table B30.

KERSTEN FAIRBANKS SILTY CLAY LOAM PAZID UNSATURATED EROSION

TYPE OF SOIL: CLAY NATURAL UNSATURATED FROZEN
RHO = 2.710 TEMP = -4.000 u = 0.541 CL = 0.278 CSOLVIS = 7 0.976 ALPHAI = ----- KA = 0.04

SAMPLE	MOISTURE CONTENT	DRY DENSITY	X MEASURED	METHOD K COMPUTED	DEVIATION (PERCENT)	METHOD K COMPUTED	DEVIATION (PERCENT)	PUR. SAT.	NE UNC
SAMPLE	2.40	0.924	0.140	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KERSTEN KUNIJ-SMITH	0.112 0.108 0.140 0.140 0.112 0.112	-15.1 -18.8 +24.8 -24.8 -24.8 -24.8	MICKEY MC GEE SMITH GEMANT VAN RUYEN GEOM. MEAN	0.858 0.858 0.858 0.858 0.887 0.887	513.6 513.6 513.6 513.6 -78.1 -78.1
SAMPLE	2.40	0.924	0.140	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KERSTEN KUNIJ-SMITH	0.112 0.108 0.140 0.140 0.112 0.112	-15.1 -18.8 +24.8 -24.8 -24.8 -24.8	MICKEY MC GEE SMITH GEMANT VAN RUYEN GEOM. MEAN	0.855 0.855 0.855 0.855 0.884 0.884	511.2 511.2 511.2 511.2 -39.9 -39.9
SAMPLE	2.50	1.118	0.189	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KERSTEN KUNIJ-SMITH	0.160 0.170 0.177 0.177 0.160 0.160	-15.1 -18.8 +25.3 -25.3 -15.1 -15.1	MICKEY MC GEE SMITH GEMANT VAN RUYEN GEOM. MEAN	1.869 1.869 1.869 1.869 0.130 0.130	471.4 471.4 471.4 471.4 -31.4 -31.4
SAMPLE	2.50	1.118	0.189	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KERSTEN KUNIJ-SMITH	0.160 0.170 0.177 0.177 0.160 0.160	-15.1 -18.8 +25.3 -25.3 -15.1 -15.1	MICKEY MC GEE SMITH GEMANT VAN RUYEN GEOM. MEAN	1.876 1.876 1.876 1.876 0.126 0.126	487.4 487.4 487.4 487.4 -33.5 -33.5
SAMPLE	2.40	1.278	0.235	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES KERSTEN KUNIJ-SMITH	0.216 0.200 0.200 0.200 0.216 0.216	-16.7 -22.9 +26.1 -26.1 -16.7 -16.7	MICKEY MC GEE SMITH GEMANT VAN RUYEN GEOM. MEAN	1.270 1.270 1.270 1.270 0.160 0.160	443.8 443.8 443.8 443.8 -73.4 -73.4

Table B30 (cont'd).

SAMPLE	2.48	1.278	0.235	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	0.435 0.604 1.014 1.017 1.044 1.044	-12.7 2.4 0.6 213.9 224.9 224.9	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	1.270 1.375 1.375 0.388 0.415 0.294	446.5 -25.0 -26.7 -26.0 -24.5 -30.0	0.528 2.058 + 0.031
SAMPLE	0.00	1.272	0.420	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	0.436 0.604 1.014 1.017 1.044 1.044	-12.7 2.4 0.6 213.9 224.9 224.9	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	1.392 1.375 1.375 0.388 0.415 0.294	211.7 -25.0 -26.7 -26.0 -24.5 -30.0	0.531 2.172 - 0.092
SAMPLE	0.00	1.272	0.420	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	0.436 0.604 1.014 1.017 1.044 1.044	-12.7 2.4 0.6 213.9 224.9 224.9	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	1.390 1.375 1.375 0.388 0.415 0.294	228.5 -25.0 -26.7 -26.0 -24.5 -30.0	0.531 2.102 - 0.060
SAMPLE	0.00	1.422	0.522	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	0.564 0.826 0.948 0.948 1.044 1.044	-8.6 7.6 224.9 224.9 224.9 224.9	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	1.437 1.375 1.375 0.415 0.415 0.294	213.0 -25.0 -26.7 -26.0 -24.5 -30.0	0.475 2.225 - 0.092
SAMPLE	6.98	1.422	0.522	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	0.564 0.826 0.948 0.948 1.044 1.044	-8.6 7.6 224.9 224.9 224.9 224.9	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	1.624 1.375 1.375 0.415 0.415 0.294	211.0 -25.0 -26.7 -26.0 -24.5 -30.0	0.475 2.214 - 0.060
SAMPLE	12.38	1.283	0.632	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	0.743 1.187 1.083 1.083 1.083 1.083	17.6 82.9 185.5 185.5 185.5 185.5	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	1.595 1.375 1.375 0.539 0.539 0.294	152.5 -25.0 -26.7 -26.0 -24.5 -30.0	0.527 0.327 - 0.090
SAMPLE	12.38	1.283	0.632	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	0.743 1.187 1.083 1.083 1.083 1.083	17.6 82.9 185.5 185.5 185.5 185.5	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	1.581 1.375 1.375 0.516 0.516 0.294	150.4 -25.0 -26.7 -26.0 -24.5 -30.0	0.527 0.316 - 0.060
SAMPLE	12.38	1.443	0.870	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	0.725 1.134 1.260 1.260 1.044 1.044	-8.3 76.9 159.9 159.9 159.9 159.9	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	1.899 1.375 1.375 0.658 0.658 0.294	116.4 -25.0 -26.7 -26.0 -24.5 -30.0	0.467 0.311 - 0.090
SAMPLE	12.38	1.443	0.870	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	0.725 1.134 1.260 1.260 1.044 1.044	-8.3 76.9 159.9 159.9 159.9 159.9	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	1.884 1.375 1.375 0.644 0.644 0.294	116.0 -25.0 -26.7 -26.0 -24.5 -30.0	0.467 0.310 - 0.060
SAMPLE	12.40	1.616	1.201	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	1.414 1.476 1.289 1.289 1.044 1.044	1.1 73.6 135.5 135.5 135.5 135.5	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	2.308 1.375 1.375 0.794 0.794 0.294	92.2 -25.0 -26.7 -26.0 -24.5 -30.0	0.464 0.341 - 0.060
SAMPLE	12.40	1.616	1.201	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	1.214 1.377 1.171 1.171 1.044 1.044	1.1 55.3 42.4 42.4 135.5 135.5	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	2.288 1.375 1.375 0.780 0.780 0.294	90.5 -25.0 -26.7 -26.0 -24.5 -30.0	0.464 0.328 - 0.060
SAMPLE	17.68	1.286	0.933	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	1.031 1.037 2.185 2.185 1.044 1.044	19.5 55.5 134.2 134.2 135.5 135.5	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	1.795 1.375 1.375 0.693 0.693 0.294	92.4 -25.0 -26.7 -26.0 -24.5 -30.0	0.525 0.470 - 0.060
SAMP.:	17.68	1.286	0.933	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	1.031 1.037 2.185 2.185 1.044 1.044	19.5 55.5 134.2 134.2 135.5 135.5	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	1.789 1.375 1.375 0.684 0.684 0.294	90.8 -25.0 -26.7 -26.0 -24.5 -30.0	0.525 0.459 - 0.060
SAMPLE	17.68	1.438	1.270	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	1.277 1.287 1.197 1.197 1.044 1.044	8.1 67.8 119.0 119.0 135.5 135.5	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	2.170 1.375 1.375 0.851 0.851 0.294	70.0 -25.0 -26.7 -26.0 -24.5 -30.0	0.499 0.395 - 0.060
SAMPLE	17.68	1.438	1.270	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	1.277 1.287 1.197 1.197 1.044 1.044	8.1 67.8 119.0 119.0 135.5 135.5	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	2.151 1.375 1.375 0.841 0.841 0.294	68.5 -25.0 -26.7 -26.0 -24.5 -30.0	0.499 0.383 - 0.060
SAMPLE	17.68	1.634	1.090	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	1.081 1.084 3.354 3.354 1.044 1.044	-6.9 74.8 97.8 97.8 135.5 135.5	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	2.799 1.375 1.375 1.025 1.025 0.294	65.1 -25.0 -26.7 -26.0 -24.5 -30.0	0.397 0.398 - 0.060
SAMPLE	17.68	1.634	1.090	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIT-SMITH	1.081 1.084 3.354 3.354 1.044 1.044	-6.9 74.8 97.8 97.8 135.5 135.5	MICKLEY MC GARN SMITH GEMANT VAN RUYEN GEOM. MEAN	2.768 1.375 1.375 1.013 1.013 0.294	63.3 -25.0 -26.7 -26.0 -24.5 -30.0	0.397 0.384 - 0.060

Table B30 (cont'd).

SAMPLE	25.00	1.277	1.281	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	1.412 2.612 17.6 17.5 17.5 17.5	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	2.092 ----- ----- ----- 0.793 -----	74.2 ----- ----- ----- 34.0 -----	0.529 0.658 -----
SAMPLE	25.00	1.277	1.281	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	1.412 2.612 17.6 17.5 17.5 17.5	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	2.074 ----- ----- ----- 0.790 -----	72.6 ----- ----- ----- 34.2 -----	0.529 0.648 -----
SAMPLE	25.30	1.435	1.590	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	1.756 2.693 19.7 19.6 19.6 19.6	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	2.672 ----- ----- ----- 0.987 -----	67.4 ----- ----- ----- 38.2 -----	0.470 0.841 -----
SAMPLE	25.30	1.435	1.590	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	1.756 2.693 19.7 19.6 19.6 19.6	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	2.642 ----- ----- ----- 0.983 -----	66.5 ----- ----- ----- 38.4 -----	0.470 0.838 -----
SAMPLE	24.70	1.517	1.849	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	1.987 2.689 17.1 17.1 17.1 17.1	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	3.037 ----- ----- ----- 1.099 -----	64.3 ----- ----- ----- 48.5 -----	0.448 0.928 -----
SAMPLE	24.70	1.517	1.849	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	1.987 2.689 17.1 17.1 17.1 17.1	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	2.991 ----- ----- ----- 1.094 -----	61.6 ----- ----- ----- 48.8 -----	0.448 0.915 -----
SAMPLE	24.70	1.444	1.908	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	2.053 2.650 17.6 17.6 17.6 17.6	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	3.462 ----- ----- ----- 1.035 -----	81.5 ----- ----- ----- 45.7 -----	0.467 1.006 -----
SAMPLE	24.70	1.446	1.908	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	2.051 2.650 17.6 17.6 17.6 17.6	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	3.340 ----- ----- ----- 1.037 -----	75.1 ----- ----- ----- 45.6 -----	0.466 0.993 -----
SAMPLE	37.30	1.273	1.948	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	2.686 2.617 17.0 17.0 17.0 17.0	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	3.038 ----- ----- ----- 0.831 -----	55.5 ----- ----- ----- 57.3 -----	0.538 0.977 -----
SAMPLE	37.30	1.273	1.948	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	2.686 2.626 17.0 17.0 17.0 17.0	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	2.965 ----- ----- ----- 0.831 -----	52.2 ----- ----- ----- 57.4 -----	0.538 0.966 -----

Table B31.

KERSTEN ALARTHUSY SILT LOAM UNSATURATED FROZEN

TYPE OF SITE	CLAY	NATURAL	UNSATURATED	FROZEN	RHO = 1.717	TDP = 76.00	CL = 0.015	CSOLIDS = 0.136	CSOLIDS = 0.136	ALPHA = 2.043	KA = 0.046	
SAMPLE	POLE SITE	DRY	X	MEASURED	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	0.163	0.163	0.136	0.136	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	0.526 ----- ----- ----- ----- -----	0.537 0.042 0.000
SAMPLE	1.7	1.346	0.173	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	0.199 2.693 17.6 17.6 17.6 17.6	13.8 14.0 14.0 14.0 14.0 14.0	0.199 2.693 17.6 17.6 17.6 17.6	0.199 2.693 17.6 17.6 17.6 17.6	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	0.402 ----- ----- ----- ----- -----	0.392 0.000	
SAMPLE	1.4	1.446	0.174	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	0.230 2.690 17.6 17.6 17.6 17.6	23.7 14.0 14.0 14.0 14.0 14.0	0.230 2.690 17.6 17.6 17.6 17.6	0.230 2.690 17.6 17.6 17.6 17.6	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	0.652 ----- ----- ----- ----- -----	0.666 0.048 0.000	
SAMPLE	1.7	1.552	0.235	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	0.319 2.655 17.0 17.0 17.0 17.0	35.8 41.0 41.0 41.0 41.0 41.0	0.319 2.655 17.0 17.0 17.0 17.0	0.319 2.655 17.0 17.0 17.0 17.0	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	0.726 ----- ----- ----- ----- -----	0.625 0.068 0.000	
SAMPLE	1.7	1.552	0.235	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNII-SMITH	0.319 2.655 17.0 17.0 17.0 17.0	35.8 41.0 41.0 41.0 41.0 41.0	0.319 2.655 17.0 17.0 17.0 17.0	0.319 2.655 17.0 17.0 17.0 17.0	MICKLEY MC GAN SMITH GEMANT VAN ROOYEN GEOM. MEAN	0.726 ----- ----- ----- ----- -----	0.625 0.068 0.000	

Table B31 (cont'd).

SAMPLE	1,401	1,460	0,526	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	0.485 0.735 1.102 ----- ----- -----	40.1 49.4 109.4 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	0.881 0.881 0.881 ----- ----- -----	67.3 ----- ----- ----- ----- -----	0.423 0.265 0.000
SAMPLE	1,411	1,470	0,537	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.176 1.458 1.486 ----- ----- -----	23.3 33.7 48.6 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.098 1.098 1.098 ----- ----- -----	17.1 ----- ----- ----- ----- -----	0.470 0.344 0.000
S.	1,421	1,481	1,087	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.171 1.456 1.460 ----- ----- -----	26.1 33.7 53.2 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.256 1.256 1.256 ----- ----- -----	15.5 ----- ----- ----- ----- -----	0.429 0.342 0.000
SAMPLE	1,431	1,473	1,526	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.895 1.403 1.097 ----- ----- -----	26.2 26.7 40.0 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.702 1.702 1.702 ----- ----- -----	11.6 ----- ----- ----- ----- -----	0.383 0.381 0.000
SAMPLE	1,431	1,473	1,288	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.435 1.456 1.423 ----- ----- -----	27.1 46.1 43.1 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.645 1.645 1.645 ----- ----- -----	12.2 ----- ----- ----- ----- -----	0.462 0.385 0.000
SAMPLE	1,431	1,473	1,550	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.469 1.490 1.424 ----- ----- -----	21.6 28.1 21.5 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.730 1.730 1.730 ----- ----- -----	12.5 ----- ----- ----- ----- -----	0.419 0.311 0.000

Table B32.

KERSTEN RANDSEY 54° BY LOAM UNSATURATED FROZEN											
TYPE	TEST	CLAY	NATURAL	UNSATURATED	FROZEN	DENSITY	MEASURED	METHOD	COMPUTED	DEVIATION	TYPE
RHO	TEST	CL	%	CL	%	RHO	RHO	RHO	RHO	(PERCENT)	RHO
SAMPLE	1,401	1,460	0,526	0,513	CL %	0.485	0,526	4,185 CSOLIDS %	1,245	4,185 ALPHA %	KA %
				2,350	KM %	0.735	0,526	2,000 CSOLIDSP %	0.485	5,227	0.046
SAMPLE	1,401	1,460	0,526	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	0.482 0.493 1.102 ----- ----- -----	45.3 45.3 109.4 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.245 1.245 1.245 ----- ----- -----	10.0 10.0 10.0 ----- ----- -----	0.496 0.093 0.000	
SAMPLE	1,411	1,470	0,537	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.176 1.458 1.486 ----- ----- -----	23.3 33.7 48.6 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.098 1.098 1.098 ----- ----- -----	17.1 17.1 17.1 ----- ----- -----	0.470 0.344 0.000	
SAMPLE	1,421	1,481	1,087	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.171 1.456 1.460 ----- ----- -----	26.1 33.7 53.2 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.256 1.256 1.256 ----- ----- -----	15.5 15.5 15.5 ----- ----- -----	0.429 0.342 0.000	
SAMPLE	1,431	1,473	1,526	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.895 1.403 1.097 ----- ----- -----	26.2 26.7 40.0 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.702 1.702 1.702 ----- ----- -----	11.6 11.6 11.6 ----- ----- -----	0.383 0.381 0.000	
SAMPLE	1,431	1,473	1,288	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.435 1.456 1.423 ----- ----- -----	27.1 46.1 43.1 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.645 1.645 1.645 ----- ----- -----	12.2 12.2 12.2 ----- ----- -----	0.462 0.385 0.000	
SAMPLE	1,431	1,473	1,550	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.469 1.490 1.424 ----- ----- -----	21.6 28.1 21.5 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.730 1.730 1.730 ----- ----- -----	12.5 12.5 12.5 ----- ----- -----	0.419 0.311 0.000	
SAMPLE	1,441	1,481	1,507	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.743 1.493 1.414 ----- ----- -----	40.8 45.6 118.7 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.028 1.028 1.028 ----- ----- -----	69.5 69.5 69.5 ----- ----- -----	0.346 0.343 0.000	
SAMPLE	1,451	1,486	1,507	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.743 1.493 2.814 ----- ----- -----	40.8 45.6 46.7 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	2.292 2.292 2.292 ----- ----- -----	58.1 58.1 58.1 ----- ----- -----	0.289 0.489 0.000	
SAMPLE	1,451	1,486	1,507	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.743 1.493 2.814 ----- ----- -----	40.8 45.6 46.7 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	2.292 2.292 2.292 ----- ----- -----	58.1 58.1 58.1 ----- ----- -----	0.289 0.489 0.000	
SAMPLE	1,461	1,494	1,480	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	0.927 1.427 2.187 ----- ----- -----	6.2 44.6 121.1 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	1.803 1.803 1.803 ----- ----- -----	82.3 82.3 82.3 ----- ----- -----	0.403 0.394 0.000	
SAMPLE	1,471	1,504	1,260	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.744 1.492 2.668 ----- ----- -----	0.3 54.2 115.2 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	2.156 2.156 2.156 ----- ----- -----	75.7 75.7 75.7 ----- ----- -----	0.346 0.326 0.000	
SAMPLE	1,481	1,514	1,507	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.744 1.492 3.123 ----- ----- -----	0.3 54.2 107.3 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	2.574 2.574 2.574 ----- ----- -----	70.8 70.8 70.8 ----- ----- -----	0.286 0.286 0.000	
SAMPLE	1,491	1,520	1,517	KERSTEN JOHANSEN DE VRIES ADJ DE VRIES RESISTOR KUNIIT-SMITH	1.756 1.493 3.451 ----- ----- -----	48.2 24.6 57.3 ----- ----- -----	MICKLEY MC GAV SMITH GERMANT VAN ROOYEN GEOM. MEAN	3.019 3.019 3.019 ----- ----- -----	19.8 19.8 19.8 ----- ----- -----	0.246 0.258 0.000	

Table B32 (cont'd).

SAMPLE	13.70	1.409	1.370	KERSTEN	1.294	+0.1	MICKLEY	2.093	51.4	0.601	-----
				JOHANSEN	2.047	-0.6	MC GAI			0.504	-----
				DE VRIES	2.616	-0.6	SCHMIT			0.504	-----
				ADJ	DE VRIES	-----	GEMANT			0.504	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.504	-----
				KUNIITS-SMITH	-----	-----				0.504	-----
SAMPLE	13.70	1.270	1.803	KERSTEN	1.618	+0.3	MICKLEY	2.512	50.3	0.340	-----
				JOHANSEN	2.506	-0.3	MC GAI			0.340	-----
				DE VRIES	3.045	-0.3	SCHMIT			0.340	-----
				ADJ	DE VRIES	-----	GEMANT			0.340	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.340	-----
				KUNIITS-SMITH	-----	-----				0.340	-----
SAMPLE	12.70	1.940	2.493	KERSTEN	2.115	+0.3	MICKLEY	3.285	55.7	0.276	-----
				JOHANSEN	3.220	-0.3	MC GAI			0.276	-----
				DE VRIES	3.515	-0.3	SCHMIT			0.276	-----
				ADJ	DE VRIES	-----	GEMANT			0.276	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.276	-----
				KUNIITS-SMITH	-----	-----				0.276	-----
SAMPLE	17.80	1.580	1.860	KERSTEN	1.576	+0.7	MICKLEY	2.352	25.0	0.167	-----
				JOHANSEN	2.548	-0.6	MC GAI			0.167	-----
				DE VRIES	2.876	-0.6	SCHMIT			0.167	-----
				ADJ	DE VRIES	-----	GEMANT			0.167	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.167	-----
				KUNIITS-SMITH	-----	-----				0.167	-----
SAMPLE	18.10	1.753	2.593	KERSTEN	2.036	+1.0	MICKLEY	3.870	40.0	0.360	-----
				JOHANSEN	3.272	-0.1	MC GAI			0.360	-----
				DE VRIES	3.411	-0.2	SCHMIT			0.360	-----
				ADJ	DE VRIES	-----	GEMANT			0.360	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.360	-----
				KUNIITS-SMITH	-----	-----				0.360	-----

Table B33.

USSR BUILDING CODE (1960) CLAY SOILS UNSATURATED FROZEN											
TYPE OF SOIL			PLAY NATURAL /SATURATED FROZEN			K=0.024			K=0.024		
R ₁	R ₂	R ₃	R ₄	R ₅	R ₆	R ₇	R ₈	R ₉	R ₁₀	R ₁₁	R ₁₂
SAMPLE	MOISTURE	DRY	K	MEASUREMENT	K	COMPUTED	DEVIATION	(PERCENT)	METHOD	K COMPUTED	DEVIATION (PERCENT)
SAMPLE	CONTENT	DENSITY	DE	DE	DE	DE	DE	DE	WICKLEY	0.674	PUR. SAT.
	%, DU	g/cm ³	ADJ	JOHANSEN	1.544	0.507	+0.6	+0.7	MC GAI	0.674	0.177
				DE VRIES	1.597	-0.551	-1.9	-2.0	SCHMIT		0.000
				RESISTOR	-----	-----	-----	-----	GEMANT		
				KUNIITS-SMITH	-----	-----	-----	-----	VAN ROOYEN		
SAMPLE	8.00	1.101	1.931	KERSTEN	1.588	+0.3	MICKLEY	0.534	-8.2	0.1928	-----
				JOHANSEN	1.580	-0.3	MC GAI			0.1928	-----
				DE VRIES	1.597	-0.3	SCHMIT			0.1928	-----
				ADJ	DE VRIES	-----	GEMANT			0.1928	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.1928	-----
SAMPLE	8.00	1.203	1.693	KERSTEN	1.654	+0.9	MICKLEY	0.507	-14.4	0.564	-----
				JOHANSEN	1.512	-0.6	MC GAI			0.564	-----
				DE VRIES	1.492	-0.6	SCHMIT			0.564	-----
				ADJ	DE VRIES	-----	GEMANT			0.564	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.564	-----
SAMPLE	8.00	1.300	1.649	KERSTEN	1.585	+0.1	MICKLEY	0.562	-21.7	0.285	-----
				JOHANSEN	1.507	-0.5	MC GAI			0.285	-----
				DE VRIES	1.507	-0.5	SCHMIT			0.285	-----
				ADJ	DE VRIES	-----	GEMANT			0.285	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.285	-----
SAMPLE	8.00	1.401	1.123	KERSTEN	1.611	+0.3	MICKLEY	0.737	+28.0	0.491	-----
				JOHANSEN	1.621	-0.2	MC GAI			0.491	-----
				DE VRIES	1.621	-0.2	SCHMIT			0.491	-----
				ADJ	DE VRIES	-----	GEMANT			0.491	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.491	-----
SAMPLE	8.00	1.501	1.197	KERSTEN	1.741	+0.2	MICKLEY	0.815	-31.0	0.655	-----
				JOHANSEN	1.761	-0.5	MC GAI			0.655	-----
				DE VRIES	1.054	-0.1	SCHMIT			0.655	-----
				ADJ	DE VRIES	-----	GEMANT			0.655	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.655	-----
SAMPLE	8.00	1.601	1.345	KERSTEN	1.641	+0.0	MICKLEY	0.900	-35.0	0.498	-----
				JOHANSEN	1.693	-0.1	MC GAI			0.498	-----
				DE VRIES	1.117	-0.1	SCHMIT			0.498	-----
				ADJ	DE VRIES	-----	GEMANT			0.498	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.498	-----
SAMPLE	8.00	1.103	1.372	KERSTEN	1.620	+0.0	MICKLEY	0.724	-17.0	0.360	-----
				JOHANSEN	1.681	-0.2	MC GAI			0.360	-----
				DE VRIES	1.117	-0.2	SCHMIT			0.360	-----
				ADJ	DE VRIES	-----	GEMANT			0.360	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.360	-----
SAMPLE	18.00	1.201	1.181	KERSTEN	1.681	+0.6	MICKLEY	0.821	-24.1	0.298	-----
				JOHANSEN	1.693	-0.2	MC GAI			0.298	-----
				DE VRIES	1.117	-0.2	SCHMIT			0.298	-----
				ADJ	DE VRIES	-----	GEMANT			0.298	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.298	-----
SAMPLE	18.00	1.301	1.250	KERSTEN	1.672	+0.6	MICKLEY	0.929	-26.0	0.327	-----
				JOHANSEN	1.622	-0.6	MC GAI			0.327	-----
				DE VRIES	1.122	-0.6	SCHMIT			0.327	-----
				ADJ	DE VRIES	-----	GEMANT			0.327	-----
				RESISTOR	-----	-----	VAN ROOYEN	GEOM, MEAN		0.327	-----
				KUNIITS-SMITH	-----	-----				0.327	-----

Table B33 (cont'd).

Table B34.

PENNER LEED CLAY SATURATED FROZEN					
TYPE OF SOIL: CLAY NATURAL SATURATED FROZEN					
RHO = 1.750 TEMP = -25.000 KU = 8.430 CL = 1.000 CSOLIUS = 1.000	KA = 0.024		1/10 ALPHA = -----	KA = 0.024	
RHO = 1.750 KICE = 2.250 KU = 8.430 CL = 1.000 CSOLIUS2 = 1.000			1/10 ALPHA = -----	KA = 0.024	
SAMPLE MOISTURE CONTENT DRY K MEASURED METHOD K COMPUTED DEVIATION (%)			METHOD K COMPUTED DEVIATION (%)	PUR. SAT. UNC.	NC
SAMPLE 20.00 1.520 1.570 KERSTEN 2.023 6.5			HICKLEY 2.084 12.2	0.689 1.000	-0.135
JOHANSEN 1.920 1.920			MC GAN 2.084 12.2		
DE VRIES 1.754 1.754			SMITH 2.084 12.2		
ADJ DE VRIES -----			GEMANT 2.084 12.2		
RESISTOR 2.097 33.6			VAN RODDEN 2.084 12.2		
KUNIT-SMITH 2.151 37.0			GEOM. MEAN 2.084 12.2		
RHO = 1.750 TEMP = -25.000 KU = 8.430 CL = 1.000 CSOLIUS = 1.000	KA = 0.024		1/10 ALPHA = -----	KA = 0.024	
RHO = 1.750 KICE = 2.250 KU = 8.430 CL = 1.000 CSOLIUS2 = 1.000			1/10 ALPHA = -----	KA = 0.024	
SAMPLE MOISTURE CONTENT DRY K MEASURED METHOD K COMPUTED DEVIATION (%)			METHOD K COMPUTED DEVIATION (%)	PUR. SAT. UNC.	NC
SAMPLE 20.00 1.520 1.710 KERSTEN 2.023 4.5			HICKLEY 2.098 22.7	0.782 1.000	-0.190
JOHANSEN 1.920 1.920			MC GAN 2.098 22.7		
DE VRIES 1.840 1.840			SMITH 2.098 22.7		
ADJ DE VRIES -----			GEMANT 2.098 22.7		
RESISTOR 2.113 23.0			VAN RODDEN 2.098 22.7		
KUNIT-SMITH 2.168 26.0			GEOM. MEAN 2.098 22.7		
RHO = 1.750 TEMP = -10.000 KU = 8.430 CL = 1.000 CSOLIUS = 1.000	KA = 0.024		1/10 ALPHA = -----	KA = 0.024	
RHO = 1.750 KICE = 2.250 KU = 8.430 CL = 1.000 CSOLIUS2 = 1.000			1/10 ALPHA = -----	KA = 0.024	
SAMPLE MOISTURE CONTENT DRY K MEASURED METHOD K COMPUTED DEVIATION (%)			METHOD K COMPUTED DEVIATION (%)	PUR. SAT. UNC.	NC
SAMPLE 20.00 1.520 1.760 KERSTEN 2.023 43.3			HICKLEY 2.122 20.0	0.783 1.000	-0.098
JOHANSEN 1.920 1.920			MC GAN 2.122 20.0		
DE VRIES 1.840 1.840			SMITH 2.122 20.0		
ADJ DE VRIES -----			GEMANT 2.122 20.0		
RESISTOR 2.137 21.4			VAN RODDEN 2.122 20.0		
KUNIT-SMITH 2.194 24.7			GEOM. MEAN 2.122 20.0		
RHO = 1.750 TEMP = -22.000 KU = 8.430 CL = 1.000 CSOLIUS = 1.000	KA = 0.024		1/10 ALPHA = -----	KA = 0.024	
RHO = 1.750 KICE = 2.250 KU = 8.430 CL = 1.000 CSOLIUS2 = 1.000			1/10 ALPHA = -----	KA = 0.024	
SAMPLE MOISTURE CONTENT DRY K MEASURED METHOD K COMPUTED DEVIATION (%)			METHOD K COMPUTED DEVIATION (%)	PUR. SAT. UNC.	NC
SAMPLE 20.00 1.520 1.820 KERSTEN 2.023 38.0			HICKLEY 2.198 20.8	0.784 1.000	-0.075
JOHANSEN 1.920 1.920			MC GAN 2.198 20.8		
DE VRIES 1.880 1.880			SMITH 2.198 20.8		
ADJ DE VRIES -----			GEMANT 2.198 20.8		
RESISTOR 2.210 21.8			VAN RODDEN 2.198 20.8		
KUNIT-SMITH 2.280 25.3			GEOM. MEAN 2.198 20.8		

Table B35.

PENNER SUUBURG SALTY CLAY SATURATED FROZEN					
TYPE OF SOIL: CLAY NATURAL SATURATED FROZEN					
RHO = 1.750 TEMP = 0.000 KU = 2.000 CL = 1.000 CSOLIUS = 1.000	KA = 0.024		2/930 ALPHA = -----	KA = 0.024	
RHO = 0.500 KICE = 2.200 KU = 8.430 CL = 1.000 CSOLIUS2 = 1.000			2/930 ALPHA = -----	KA = 0.024	
SAMPLE MOISTURE CONTENT DRY K MEASURED METHOD K COMPUTED DEVIATION (%)			METHOD K COMPUTED DEVIATION (%)	PUR. SAT. UNC.	NC
SAMPLE 29.00 1.520 1.550 KERSTEN 2.218 4.5			HICKLEY 2.613 68.5	0.441 1.000	-0.448
JOHANSEN 2.000 2.000			MC GAN 2.613 68.5		
DE VRIES 1.490 1.490			SMITH 2.613 68.5		
ADJ DE VRIES -----			GEMANT 2.613 68.5		
RESISTOR 2.029 69.6			VAN RODDEN 2.613 68.5		
KUNIT-SMITH 2.559 65.1			GEOM. MEAN 2.613 68.5		
RHO = 1.750 TEMP = 22.000 KU = 8.430 CL = 1.000 CSOLIUS = 1.000	KA = 0.024		2/940 ALPHA = -----	KA = 0.024	
RHO = 0.500 KICE = 2.270 KU = 8.430 CL = 1.000 CSOLIUS2 = 1.000			2/940 ALPHA = -----	KA = 0.024	
SAMPLE MOISTURE CONTENT DRY K MEASURED METHOD K COMPUTED DEVIATION (%)			METHOD K COMPUTED DEVIATION (%)	PUR. SAT. UNC.	NC
SAMPLE 29.00 1.520 2.130 KERSTEN 2.218 4.5			HICKLEY 2.686 22.3	0.487 1.000	-0.155
JOHANSEN 2.000 2.000			MC GAN 2.686 22.3		
DE VRIES 1.436 1.436			SMITH 2.686 22.3		
ADJ DE VRIES -----			GEMANT 2.686 22.3		
RESISTOR 2.022 23.1			VAN RODDEN 2.686 22.3		
KUNIT-SMITH 2.548 19.8			GEOM. MEAN 2.686 22.3		
RHO = 1.750 TEMP = 25.000 KU = 8.430 CL = 1.000 CSOLIUS = 1.000	KA = 0.024		2/950 ALPHA = -----	KA = 0.024	
RHO = 0.500 KICE = 2.250 KU = 8.430 CL = 1.000 CSOLIUS2 = 1.000			2/950 ALPHA = -----	KA = 0.024	
SAMPLE MOISTURE CONTENT DRY K MEASURED METHOD K COMPUTED DEVIATION (%)			METHOD K COMPUTED DEVIATION (%)	PUR. SAT. UNC.	NC
SAMPLE 29.00 1.520 2.220 KERSTEN 2.218 4.5			HICKLEY 2.617 17.0	0.472 1.000	-0.090
JOHANSEN 2.000 2.000			MC GAN 2.617 17.0		
DE VRIES 2.136 2.136			SMITH 2.617 17.0		
ADJ DE VRIES -----			GEMANT 2.617 17.0		
RESISTOR 2.033 18.6			VAN RODDEN 2.617 17.0		
KUNIT-SMITH 2.559 18.3			GEOM. MEAN 2.617 17.0		
RHO = 1.750 TEMP = -10.000 KU = 8.430 CL = 1.000 CSOLIUS = 1.000	KA = 0.024		2/970 ALPHA = -----	KA = 0.024	
RHO = 0.500 KICE = 2.320 KU = 8.430 CL = 1.000 CSOLIUS2 = 1.000			2/970 ALPHA = -----	KA = 0.024	
SAMPLE MOISTURE CONTENT DRY K MEASURED METHOD K COMPUTED DEVIATION (%)			METHOD K COMPUTED DEVIATION (%)	PUR. SAT. UNC.	NC
SAMPLE 29.00 1.520 2.340 KERSTEN 2.218 4.5			HICKLEY 2.643 12.0	0.472 1.000	-0.098
JOHANSEN 2.000 2.000			MC GAN 2.643 12.0		
DE VRIES 2.156 2.156			SMITH 2.643 12.0		
ADJ DE VRIES -----			GEMANT 2.643 12.0		
RESISTOR 2.058 18.6			VAN RODDEN 2.643 12.0		
KUNIT-SMITH 2.580 18.3			GEOM. MEAN 2.643 12.0		

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Table B35 (cont'd).

RHO =	2,780	TEMP =	-24,000	CL =	0,000	CSOLIDUS =	3,080	ALPHA =	-----	KA =	0,024
KH =	0,960	KICE =	2,280	KD =	0,430	KU =	2,080	CSOLIDUS2 =	3,080	ALPHA2 =	-----
SAMPLE	20.00	DNY	1.020	K MEASURED	2.380	METHOD K COMPUTED	2.118	DEVIATION	(PERCENT)	MICKLEY	2.099
						JOHANSEN	2.483		-6.0	MC GAN	-----
						DE VRIES	2.424		-1.0	SMITH	-----
						ADJ DE VRIES	-----			GEMANT	-----
						MESISTOR	2.112		-13.9	VAN ROOYEN	-----
						KUNIT-SMITH	2.050		-17.2	GEOM. MEAN	2.099
											-13.4

Table B36.

PENNER ET AL SOIL 7 SATURATED FROZEN

TYPE OF SOIL:	CLAY NATURAL SATURATED	FROZEN
RHO = 2,780 TEMP = -24,000 CL = 0,000 CSOLIDUS = 3,080 ALPHA = ----- KA = 0,024		
KH = 0,960 KICE = 2,280 KD = 0,430 KU = 2,080 CSOLIDUS2 = 3,080		
SAMPLE MOISTURE CONTENT DNY K MEASURED		
SAMPLE 13.30 1.934 2.581		
KERSTEN 2.186 -16.3 MICKLEY 2.787 8.0 PUR. SAT. UNC		
JOHANSEN 2.422 -6.2 MC GAN ----- 0.272 1.000 -----		
DE VRIES 2.449 -5.1 SMITH ----- 0.180 0.180 -----		
ADJ DE VRIES ----- ----- GEMANT ----- -----		
MESISTOR 2.091 -8.5 VAN ROOYEN ----- -----		
KUNIT-SMITH 2.071 7.6 GEOM. MEAN 2.788 8.0 -----		
SAMPLE 13.30 + 1.917 2.581		
KERSTEN 2.097 -18.6 MICKLEY 2.782 7.8 PUR. SAT. UNC		
JOHANSEN 2.783 7.8 MC GAN ----- 0.278 1.000 -----		
DE VRIES 2.789 8.1 SMITH ----- 0.278 0.278 -----		
ADJ DE VRIES ----- ----- GEMANT ----- -----		
MESISTOR 2.790 8.3 VAN ROOYEN ----- -----		
KUNIT-SMITH 2.669 7.3 GEOM. MEAN 2.783 7.8 -----		

Table B37.

SLUSARCHUK-WATSON UNDISTURBED PERMAFROST

TYPE OF SOIL:	CLAY NATURAL SATURATED	FROZEN
RHO = 2,780 TEMP = -24,000 CL = 0,000 CSOLIDUS = 3,080 ALPHA = ----- KA = 0,024		
KH = 0,970 KICE = 2,280 KD = 0,430 KU = 2,080 CSOLIDUS2 = 3,080		
SAMPLE MOISTURE CONTENT DNY K MEASURED		
SAMPLE 38.98 1.317 2.120		
KERSTEN 2.149 1.4 MICKLEY 2.154 1.0 PUR. SAT. UNC		
JOHANSEN 2.154 1.0 MC GAN ----- 0.530 1.000 -----		
DE VRIES 2.155 1.0 SMITH ----- 0.000 0.000 -----		
ADJ DE VRIES ----- ----- GEMANT ----- -----		
MESISTOR 2.158 1.8 VAN ROOYEN ----- -----		
KUNIT-SMITH 2.188 3.2 GEOM. MEAN 2.154 1.0 -----		
SAMPLE 38.98 1.261 2.234		
KERSTEN 2.164 -2.4 MICKLEY 2.158 -3.2 PUR. SAT. UNC		
JOHANSEN 2.188 -3.2 MC GAN ----- 0.543 1.000 -----		
DE VRIES 2.160 -3.0 SMITH ----- 0.000 0.000 -----		
ADJ DE VRIES ----- ----- GEMANT ----- -----		
MESISTOR 2.162 -3.1 VAN ROOYEN ----- -----		
KUNIT-SMITH 2.193 -1.7 GEOM. MEAN 2.158 -3.2 -----		

Table B38.

USSR BUILDING CODE (1960) CLAY SOILS SATURATED FROZEN

TYPE OF SOIL:	CLAY NATURAL SATURATED	FROZEN
RHO = 2,780 TEMP = -24,000 CL = 0,000 CSOLIDUS = 3,080 ALPHA = ----- KA = 0,024		
KH = 0,960 KICE = 2,280 KD = 0,430 KU = 2,080 CSOLIDUS2 = 3,080		
SAMPLE MOISTURE CONTENT DNY K MEASURED		
SAMPLE 48.00 1.449 2.790		
KERSTEN 2.012 -27.0 MICKLEY 2.092 -25.0 PUR. SAT. UNC		
JOHANSEN 2.092 -25.0 MC GAN ----- 0.343 1.000 -----		
DE VRIES 2.093 -25.0 SMITH ----- 0.000 0.000 -----		
ADJ DE VRIES ----- ----- GEMANT ----- -----		
MESISTOR 2.099 -24.9 VAN ROOYEN ----- -----		
KUNIT-SMITH 2.112 -24.3 GEOM. MEAN 2.092 -25.0 -----		
SAMPLE 48.00 + 1.242 2.380		
KERSTEN 2.119 -10.2 MICKLEY 2.147 -9.0 PUR. SAT. UNC		
JOHANSEN 2.144 -9.8 MC GAN ----- 0.361 1.000 -----		
DE VRIES 2.140 -9.8 SMITH ----- 0.000 0.000 -----		
ADJ DE VRIES ----- ----- GEMANT ----- -----		
MESISTOR 2.151 -9.9 VAN ROOYEN ----- -----		
KUNIT-SMITH 2.180 -7.6 GEOM. MEAN 2.147 -9.0 -----		

Table B39.

JOHANSEN CRUSHED ROCK DRY											
TYPE - S 112 COARSE NATURAL DRY			RHO = 1.672 TEMP = 20.21°C KU = 2.000 KU = 2.000 CSOLIDS = 0.000 CSOLIDS = 0.000			ALPHA = 0.065 KA = 0.026					
SAMPLE	WATER CONTENT	DRY DENSITY	K MEASURED	METHOD & COMPUTED	DEVIATION (PERCENT)	METHOD & COMPUTED	DEVIATION (PERCENT)	POR. SAT.	NC WUE		
SAMPLE 1	0.100	1.683	0.354	KERSTEN	---	---	MICKLEY	1.659	36.9	0.319	0.000
			DE VRIES	0.314	-7.8	MC GAW	1.656	-3.4	0.000	0.000	
			ADJ DE VRIES	0.384	27.0	SMITH	0.387	-3.4			
			RESISTON	0.386	27.4	GEMANT	0.386	-3.4			
			KUNITS-SMITH	0.392	29.3	VAN BOOYEN	0.346	14.3			
						GEOM. MEAN	0.366	-3.4			
SAMPLE 2	0.134	1.682	0.354	KERSTEN	---	---	MICKLEY	1.659	61.7	0.270	0.000
			DE VRIES	0.376	23.9	MC GAW	2.737	750.1	0.000	0.000	
			ADJ DE VRIES	0.260	-6.9	SMITH	0.304	-6.9			
			RESISTON	0.384	24.9	GEMANT	0.384	-3.4			
			KUNITS-SMITH	0.384	24.9	VAN BOOYEN	0.431	33.6			
						GEOM. MEAN	0.391	-3.4			
SAMPLE 3	0.100	1.683	0.354	KERSTEN	---	---	MICKLEY	1.707	34.6	0.237	0.000
			DE VRIES	0.320	-7.8	MC GAW	1.698	62.2	0.000	0.000	
			ADJ DE VRIES	0.344	-6.9	SMITH	0.358	-2.9			
			RESISTON	0.360	17.2	GEMANT	0.360	-2.7			
			KUNITS-SMITH	0.360	109.0	VAN BOOYEN	0.491	22.7			
						GEOM. MEAN	0.386	-2.7			

Table B40.

JOHANSEN CAL-D SAT DRY											
TYPE - S 112 SATURATED NATURAL DRY			RHO = 1.672 TEMP = 20.21°C KU = 2.000 KU = 2.000 CSOLIDS = 0.000 CSOLIDS = 0.000			ALPHA = 0.073 KA = 0.026					
SAMPLE	WATER CONTENT	DRY DENSITY	K MEASURED	METHOD & COMPUTED	DEVIATION (PERCENT)	METHOD & COMPUTED	DEVIATION (PERCENT)	POR. SAT.	NC WUE		
SAMPLE 1	0.100	1.683	0.354	KERSTEN	---	---	MICKLEY	1.553	358.3	0.320	0.000
			JOHANSEN	0.320	-2.8	MC GAW	2.67	668.0	0.000	0.000	
			DE VRIES	0.304	-6.0	SMITH	0.250	-23.1			
			ADJ DE VRIES	0.350	12.7	GEMANT	0.350	-3.4			
			RESISTON	0.350	12.7	VAN BOOYEN	0.386	13.3			
			KUNITS-SMITH	0.395	16.3	GEOM. MEAN	0.356	-2.7			

Table B41.

JOHANSEN CRUSHER ROCK DRY											
TYPE - S 112 COARSE CRUSHED DRY			RHO = 1.672 TEMP = 20.21°C KU = 2.000 KU = 2.000 CSOLIDS = 0.000 CSOLIDS = 0.000			ALPHA = 0.065 KA = 0.026					
SAMPLE	WATER CONTENT	DRY DENSITY	K MEASURED	METHOD & COMPUTED	DEVIATION (PERCENT)	METHOD & COMPUTED	DEVIATION (PERCENT)	POR. SAT.	NC WUE		
SAMPLE 1	0.100	1.683	0.369	KERSTEN	---	---	MICKLEY	0.904	238.0	0.395	0.000
			JOHANSEN	0.223	-7.3	MC GAW	1.699	331.5	0.000	0.000	
			DE VRIES	0.307	20.7	SMITH	0.250	-23.1			
			ADJ DE VRIES	0.304	-6.0	GEMANT	0.304	-16.2			
			RESISTON	0.350	16.3	VAN BOOYEN	0.331	-7.3			
			KUNITS-SMITH	0.393	32.0	GEOM. MEAN	0.313	-13.1			
SAMPLE 2	0.134	1.682	0.369	KERSTEN	---	---	MICKLEY	1.043	284.7	0.392	0.000
			DE VRIES	0.306	-10.7	MC GAW	1.884	449.5	0.000	0.000	
			ADJ DE VRIES	0.296	-32.7	SMITH	0.234	-27.8			
			RESISTON	0.303	-18.0	GEMANT	0.293	-33.1			
			KUNITS-SMITH	0.317	-30.4	VAN BOOYEN	0.203	-50.0			

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